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GEOMETRIC STRUCTURE OF THE EARTH'S GRAVITATIONAL FIELD  
AS DERIVED FROM ARTIFICIAL SATELLITES

by

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**Abstract.**--The structure of the Earth's gravitational potential and its gradient is studied at sea level and in outer space (1,000 km, 10,000 km, and 100,000 km elevation above sea level), with particular consideration given to the surfaces of constant potential and of constant gravitation. Their shape, curvature, etc., are computed and the results presented in tables and maps using the zonal part of the gravitational field as a reference frame. In a concluding section the geometry of the orthogonal trajectories of the equipotential surfaces (direction of the gradient field) is briefly touched. Among others, the curvature and torsion of the verticals are derived and the trajectories are integrated up to 10 km height for the gravity field and 100,000 km for the gravitational field. All numerical computations are based on Izsak's and Kozai's latest harmonic coefficients of the gravitational potential (Izsak, 1965; Kozai, 1964).

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The gravitational potential of outer space can be described by the function

$$U = \frac{GM}{r} \left\{ 1 + \sum_{n=2}^{\infty} \sum_{m=0}^n \left( \frac{a}{r} \right)^n (C_{nm} \cos m\lambda + S_{nm} \sin m\lambda) P_{nm}(\sin \varphi) \right\}, \quad (1)$$

with

$a$  = equatorial radius,

$GM$  = product of the gravitational constant and the mass of the Earth

$r$  = geocentric radius, i.e., the distance from the gravity center to a point in free space,

$\varphi$  = geocentric latitude, i.e., the complement of the angle that is included by  $r$  and the rotation axis,

$\lambda$  = geocentric longitude, i.e., the angle between a meridian plane through  $r$  and the meridian plane through Greenwich,

$C_{nm}, S_{nm}$  = harmonic coefficients,

$P_{nm}(\sin \varphi)$  = Legendre's associated function.

Equation (1) is valid in the case of

$$\operatorname{div} \operatorname{grad} U = 0, \quad (2)$$

that is, in empty space where Laplace's equation is satisfied.

In our analysis we make two specifications:

- 1) we disregard the mass of the atmosphere, and
- 2) we extend the above function down to sea level, ignoring thereby the effect of the interfering topography. This gravitational potential is described together with its gradient field, in four different sections of outer space:

at sea level, at 1,000 km elevation, at 10,000 km elevation, and at 100,000 km elevation. At sea level we shall also consider the case where the centrifugal potential Z is included:

$$Z = \frac{\omega^2 r^2}{2} \cos^2 \varphi , \quad (3)$$

where  $\omega$  = angular velocity of the Earth.

The numerical computations are performed with a set of harmonic coefficients given in the Appendix. Although these values will be subject to continuous improvement, they already give a rather close picture of the general structure of the Earth's gravitational field. In the following sections we comment only briefly on the analysis used. It is, however, not intended to give a collection of formulas, nor should the subject be treated from a theoretical point of view.

## I. POTENTIAL FIELD - EQUIPOTENTIAL SURFACES

The geopotential in free space is a function of the radius vector  $\vec{r}$  and the harmonic coefficients  $C_{nm}$  and  $S_{nm}$ :

$$V = U + Z = V(r, \varphi, \lambda, C_{nm}, S_{nm}; \omega) \quad . \quad (4)$$

By neglecting the influence of the nonzonal terms, we get a new potential function:

$$\begin{aligned} \bar{V} &= \bar{V}(\bar{r}, \varphi, C_{no}; \omega) \quad , \\ &= \frac{GM}{\bar{r}} \left\{ 1 + \sum_{n=2}^{\infty} \left( \frac{a}{\bar{r}} \right)^n C_{no} P_{no}(\sin \varphi) \right\} + \frac{\omega^2 \bar{r}^2}{2} \cos^2 \varphi \quad , \end{aligned} \quad (5)$$

which is rotational symmetric and approximates the potential  $V$  in the mean. To relate both equations (4) and (5), we put

$$V = \bar{V}$$

and determine  $\bar{V}$  in such a way that at sea level

$$\bar{r} = 6,378,165 \text{ m for } \varphi = 0, \text{ and } \omega \neq 0 \text{ or } \omega = 0, \quad ^3$$

and similarly in outer space

$$\left. \begin{array}{l} \bar{r} = 6,378,165 \text{ m} + 1,000' \text{ km} \\ \bar{r} = 6,378,165 \text{ m} + 10,000 \text{ km} \\ \bar{r} = 6,378,165 \text{ m} + 100,000 \text{ km} \end{array} \right\} \text{ for } \begin{array}{l} \varphi = 0 \\ \omega = 0 \end{array} .$$

<sup>3</sup> $\omega \neq 0$  simply means that the centrifugal potential is taken into consideration.

To get the geometric structure of an equipotential surface  $V = \text{const}$ , we first compute the relatively simple (rotational symmetric) structure of  $\bar{V} = \text{const}$  and add a small correction term, leading to the geometrical structure of  $V = \text{const}$ . The advantage of this stepwise procedure is also a practical one: the geometric results of  $\bar{V} = \text{const}$  can be presented in tables, while the small corrections are easily plotted as equilevel curves in maps.

### 1. Shape of the equipotential surfaces

To determine the geocentric radii of an equipotential surface  $V = \text{const}$ , we split  $r$  into a spheroidal part  $\bar{r}$  and a correction term  $\Delta r$  (see Figure 1):

$$r(\varphi, \lambda) = \bar{r}(\varphi) + \Delta r(\varphi, \lambda) . \quad (6)$$

While  $\bar{r}$  as a function of  $\varphi$  only is computed from equation (5), we obtain, by introducing the latter into equation (4),

$$V = V(\bar{r} + \Delta r, \varphi, \lambda; C_{nm}, S_{nm}; \omega) , \quad (7)$$

the "undulations"  $\Delta r$  in the geocentric latitude  $\varphi$  and longitude  $\lambda$ .

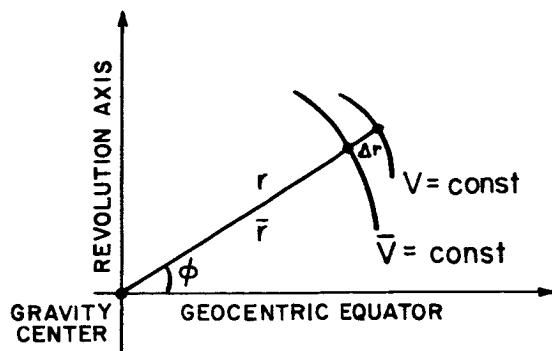


Figure 1.

Example: The geocentric radius  $r$  in the geocentric latitude  $\varphi = 20^\circ$  and longitude  $\lambda = 40^\circ$  is for the equipotential surface  $V(\omega = 0) = \text{const}$  at 1,000 km elevation above sea level:

$$\left. \begin{aligned} r_\varphi &= 20^\circ = \bar{r}_\varphi = 20^\circ + \Delta r_\varphi = 20^\circ \\ \lambda &= 40^\circ \qquad \qquad \qquad \lambda = 40^\circ \\ &\qquad \qquad \qquad = 7,377,109 + 4 \\ &\qquad \qquad \qquad = 7,377,113 \text{ m} \end{aligned} \right\} \text{see 1a), b) .}$$

- a) Spheroidal<sup>4</sup> (zonal) part  $\bar{r}$  of the geocentric radius — shape of the equipotential surface  $\bar{V} = \text{const}$ .

Table 1.

$\varphi^\circ$	Sea level		Elevation above sea level ( $\omega = 0$ )		
	$\omega \neq 0$	$\omega = 0$	1,000 km	10,000 km	100,000 km
90	6 356 792.84	6 367 812.46	7 369 216.03	16 374 133.15	106 377 544.02
80	6 357 432.57	6 368 122.11	7 369 484.57	16 374 254.56	106 377 562.74
70	6 359 277.55	6 369 015.89	7 370 258.62	16 374 604.17	106 377 616.65
60	6 362 110.55	6 370 389.36	7 371 446.05	16 375 139.93	106 377 699.24
50	6 365 590.83	6 372 075.63	7 372 903.81	16 375 797.32	106 377 800.56
40	6 369 299.88	6 373 871.18	7 374 456.29	16 376 497.19	106 377 908.38
30	6 372 791.63	6 375 560.68	7 375 916.75	16 377 155.19	106 378 009.71
20	6 375 643.11	6 376 939.98	7 377 109.46	16 377 691.97	106 378 092.33
10	6 377 511.56	6 377 846.07	7 377 890.96	16 378 042.72	106 378 146.26
0	6 378 165.00	6 378 165.00	7 378 165.00	16 378 165.00	106 378 165.00
-10	6 377 518.83	6 377 853.31	7 377 896.75	16 378 043.92	106 378 146.29
-20	6 375 657.39	6 376 954.22	7 377 119.49	16 377 693.99	106 378 092.38
-30	6 372 804.89	6 375 573.89	7 375 926.91	16 377 157.34	106 378 009.77
-40	6 369 308.04	6 373 879.32	7 374 463.20	16 376 498.68	106 377 908.42
-50	6 365 593.98	6 372 078.76	7 372 905.38	16 375 797.47	106 377 800.56
-60	6 362 102.61	6 370 381.43	7 371 439.32	16 375 138.34	106 377 699.20
-70	6 359 254.55	6 368 992.98	7 370 241.97	16 374 600.88	106 377 616.57
-80	6 357 398.44	6 368 088.12	7 369 460.19	16 374 250.01	106 377 562.64
-90	6 356 754.75	6 367 774.51	7 369 188.79	16 374 128.14	106 377 543.90

<sup>4</sup>The word spheroid is used in the following for "zonal part" and vice versa.

The geocentric radii at the poles are always smaller than the corresponding equatorial radius. At sea level the difference is about 21.4 km ( $\omega \neq 0$ ) and about 10.4 km ( $\omega = 0$ ), respectively, and decreases at 100,000 km elevation to 0.62 km ( $\omega = 0$ ). For infinity ( $\omega = 0$ ) the difference is obviously zero.

The phrase "sea level" has, of course, only physical meaning for  $\omega \neq 0$ ; in the case  $\omega = 0$ , only the equatorial radius  $\bar{r}(\varphi = 0)$  is supposed to be at sea level (see also page 4).

b) Corrections (undulations),  $\Delta r$

See Contour Map 1, page 70. The undulations are referred to the spheroidal part of the equipotential surface (equation (5)). As shown in Table 2, we have at sea level heights up to 60 m and depths down to -70 m. In outer space these undulations are smoothed depending on the height above sea level. At 100,000 km elevation only the  $C_{22}$  and  $S_{22}$  terms show up clearly and give the gravitational field some similarity with that of a 3-axial ellipsoid. Beyond 100,000 km the potential field assumes more and more a spherical shape according to equation (4). At ground level, only the case  $\Delta r$  ( $\omega \neq 0$ ) has been plotted, because  $\Delta r$  ( $\omega = 0$ ) is equal to  $\Delta r$  ( $\omega \neq 0$ ) within some decimeters.

Table 2.

Elevation above sea level (km)	Range of $\Delta r$ (m)	
0	60	-70
1,000	47	-51
10,000	16	-14
100,000	1.8	-1.8

c) Difference of the geocentric radii of the Northern and Southern Hemispheres of the spheroid

Table 3.

$\varphi^{\circ}$	$\bar{r}(\varphi) - \bar{r}(-\varphi)$ (m)	Sea level		Elevation above sea level ( $\omega = 0$ )		
		$\omega \neq 0$	$\omega = 0$	1,000 km	10,000 km	100,000 km
90	+38.09	+37.95	+27.24	+5.01	+0.117	
80	+34.13	+33.99	+24.38	+4.55	+0.106	
70	+23.00	+22.91	+16.65	+3.29	+0.078	
60	+ 7.94	+ 7.93	+ 6.73	+1.59	+0.038	
50	- 3.15	- 3.13	- 1.57	-0.15	-0.003	
40	- 8.16	- 8.14	- 6.91	-1.49	-0.035	
30	-13.26	-13.21	-10.16	-2.15	-0.051	
20	-14.28	-14.24	-10.03	-2.02	-0.048	
10	- 7.27	- 7.24	- 5.79	-1.20	-0.029	
0	0.00	0.00	0.00	0.00	0.000	

The differences  $\bar{r}(\varphi) - \bar{r}(-\varphi)$  shown in Table 3 point out an asymmetric spheroid whose northern part is longer and tighter while its southern part is shorter but broader.<sup>5</sup>

d) Difference of a spheroid and an ellipsoid with equal axes

The deviation of an equipotential spheroid from an ellipsoid with coincident axes is obtained by comparison of their geocentric radii  $\bar{r}$  and  $E$  along a meridian plane (see Table 4). At the Northern Hemisphere the equipotential spheroid is completely surrounded by the corresponding ellipsoid, while for the Southern Hemisphere the ellipsoid stays within the spheroid. The largest deviation - 19.2 m - is found at sea level ( $\omega \neq 0$ ) at about  $-50^{\circ}$  latitude.

<sup>5</sup>Often called pear shaped; this is, however, not quite correct because these spheroids have no negative gaussian curvature (see page 30).

Table 4.

$\varphi^\circ$	$\bar{r} - E$ (m)	Sea level		Elevation above sea level ( $\omega = 0$ )		
		$\omega \neq 0$	$\omega = 0$	1,000 km	10,000 km	100,000 km
90	0.00	0.00	0.00	0.00	0.0000	
80	- 1.59	- 1.77	- 0.83	- 0.13	- 0.0034	
70	- 4.27	- 4.97	- 2.56	- 0.46	- 0.0125	
60	- 5.18	- 6.51	- 4.17	- 0.91	- 0.0243	
50	- 6.40	- 8.13	- 5.77	- 1.34	- 0.0351	
40	- 8.56	- 10.28	- 7.26	- 1.60	- 0.0410	
30	-10.13	-11.45	- 7.95	- 1.58	- 0.0396	
20	-10.69	-11.40	- 7.02	- 1.25	- 0.0306	
10	- 5.85	- 6.03	- 3.72	- 0.67	- 0.0161	
0	0.00	0.00	0.00	0.00	0.0000	
-10	+ 2.58	+ 2.37	+ 2.90	+ 0.68	+ 0.0163	
-20	+ 8.09	+ 7.30	+ 6.20	+ 1.35	+ 0.0312	
-30	+12.71	+11.28	+ 9.05	+ 1.82	+ 0.0407	
-40	+15.43	+13.58	+10.94	+ 1.96	+ 0.0423	
-50	+19.20	+17.31	+11.81	+ 1.75	+ 0.0364	
-60	+15.51	+14.06	+ 9.56	+ 1.26	+ 0.0254	
-70	+ 6.41	+ 5.64	+ 4.87	+ 0.67	+ 0.0130	
-80	+ 1.25	+ 1.03	+ 1.22	+ 0.19	+ 0.0036	
-90	0.00	0.00	0.00	0.00	0.0000	

- e) Difference of the spheroid ( $\omega \neq 0$ ) at sea level and an ellipsoid with the flattening  $f = 1/298.25$

The undulations of the "satellite geoid" referred to the ellipsoid with the above flattening  $f = 1/298.25$  are obtained by adding the values  $\bar{r} - E$  of the Table 5 to the undulations at sea level (page 70).

Example: At  $\varphi = 20^\circ$  and  $\lambda = 40^\circ$  we obtain the undulation in question:

$$\begin{array}{ll} \Delta r & 13 \text{ m} \\ \bar{r} - E & \underline{-9 \text{ m}} \\ \text{undulation} & 4 \text{ m} \end{array}$$

Table 5.

$\psi^\circ$	$\bar{r} - E$ (m)
90	+13.13
80	+11.16
70	+ 7.35
60	+ 4.70
50	+ 1.34
40	- 3.10
30	- 6.81
20	- 9.14
10	- 5.44
0	0.00
-10	+ 1.83
-20	+ 5.15
-30	+ 6.43
-40	+ 5.06
-50	+ 4.50
-60	- 3.24
-70	-15.65
-80	-22.96
-90	-24.95

The spheroid and the ellipsoid were supposed to have the same equatorial radius  $\bar{r} = 6,378,165$  m. However, if we refer the satellite geoid to an ellipsoid with the same potential and the same zonal harmonic coefficient of the second degree  $C_{20} = -J_2$ , we have to add to the above values a constant term of about 0.75 m. It should be noted that the flattening used is a rounded value of Kozai's (1964, p. 12).

f) Difference of the spheroids ( $\omega = 0$ ) and ( $\omega \neq 0$ ) at sea level

The opposite results are obtained from the geocentric radii of section a) at sea level along a meridian. The spheroid ( $\omega \neq 0$ ) is completely surrounded by the corresponding spheroid ( $\omega = 0$ ) so that the difference of the geocentric radii increases from 0 at the equator according to definition to 11.02 km at the poles.

Table 6.

$\phi^\circ$	$\bar{r} (\omega = 0)$ - $\bar{r} (\omega \neq 0)$ (m)
90	11,019.61
80	10,689.54
70	9,738.34
60	8,278.81
50	6,484.80
40	4,571.30
30	2,769.05
20	1,296.87
10	334.51
0	0.00
-10	334.48
-20	1,296.83
-30	2,769.00
-40	4,571.28
-50	6,484.78
-60	8,278.82
-70	9,738.43
-80	10,689.68
-90	11,019.76

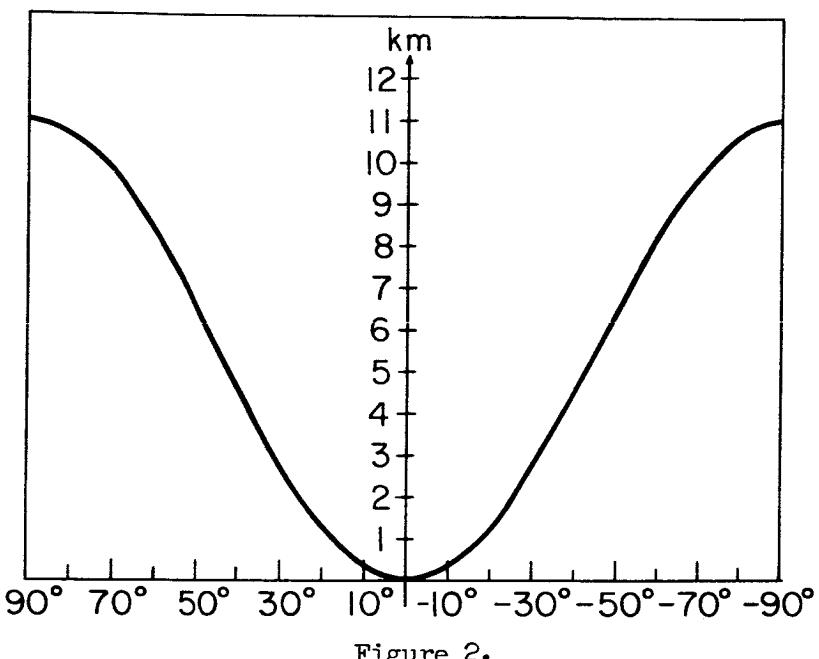


Figure 2.

g) Mean flattening of the spheroids ( $\omega = 0$ ) as a function of elevation above sea level

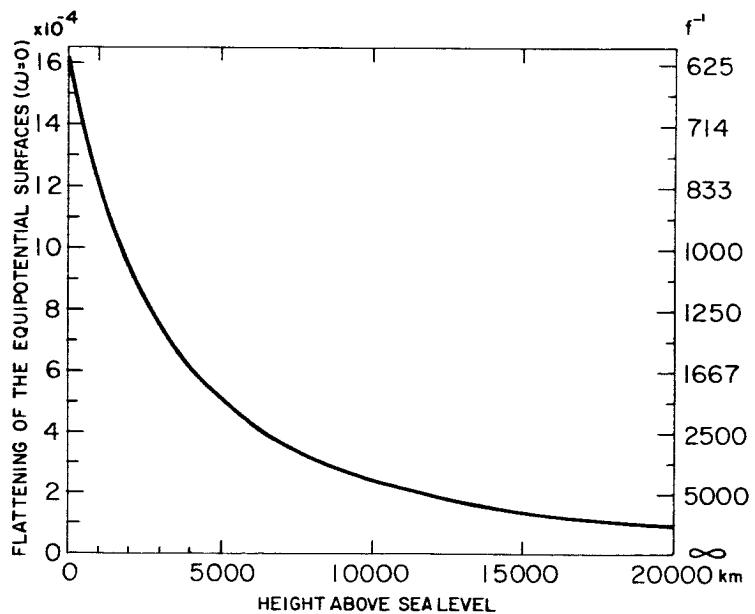


Figure 3.

Definition of the mean flattening:

$$f = 1 - \frac{\bar{r}_{\text{north}} + \bar{r}_{\text{south}}}{2 \bar{r}_{\text{equator}}} . \quad (8)$$

As we can see from Figure 3, between sea level and 3,000 km elevation the flattening of the equipotential spheroids ( $\omega = 0$ ) decreases very strongly and tends afterward asymptotically to zero at infinite height.

- h) Distance of the actual equator of the spheroids from the gravity center

Because of the odd zonal coefficients  $C_{n_0}$  ( $n = 3, 5, \dots$ ), the actual equator (surface normal is orthogonal to revolution axis) of the spheroid  $\bar{V} = \text{const}$  does not coincide with the geocentric equator (see Figure 4).

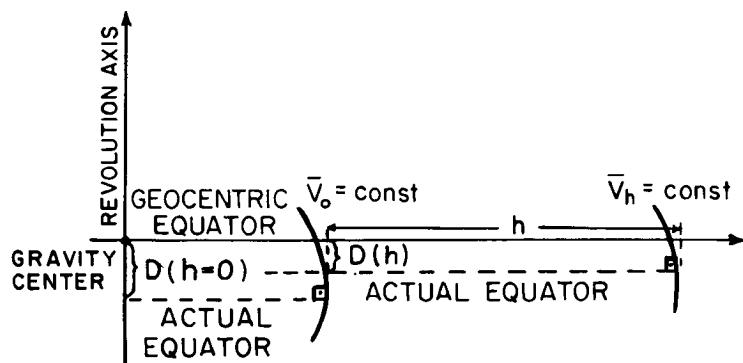


Figure 4.

The polar coordinates  $\bar{r}, \varphi$  of the actual equator derive from

$$\bar{r} \frac{\partial \bar{V}}{\partial \varphi} \sin \varphi + \frac{\partial \bar{V}}{\partial \varphi} \cos \varphi = 0,$$

and the distance D:

$$D \approx a \left[ -\frac{3}{2} c_{30} \left(1 + \frac{h}{a}\right)^{-2} + \frac{15}{8} c_{50} \left(1 + \frac{h}{a}\right)^{-4} - \frac{35}{16} c_{70} \left(1 + \frac{h}{a}\right)^{-6} \right. \\ \left. + \frac{315}{128} c_{90} \left(1 + \frac{h}{a}\right)^{-8} - \frac{693}{256} c_{110} \left(1 + \frac{h}{a}\right)^{-10} + \frac{3003}{1024} c_{130} \left(1 + \frac{h}{a}\right)^{-12} - \dots \right], \quad (9)$$

where  $h$  = average elevation of the spheroid above sea level. Figure 5 shows that the actual equator moves at first slightly southward and then reapproaches asymptotically the geocentric equator as the flattening decreases.

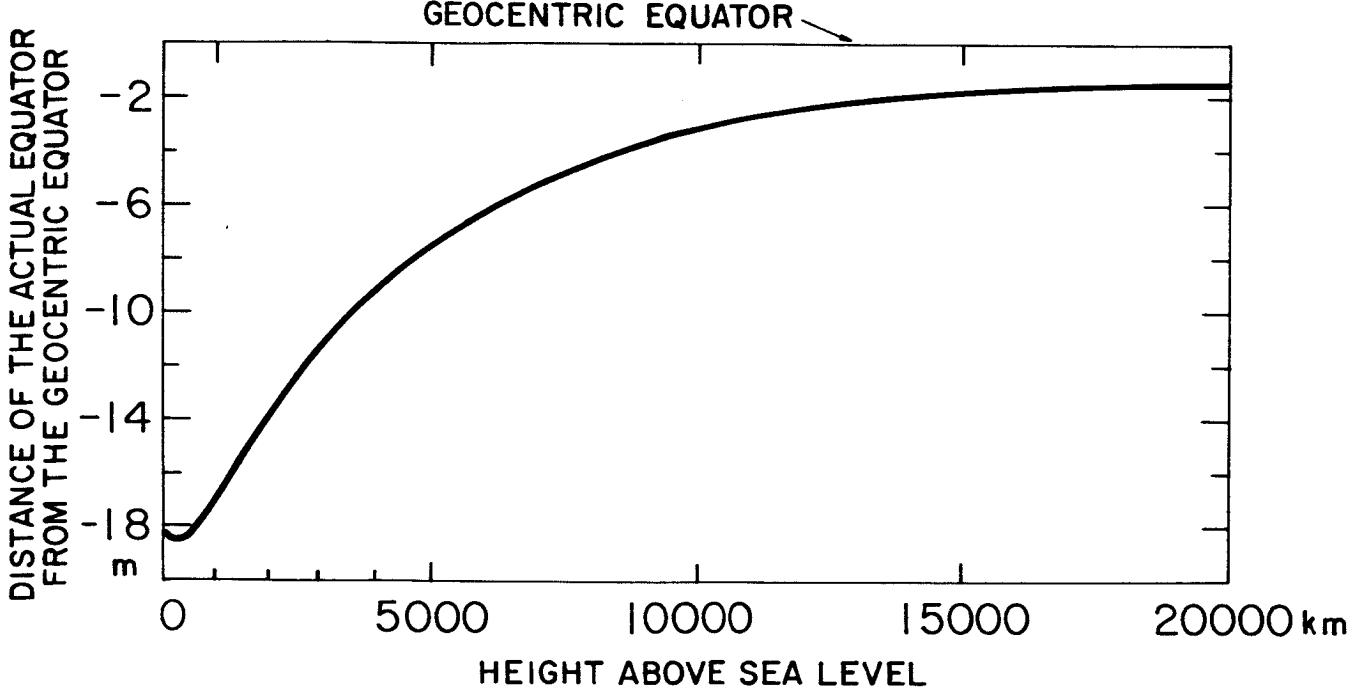


Figure 5.

2. Gravitation<sup>6</sup> (gravity) along an equipotential surface

By partial differentiation of equation (4) with respect to the rectangular coordinates  $x^i$  we obtain

$$\text{grad } V = \frac{\partial V}{\partial x^1} \vec{e}^1 + \frac{\partial V}{\partial x^2} \vec{e}^2 + \frac{\partial V}{\partial x^3} \vec{e}^3 \quad (10)$$

if

$$\left. \begin{array}{l} x^1 = r \cos \varphi \cos \lambda \\ x^2 = r \cos \varphi \sin \lambda \\ x^3 = r \sin \varphi \end{array} \right\} \quad \text{see page 109.} \quad (11)$$

and similarly with equation (5)

$$\text{grad } \bar{V} = \frac{\partial \bar{V}}{\partial \bar{x}^1} \vec{e}^1 + \frac{\partial \bar{V}}{\partial \bar{x}^2} \vec{e}^2 + \frac{\partial \bar{V}}{\partial \bar{x}^3} \vec{e}^3 . \quad (12)$$

Introducing  $\bar{r}$  and  $r = \bar{r} + \Delta r$  of the previous section into equations (12) and (10), we derive the variation of the gravitation along the spheroid:

$$\begin{aligned} \bar{g} &= |\text{grad } \bar{V}| = \bar{g}(\bar{r}, \varphi, c_{no}; \omega) \\ &= \left[ \left( \frac{\partial \bar{V}}{\partial \bar{r}} \right)^2 + \frac{1}{\bar{r}^2} \left( \frac{\partial \bar{V}}{\partial \varphi} \right)^2 \right]^{\frac{1}{2}}, \end{aligned} \quad (13)$$

---

<sup>6</sup> "Gravitation" is used for  $\omega = 0$  and "gravity" for  $\omega \neq 0$ .

and the correction  $\Delta g = g - \bar{g}$  :

$$\begin{aligned}\Delta g &= |\text{grad } V| - |\text{grad } \bar{V}| = \Delta g(r, \bar{r}, \varphi, \lambda, C_{nm}, S_{nm}, \omega) \\ &= \left[ \left( \frac{\partial V}{\partial r} \right)^2 + \frac{1}{r^2} \left( \frac{\partial V}{\partial \varphi} \right)^2 + \frac{1}{r^2 \cos^2 \varphi} \left( \frac{\partial V}{\partial \lambda} \right)^2 \right]^{\frac{1}{2}} - \left[ \left( \frac{\partial \bar{V}}{\partial \bar{r}} \right)^2 + \frac{1}{\bar{r}^2} \left( \frac{\partial \bar{V}}{\partial \varphi} \right)^2 \right]^{\frac{1}{2}}.\end{aligned}\quad (14)$$

Example: In section a) below we find, at the equipotential spheroid ( $\omega = 0$ ) in the latitude  $\varphi = 20^\circ$  for 1,000 km elevation,

$$\bar{g} = 733.0085 \text{ cm sec}^{-2}$$

If we add to this value the  $\Delta g$  correction (of b) below in  $\varphi = 20^\circ$  and  $\lambda = 40^\circ$ , then we obtain the gravitation  $g_\varphi = 20^\circ$  at the equipotential surface  $V = \text{const}$  at  $\lambda = 40^\circ$  1,000 km elevation:

$$\bar{g} = 733.008 \text{ cm sec}^{-2}$$

$$\frac{\Delta g}{g} = + 0.004$$

$$g = 733.012 \text{ cm sec}^{-2}$$

a) Spheroidal (zonal) part  $\bar{g}$

Table 7.

$\phi^{\circ}$	(cm $\text{sec}^{-2}$ )	Sea level		Elevation above sea level ( $\omega = 0$ )		
		$\omega \neq 0$	$\omega = 0$	1,000 km	10,000 km	100,000 km
90	983.222	979.833	732.2265	148.59713	3.522 376 640	
80	983.066	979.879	732.2528	148.59823	3.522 377 260	
70	982.616	980.014	732.3291	148.60139	3.522 379 044	
60	981.931	980.226	732.4468	148.60625	3.522 381 778	
50	981.087	980.485	732.5914	148.61221	3.522 385 132	
40	980.185	980.760	732.7453	148.61855	3.522 388 702	
30	979.335	981.020	732.8902	148.62452	3.522 392 057	
20	978.638	981.231	733.0085	148.62939	3.522 394 793	
10	978.187	981.372	733.0869	148.63258	3.522 396 579	
0	978.033	981.424	733.1148	148.63370	3.522 397 200	
-10	978.188	981.374	733.0879	148.63260	3.522 396 581	
-20	978.644	981.236	733.0107	148.62943	3.522 394 796	
-30	979.338	981.024	732.8920	148.62456	3.522 392 060	
-40	980.185	980.761	732.7464	148.61858	3.522 388 704	
-50	981.090	980.488	732.5921	148.61221	3.522 385 132	
-60	981.930	980.225	732.4459	148.60622	3.522 381 776	
-70	982.608	980.006	732.3256	148.60133	3.522 379 039	
-80	983.053	979.867	732.2474	148.59814	3.522 377 252	
-90	983.209	979.820	732.2204	148.59703	3.522 376 632	

As we can see from Table 7, the gravity  $\bar{g}$  ( $\omega \neq 0$ ) increases at sea level along an equipotential surface  $\bar{V}$  ( $\omega \neq 0$ ) = const, with about 5.2 gal from the equator to the poles; further, the value of  $\bar{g}$  at the North Pole is around 13 mgal larger than the corresponding value at the South Pole. If we exclude the centrifugal force, then the gravitation  $\bar{g}$  ( $\omega = 0$ ) along the spheroid  $\bar{V}$  ( $\omega = 0$ ) = const decreases toward the poles. The difference between pole and equator amounts at sea level to about 1.6 gal, while the value at the North Pole is again 13 mgal larger than  $\bar{g}$  at the South Pole. With increasing elevation these numbers tend to zero for infinity.

Note that  $\bar{g}$  is always computed along an equipotential surface  $\bar{V} = \text{const}$  ( $\omega \neq 0$  or  $\omega = 0$ ) compared to section IIA where the surface  $\bar{g} = \text{const}$  is considered.

b) Corrections,  $\Delta g$

See Contour Map 2. At higher elevations  $\Delta g$  decreases percentagewise faster than the undulations  $\Delta r$  (see Table 8). At sea level we have values of  $\Delta g$  running from -26 mgal to 22 mgal, while at 100,000 km elevation  $\Delta g$  varies only from -0.000062 to 0.000066 mgal. Similarly to  $\Delta r$  the gravitational correction  $\Delta g$  ( $\omega = 0$ ) at sea level is very close to  $\Delta g$  ( $\omega \neq 0$ ), and hence only one case was plotted.

Table 8.

Elevation above sea level (km)	Range of $\Delta g$ (mgal)	
0	22	-26
1,000	8	-10
10,000	0.18	-0.17
100,000	0.000066	-0.000062

3. Oscillation of the surface normals; latitude and longitude curves

The surface normals are collinear with the gradients of equations (10) and (12). From Figure 6 we see that the angle  $\nu$  of two corresponding normal vectors (in  $\varphi$  and  $\lambda$ ) on  $V = \text{const}$  and  $\bar{V} = V$  follows from the scalar product

$$\cos \nu = \frac{1}{\bar{g} g} \cdot \text{grad } \bar{V} \cdot \text{grad } V , \quad (15)$$

and its components in and orthogonally to a meridian plane from the expressions

$$\xi = v \cos \alpha , \quad (16)$$

$$\eta = v \sin \alpha ;$$

$\alpha$  is the azimuth of the projection of  $v$  on  $\bar{V} = \text{const}$ , with its cosine

$$\cos \alpha = - \frac{(\text{grad } \bar{V} \times \text{grad } V) \cdot (\text{grad } \bar{V} \times \vec{e}^3)}{|\text{grad } \bar{V} \times \text{grad } V| |\text{grad } \bar{V} \times \vec{e}^3|} , \quad (17)$$

whereby  $\vec{e}^3$  is the unit vector in the  $x^3$  direction. Because  $v$ ,  $\xi$ , and  $\eta$  are

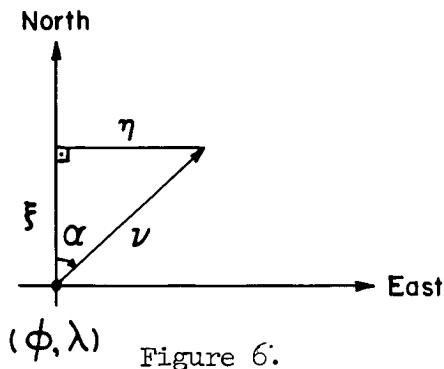


Figure 6.

small values, equations (15) and (16) can be replaced by the differential expressions

$$v = \left[ \left( \frac{1}{r} \frac{\partial \bar{r}}{\partial \varphi} - \frac{1}{r} \frac{\partial r}{\partial \varphi} \right)^2 + \frac{1}{r^2 \cos^2 \varphi} \left( \frac{\partial r}{\partial \lambda} \right)^2 \right]^{\frac{1}{2}},$$

$$\xi = \frac{1}{r} \frac{\partial \bar{r}}{\partial \varphi} - \frac{1}{r} \frac{\partial r}{\partial \varphi}, \quad (18)$$

$$\eta = - \frac{1}{r \cos \varphi} \frac{\partial r}{\partial \lambda},$$

wherein

$$\frac{\partial r}{\partial p^i} = - \frac{\partial V / \partial p^i}{\partial V / \partial r} \quad (\text{analogously for } \bar{r}), \quad (19)$$

with  $\varphi = p^1$  and/or  $\lambda = p^2$ . The sign of  $\xi$ ,  $\eta$  was defined according to geodetic usage. If we use the normal of an equipotential surface for latitude and longitude definition, as in section 3d), then  $\xi$  is positive if the latitude  $B$  of a point on  $V = \text{const}$  is larger than the latitude  $\bar{B}$  of the corresponding point on  $\bar{V} = \text{const}$ , and analogously for the longitude (Bomford, 1962).

### a) Oscillation components $\xi$ and $\eta$

See Contour Maps 3 and 4. The absolute oscillation  $v = (\xi^2 + \eta^2)^{\frac{1}{2}}$  takes the largest values where the inclination of the undulations is a maximum. If we split  $v$  up into its components, then  $\xi$  represents the inclination in the meridional direction while  $\eta$  accounts for the effect orthogonally to it (see Tables 9 and 10). Superposing the  $\xi$ ,  $\eta$  graphs with the corresponding  $\Delta r$  graphs, the zero curves of  $\xi$  intersect the undulation lines along points where the tangents of the  $\Delta r$  lines coincide with the meridians; analogously, the zero curves of  $\eta$  intersect at points where the latitude lines touch the  $\Delta r$  lines.

Table 9.

Elevation above sea level (km)	Range of $\xi$ (seconds of arc)	
0	5".7	-5".4
1,000	3".1	-2".9
10,000	0".25	-0".29
100,000	0".0039	-0".0038

Table 10.

Elevation above sea level (km)	Range of $\eta$ (seconds of arc)	
0	5".7	-6".8
1,000	2".9	-3".9
10,000	0".34	-0".42
100,000	0".0070	-0".0074

- b) The angle between the spheroid normal ( $\omega \neq 0$ ) and the ellipsoid normal ( $f^{-1} = 298.25$ ) in the same geocentric latitude; deflection of the vertical

By adding  $\delta\xi$  of Table 11 to  $\xi$  of sea level, we get the oscillations of the surface normals of the satellite geoid relative to Kozai's reference ellipsoid (see Figure 7);  $\eta$  remains unchanged. In geodesy these values  $\xi + \delta\xi$  and  $\eta$  are known as components of the deflection of the vertical (Bomford, 1962).

Example: At  $\varphi = 20^\circ$  and  $\lambda = 40^\circ$  we find (see pages 74, 76) the deflection components of the vertical:

$$\begin{array}{rcl} \xi & = & -1''5 \\ \underline{\delta\xi} & = & +0''02 \\ \xi + \delta\xi & = & -1''5 \end{array} \quad \begin{array}{rcl} \eta & = & 2''0 \\ - & & - \\ \eta & = & 2''0 \end{array}$$

Table 11.

$\varphi^\circ$	$\delta\xi$
90	0''00
80	-0''66
70	-0''63
60	-0''45
50	-0''81
40	-0''75
30	-0''66
20	+0''02
10	+1''18
0	+0''59
-10	+0''37
-20	+0''67
-30	-0''22
-40	-0''07
-50	-0''52
-60	-2''24
-70	-1''98
-80	-0''78
-90	0''00

$$\delta\xi = \cancel{\vec{n}_{\text{ELL}}} (\vec{n}_{\text{ELL}}, \vec{n} (\omega \neq 0))$$

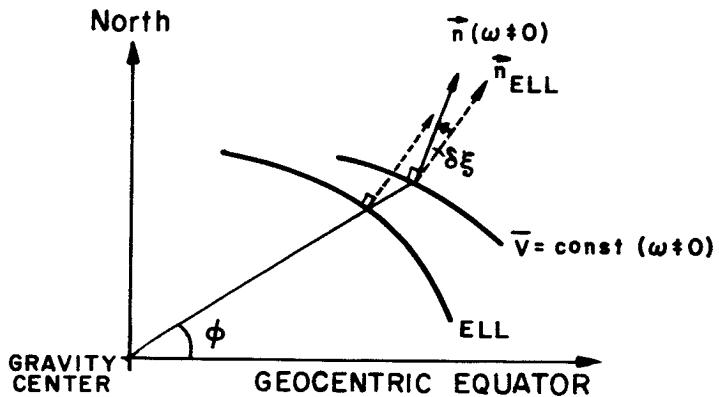


Figure 7.

- c) Angle between the spheroid normal ( $\omega \neq 0$ ) and the spheroid normal ( $\omega = 0$ ) along  $\bar{v} = \text{const } (\omega \neq 0)$  at sea level

Table 12 and Figure 8 show the angles with which the spheroids ( $\omega = 0$ ) intersect the spheroid ( $\omega \neq 0$ ) at sea level. The sign of  $\Delta\xi$  follows again from the definition of  $\xi$ .

Table 12.

$\varphi^\circ$	$\Delta\xi$
90	0' 0"
80	2' 1"
70	3' 48"
60	5' 8"
50	5' 51"
40	5' 52"
30	5' 10"
20	3' 50"
10	2' 3"
0	0' 0"
-10	-2' 3"
-20	-3' 50"
-30	-5' 10"
-40	-5' 52"
-50	-5' 51"
-60	-5' 8"
-70	-3' 48"
-80	-2' 1"
-90	0' 0"

$$\Delta\xi = \vec{\xi} (\vec{n} (\omega = 0), \vec{n} (\omega \neq 0))$$

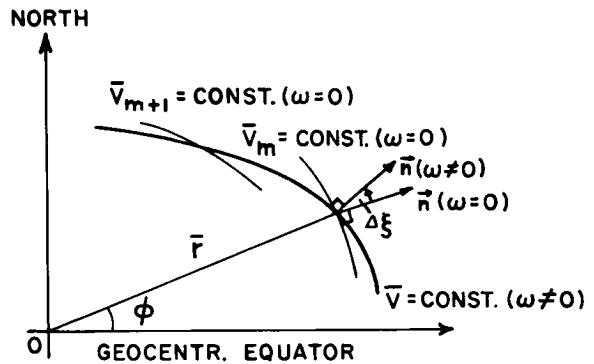


Figure 8.

d) Shape of the latitude and longitude curves

In this section we consider the latitude and longitude of a point defined by the surface normal rather than by the geocentric radius. Let us call latitude B the complement of the angle included by a surface normal and the revolution axis, and analogously, let us call longitude L the complement of the angle of a plane determined by a surface normal and the revolution axis with the  $x^2$  axis. Then the latitude and longitude curves of an equipotential surface have the equations:

$$B = \text{const},$$

$$L = \text{const.}$$

To obtain their amplitudes relative to the spheroidal part we consider equations (10) to (12); the explicit expression for the latitude curves,

$$B = \text{const} = \arcsin \frac{\sin \varphi - \frac{1}{r} \frac{\partial r}{\partial \varphi} \cos \varphi}{\left[ 1 + \frac{1}{r^2} \left( \frac{\partial r}{\partial \varphi} \right)^2 + \frac{1}{r^2 \cos^2 \varphi} \left( \frac{\partial r}{\partial \lambda} \right)^2 \right]^{\frac{1}{2}}} , \quad (20)$$

shows that  $B$  is mainly a function of  $\varphi$  plus a small correction term accounting for the undulation effect in  $V = \text{const}$ . Within a small range

$$B = f(\varphi) + \Delta , \quad (21)$$

with

$$\Delta = \text{const} ; \quad (22)$$

by differentiation

$$dB = \frac{df}{d\varphi} d\varphi . \quad (23)$$

Equating  $\xi$  to  $dB$ , we get the corresponding angle  $d\varphi$  by which the latitude curve  $B = \text{const}$  is shifted along the meridian. Because  $\frac{df}{d\varphi}$  is close to 1, its amplitude can be expressed by

$$A_B \approx -r \cdot \xi , \quad (24)$$

and similarly the amplitude of the longitude curve

$$A_L \approx -r \eta . \quad (25)$$

In the  $\xi$  and  $\eta$  plottings (see Contour Maps 3 and 4) only the amount of the equi-level line numbers must be changed, according to Table 13, to obtain the corresponding graphs of  $A_B$  and  $A_L$ .

Table 13.

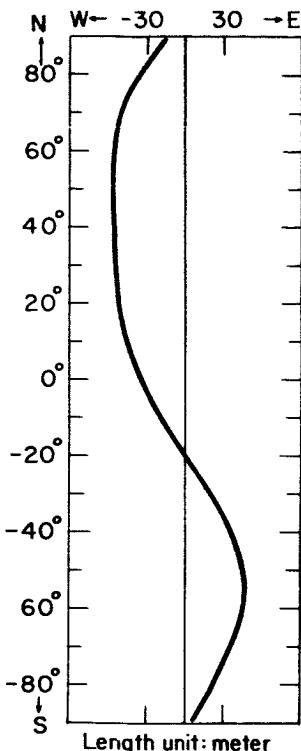
Sea level		1,000 km elevation		10,000 km elevation		100,000 km elevation	
$\xi, \eta$	$A_B, A_L$ (m)	$\xi, \eta$	$A_B, A_L$ (m)	$\xi, \eta$	$A_B, A_L$ (m)	$\xi, \eta$	$A_B, A_L$ (m)
5"	-155			0":30	-24	0":006	-3.1
4"	-124			0":25	-20	0":004	-2.1
3"	-93	3"	-107	0":20	-16	0":003	-1.6
2"	-62	2"	-72	0":15	-12	0":002	-1.0
1"	-31	1"	-36	0":10	-8	0":001	-0.5
0"	0	0"	0	0":05	-4	0":001	0
-1"	31	-1"	36	0":05	4	-0":001	0.5
-2"	62	-2"	72	-0":05	8	-0":002	1.0
-3"	93	-3"	107	-0":10	12	-0":003	1.6
-4"	124	-4"	143	-0":15	16	-0":004	2.1
-5"	155			-0":20	20	-0":006	3.1
-6"	186			-0":25	24		
				-0":30	28		
				-0":35			
				-0":40	32		

$A_B$  + latitude curve oscillates northward  
 $A_B$  - latitude curve oscillates southward

$A_L$  + longitude curve oscillates eastward  
 $A_L$  - longitude curve oscillates westward

The zero curves of  $A_B$ ,  $A_L$  and the level lines of the undulations  $\Delta r$  intersect each other as already mentioned for  $\xi, \eta$ .

Examples: 1) The latitude curve  $B = 20^\circ$ , on the satellite geoid, is at the geocentric longitude  $\lambda = 40^\circ$  about  $A_B = 46$  m northward of the corresponding  $B = 20^\circ$  latitude circle of its spheroid.

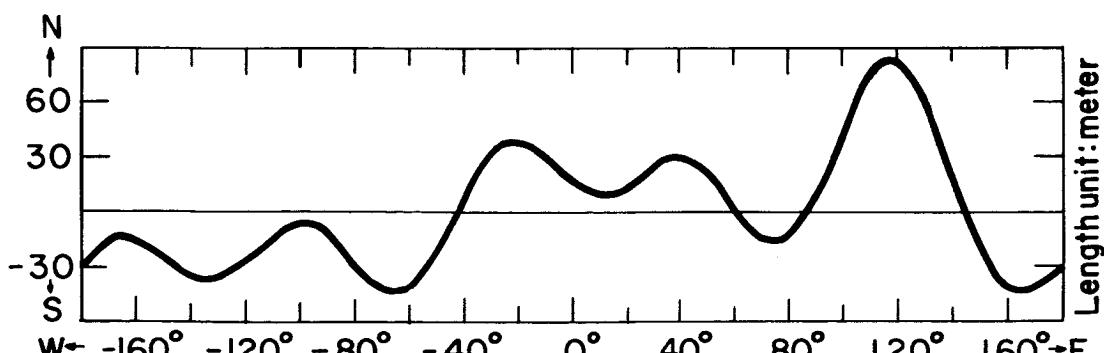


a)

2) See Figure 9.

Figure 9.--a) Shape of the longitude curve  
 $L = 0$  at sea level.

b) Shape of the equator curve  
 $B = 0$  at sea level.



b)

#### 4. Gaussian and mean curvature

With the position vector

$$\vec{x} = \vec{x}(p^1, p^2) \quad p^1, p^2 \text{ stand for } \varphi, \lambda \quad (26)$$

of an equipotential surface  $V = \text{const}$  and its unit normal (to the surface)

$$\vec{n} = -\frac{\underline{\text{grad}} V}{g} , \quad (27)$$

we get by partial differentiation of equation (26) a covariant vector

$$\vec{x}_i = \frac{\partial \vec{x}}{\partial p^i} , \quad (28)$$

which leads to the metric tensor

$$a_{ij} = \vec{x}_i \cdot \vec{x}_j . \quad (29)$$

Similarly, we proceed with the normal vector  $\vec{n}$ :

$$\vec{n}_i = \frac{\partial \vec{n}}{\partial p^i} , \quad (30)$$

wherein

$$\frac{\partial^2 r}{\partial p^i \partial p^j} = -\frac{1}{\partial r} \left( \frac{\partial^2 V}{\partial p^i \partial p^j} + \frac{\partial^2 V}{\partial r \partial p^i} \frac{\partial r}{\partial p^j} + \frac{\partial^2 V}{\partial r \partial p^j} \frac{\partial r}{\partial p^i} + \frac{\partial^2 V}{\partial r^2} \frac{\partial r}{\partial p^i} \frac{\partial r}{\partial p^j} \right) . \quad (31)$$

The covariant vector  $\vec{n}_i$  defines together with equation (28) a covariant tensor

$$b_{ij} = - \vec{n}_i \vec{x}_j , \quad (32)$$

from which we obtain with equation (29) the gaussian curvature in question:

$$K = \frac{\det b_{ij}}{\det a_{ij}} . \quad (33)$$

An analogous expression is found for the mean curvature

$$H = \frac{1}{2} a^{ij} b_{ij} , \quad (34)$$

with  $a^{ij}$  the contravariant complement of equation (29).

In our numerical computations we did not consider  $K$  and  $H$  directly, but the radii of curvature instead:

$$\frac{1}{\sqrt{K}} = \frac{1}{\sqrt{H}}(r, \varphi, \lambda, C_{nm}, S_{nm}; \omega) , \quad (35)$$

and

$$\frac{1}{H} = \frac{1}{H}(r, \varphi, \lambda, C_{nm}, S_{nm}; \omega) , \quad (36)$$

which we split up into a zonal part  $\frac{1}{\sqrt{K}}$ ,  $\frac{1}{H}$  and a correction term

$$\Delta \frac{1}{\sqrt{K}} = \frac{1}{\sqrt{K}} - \frac{1}{\bar{H}} \quad (37)$$

and

$$\Delta \frac{1}{H} = \frac{1}{H} - \frac{1}{\bar{H}} \quad . \quad (38)$$

Although  $\frac{1}{\sqrt{K}}$  and  $\frac{1}{H}$  are fairly different along the equipotential surface  $\bar{V} = \text{const}$  except at the pole  $\frac{1}{\sqrt{K}} = \frac{1}{H}$ ,<sup>7</sup> the corrections  $\Delta \frac{1}{\sqrt{K}}$  and  $\Delta \frac{1}{H}$  agree everywhere within several meters. This is easily explained by the following: calling  $\bar{R}_1$  and  $\bar{R}_2$  the principle radii of curvature of the spheroid, and

$$\begin{aligned} R_1 &= \bar{R}_1 + \Delta R_1 \\ R_2 &= \bar{R}_2 + \Delta R_2 \end{aligned} \quad (39)$$

the analogous radii for the equipotential surface  $V = \text{const}$ , we obtain from equations (37) and (39)

$$\begin{aligned} \Delta \frac{1}{\sqrt{K}} &= (R_1 R_2)^{\frac{1}{2}} - (\bar{R}_1 \bar{R}_2)^{\frac{1}{2}} \approx (\bar{R}_1 \bar{R}_2)^{\frac{1}{2}} \left( 1 + \frac{\Delta R_1}{\bar{R}_1} + \frac{\Delta R_2}{\bar{R}_2} \right)^{\frac{1}{2}} - \\ &- (\bar{R}_1 \bar{R}_2)^{\frac{1}{2}} \approx \frac{1}{2}(\Delta R_1 + \Delta R_2) \quad , \end{aligned} \quad (40)$$

---

<sup>7</sup>At ground level; at higher elevations, correspondingly closer.

and similarly

$$\Delta \frac{1}{H} = \frac{2R_1 R_2}{R_1 + R_2} - \frac{2\bar{R}_1 \bar{R}_2}{\bar{R}_1 + \bar{R}_2} \approx \frac{1}{2} (\Delta R_1 + \Delta R_2) ; \quad (41)$$

this means

$$\Delta \frac{1}{\sqrt{K}} \approx \Delta \frac{1}{H} \quad . \quad (42)$$

In the contour maps we considered, therefore, only  $\Delta \frac{1}{\sqrt{K}}$  in the different sections of outer space.

Example: With the help of sections a) and b) we obtain for the radii of the gaussian and mean curvature at  $\varphi = 20^\circ$ ,  $\lambda = 40^\circ$  at sea level ( $w \neq 0$ )

$$\begin{array}{rcl} \frac{1}{\sqrt{K}} & = & 6,361,929 \text{ m} \\ \Delta \frac{1}{\sqrt{K}} & = & - 330 \\ \hline \frac{1}{\sqrt{K}} & = & 6,361,599 \text{ m} \end{array} \quad \begin{array}{rcl} \frac{1}{H} & = & 6,361,902 \text{ m} \\ \Delta \frac{1}{H} & = & - 330 \\ \hline \frac{1}{H} & = & 6,361,572 \text{ m} \end{array}$$

or the curvature

$$\begin{aligned} \bar{K} &= 0.247\ 071\ 3 \times 10^{-13} \text{ m}^{-2} & K &= 0.247\ 097\ 0 \times 10^{-13} \text{ m}^{-2} \\ \bar{H} &= 0.157\ 185\ 7 \times 10^{-6} \text{ m}^{-1} & H &= 0.157\ 193\ 8 \times 10^{-6} \text{ m}^{-1} \end{aligned}$$

These values may also be used in geodesy for best approximating spheres in local geodetic reference systems, etc.

a) Radius of curvature of the spheroid

Table 14.  
Radius of the gaussian curvature (in meters)

$\phi^\circ$	Sea level		Elevation above sea level ( $\omega = 0$ )		
	$\omega \neq 0$	$\omega = 0$	1,000 km	10,000 km	100,000 km
90	6 399 497	6 388 407	7 387 065.7	16 382 189.50	106 378 785.7547
80	6 398 257	6 387 831	7 386 545.1	16 381 947.14	106 378 748.3272
70	6 394 651	6 386 131	7 385 022.6	16 381 249.23	106 378 640.5556
60	6 388 942	6 383 350	7 382 643.3	16 380 179.61	106 378 475.4296
50	6 381 954	6 379 959	7 379 725.1	16 378 866.90	106 378 272.8535
40	6 374 557	6 376 400	7 376 629.5	16 377 469.00	106 378 057.2485
30	6 367 540	6 372 999	7 373 718.5	16 376 154.05	106 377 854.6113
20	6 361 929	6 370 340	7 371 355.0	16 375 080.35	106 377 689.3805
10	6 358 110	6 368 452	7 369 762.9	16 374 377.58	106 377 581.4904
0	6 356 693	6 367 706	7 369 183.1	16 374 131.19	106 377 543.9652
-10	6 358 130	6 368 472	7 369 743.9	16 374 371.64	106 377 581.3461
-20	6 361 781	6 370 192	7 371 288.5	16 375 070.40	106 377 689.1396
-30	6 367 496	6 372 954	7 373 675.6	16 376 143.43	106 377 854.3560
-40	6 374 602	6 376 444	7 376 617.7	16 377 461.50	106 378 057.0731
-50	6 381 861	6 379 868	7 379 696.4	16 378 865.87	106 378 272.8385
-60	6 388 932	6 383 340	7 382 654.1	16 380 187.25	106 378 475.6191
-70	6 394 817	6 386 294	7 385 118.8	16 381 265.84	106 378 640.9439
-80	6 398 472	6 388 043	7 386 695.1	16 381 970.55	106 378 748.8591
-90	6 399 725	6 388 634	7 387 230.4	16 382 215.42	106 378 786.3390

The gaussian and the mean curvature of the potential spheroids resulted in positive values (see Tables 14 and 15). Hence the radius of the gaussian curvature is always larger than or equal to the radius of the mean curvature over the whole surface. At the poles both radii are identical; that is, the principal radii of curvature are equal (spherical surface section) and Dupin's indicatrix is circular. For infinite elevation the radius of the gaussian curvature becomes equal to the radius of the mean curvature according to the spherical shape assumed by the equipotential spheroid.

Table 15.  
Radius of the mean curvature (in meters)

$\varphi^\circ$	$\frac{1}{H}$	Sea level		Elevation above sea level ( $\omega = 0$ )		
		$\omega \neq 0$	$\omega = 0$	1,000 km	10,000 km	100,000 km
90	6 399 497	6 388 407	7 387 065.7	16 382 189.50	106 378 785.7547	
80	6 398 257	6 387 831	7 386 545.1	16 381 947.14	106 378 748.3272	
70	6 394 651	6 386 131	7 385 022.6	16 381 249.23	106 378 640.5556	
60	6 388 940	6 383 349	7 382 643.0	16 380 179.58	106 378 475.4295	
50	6 381 948	6 379 958	7 379 724.2	16 378 866.82	106 378 272.8532	
40	6 374 545	6 376 397	7 376 627.7	16 377 468.83	106 378 057.2479	
30	6 367 520	6 372 994	7 373 715.5	16 376 153.77	106 377 854.6103	
20	6 361 902	6 370 334	7 371 350.8	16 375 079.96	106 377 689.3791	
10	6 358 076	6 368 444	7 369 757.8	16 374 377.12	106 377 581.4886	
0	6 356 657	6 367 697	7 369 177.6	16 374 130.69	106 377 543.9634	
-10	6 358 096	6 368 464	7 369 738.7	16 374 371.17	106 377 581.3444	
-20	6 361 752	6 370 185	7 371 284.3	16 375 070.02	106 377 689.1382	
-30	6 367 475	6 372 949	7 373 672.5	16 376 143.15	106 377 854.3549	
-40	6 374 590	6 376 441	7 376 615.9	16 377 461.33	106 378 057.0725	
-50	6 381 855	6 379 866	7 379 695.5	16 378 865.78	106 378 272.8382	
-60	6 388 930	6 383 340	7 382 653.7	16 380 187.22	106 378 475.6190	
-70	6 394 816	6 386 294	7 385 118.8	16 381 265.83	106 378 640.9439	
-80	6 398 472	6 388 043	7 386 695.1	16 381 970.55	106 378 748.8591	
-90	6 399 725	6 388 634	7 387 230.4	16 382 215.42	106 378 786.3390	

b) Corrections,  $\Delta \frac{1}{\sqrt{K}}$  and  $\Delta \frac{1}{H}$

See Contour Map 5. The corrections  $\Delta \frac{1}{\sqrt{K}}$  and  $\Delta \frac{1}{H}$  usually have their maxima where the undulations  $\Delta r$  have their minima, and vice versa. However, there are also considerable deviations possible, depending on the shape of the undulations. At the poles the radius of curvature is not identical with the spheroidal part, but the difference is practically negligible. The diminution of  $\Delta \frac{1}{\sqrt{K}}$ ,  $\Delta \frac{1}{H}$  with higher elevation is faster than that of the undulations.

Table 16.

Elevation above sea level (km)	Range of $\Delta \frac{1}{\sqrt{K}}$ , $\Delta \frac{1}{H}$ (m)	
0	567	-470
1,000	331	-271
10,000	43	-45
100,000	3.8	-4.1

(see Table 16). The numbers drop at sea level from -470 m and/or 567 m to -4.1 m and/or 3.8 m at 100,000 km elevation. Again at sea level only the case of  $w \neq 0$  was considered, because of the close agreement with that of  $w = 0$ .

## II. GRADIENT FIELD

### A. Magnitude of the Gradient Field - Equigravitational (Equigravity) Surfaces

From the equation

$$g = \left[ \left( \frac{\partial V}{\partial r} \right)^2 + \frac{1}{r^2} \left( \frac{\partial V}{\partial \varphi} \right)^2 + \frac{1}{r^2 \cos^2 \varphi} \left( \frac{\partial V}{\partial \lambda} \right)^2 \right]^{\frac{1}{2}}, \quad (43)$$

we get the magnitude of the gradient  $V$  in each point of outer space. If we put

$$g = \text{const}, \quad (44)$$

then equation (43) describes an equigravitational (equigravity) surface whose shape will be determined in the next section. Analogous to section I we define a rotational surface

$$\bar{g} = \text{const} = \left[ \left( \frac{\partial \bar{V}}{\partial \bar{r}} \right)^2 + \frac{1}{\bar{r}^2} \left( \frac{\partial \bar{V}}{\partial \varphi} \right)^2 \right]^{\frac{1}{2}}, \quad (45)$$

which follows from equation (43) by simply neglecting the nonzonal terms. Further,  $g$  and  $\bar{g}$  are related by the equation

$$g = \bar{g} = \text{const}, \quad (46)$$

wherein  $\bar{r}$  assumes the following values: at sea level

$$\bar{r} = 6,378,165 \text{ m}, \text{ for } \varphi = 0; \omega \neq 0 \text{ or } \omega = 0,$$

and for the outer space

$$\left. \begin{array}{l} \bar{r} = 6,378,165 \text{ m} + 1,000 \text{ km} \\ \bar{r} = 6,378,165 \text{ m} + 10,000 \text{ km} \\ \bar{r} = 6,378,165 \text{ m} + 100,000 \text{ km} \end{array} \right\} \quad \begin{array}{l} \varphi = 0 \\ \omega = 0 \end{array} .$$

The results concerning the geometric structure of the equigravitational (equigravity) surfaces are again split up into two parts. At first the portion of  $\bar{g} = \text{const}$  is computed and then a correction term is added, leading to the geometrical structure of the equigravitational (equigravity) surfaces  $g = \text{const}$ . The only difference from section I is that  $\bar{g} = \text{const}$  no longer approximates  $g = \text{const}$  in the mean. However, this discrepancy is very small and can often be neglected.

1. Shape of the equigravitational (equigravity) surfaces

Similar to section I we obtain the spheroidal part  $\bar{r}$  of the geocentric radius (see Figure 10) from the equation

$$\bar{g} = \bar{g}(\bar{r}, \varphi, C_{no}; \omega) = \text{const} , \quad (47)$$

and the correction term from

$$g = g(\bar{r} + \Delta r, \varphi, \lambda, C_{nm}, S_{nm}; \omega) = \text{const} . \quad (48)$$

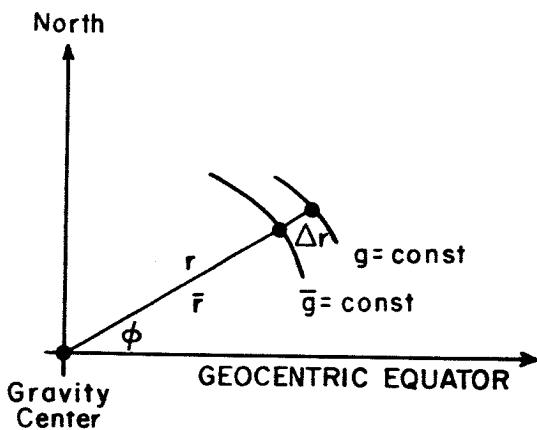


Figure. 10.

Their sum,

$$r(\varphi, \lambda) = \bar{r}(\varphi) + \Delta r(\varphi, \lambda) ,$$

leads to the geocentric radius along an equigravitational (equigravity) surface whose zonal part has the above-mentioned characteristics (pages 33-34).

Example: With the help of sections a) and b) we obtain the geocentric radius  $r$  at  $\varphi = 20^\circ$  and  $\lambda = 40^\circ$  at 10,000 km elevation ( $\omega = 0$ ):

$$\begin{array}{rcl} \bar{r} & = & 16,377,455 \text{ m} \\ \Delta r & = & -4 \\ r_{\varphi = 20^\circ} & = & 16,377,451 \text{ m} \\ \lambda = 40^\circ & & \end{array}$$

a) Spheroidal (zonal) part of the geocentric radius

Table 17.

$\varphi^\circ$	$\bar{r}$ (m)	Sea level		Elevation above sea level ( $\omega = 0$ )		
		$\omega \neq 0$	$\omega = 0$	1,000 km	10,000 km	100,000 km
90	6 373 689.54	6 362 631.37	7 364 739.30	16 372 117.71	106 377 233.56	
80	6 373 816.80	6 363 091.80	7 365 140.74	16 372 299.76	106 377 261.64	
70	6 374 191.84	6 364 426.28	7 366 300.04	16 372 824.01	106 377 342.49	
60	6 374 784.84	6 366 489.70	7 368 081.84	16 373 627.39	106 377 466.37	
50	6 375 512.59	6 369 021.52	7 370 269.12	16 374 613.21	106 377 618.34	
40	6 376 284.31	6 371 713.45	7 372 597.30	16 375 662.79	106 377 780.07	
30	6 377 013.98	6 374 247.99	7 374 786.99	16 376 649.69	106 377 932.06	
20	6 377 605.10	6 376 310.69	7 376 575.00	16 377 454.93	106 378 055.98	
10	6 378 011.24	6 377 677.53	7 377 750.54	16 377 981.27	106 378 136.88	
0	6 378 165.00	6 378 165.00	7 378 165.00	16 378 165.00	106 378 165.00	
-10	6 378 022.37	6 377 688.70	7 377 761.57	16 377 983.67	106 378 136.94	
-20	6 377 638.43	6 376 344.11	7 376 596.06	16 377 458.94	106 378 056.08	
-30	6 377 038.57	6 374 272.61	7 374 806.58	16 376 653.97	106 377 932.16	
-40	6 376 294.59	6 371 723.70	7 372 609.49	16 375 665.78	106 377 780.14	
-50	6 375 524.34	6 369 033.32	7 370 273.96	16 374 613.55	106 377 618.34	
-60	6 374 773.82	6 366 478.63	7 368 070.50	16 373 624.25	106 377 466.29	
-70	6 374 142.71	6 364 376.89	7 366 265.76	16 372 817.40	106 377 342.33	
-80	6 373 743.74	6 363 018.43	7 365 088.90	16 372 290.59	106 377 261.42	
-90	6 373 608.45	6 362 549.97	7 364 681.22	16 372 107.59	106 377 233.32	

At sea level the geocentric radii at the poles are about 4.5 km ( $\omega \neq 0$ ) and/or 15.6 km ( $\omega = 0$ ) smaller than the corresponding equatorial radius, as shown in Table 17. At higher elevations this difference decreases and finally becomes zero for infinity. Note that the polar radii at sea level are bigger for ( $\omega \neq 0$ ) than for ( $\omega = 0$ ), which is opposite to the results obtained for the equipotential surfaces.

b) Correction term,  $\Delta r$

See Contour Map 6, page 80 . The size of the corrections  $\Delta r$  is much bigger than the undulations of section II. They run now at sea level from -158 m to 116 m (see Table 18); this roughing process is easily explained by the differentiation involved (see equation (43)). At 100,000 km, however, the  $C_{22}$  and the  $S_{22}$  terms are again the most significant ones.

Table 18.

Elevation above sea level (km)	Range of $\Delta r$ (m)		
0	116	:	-158
1,000	85	:	-104
10,000	26	:	- 24
100,000	2.9	:	- 2.7

c) Difference between the geocentric radii of the Northern and Southern Hemispheres of the spheroid

See page 36. From Table 19 we see that the equigravitational (equigravity) surfaces show a stronger asymmetry of the Northern compared to the Southern Hemispheres than the equipotential surfaces (see section IIc)).

Table 19.

$ \varphi^\circ $	$\frac{\bar{r}(\varphi) - \bar{r}(-\varphi)}{(m)}$	Sea level		Elevation above sea level ( $\omega = 0$ )		
		$\omega \neq 0$	$\omega = 0$	1,000 km	10,000 km	100,000 km
90	+81.09	+81.40	+58.08	+10.12	+0.234	
80	+73.06	+73.37	+51.84	+ 9.17	+0.213	
70	+49.13	+49.39	+34.28	+ 6.61	+0.155	
60	+11.02	+11.07	+11.34	+ 3.14	+0.076	
50	-11.75	-11.80	- 4.84	- 0.34	-0.006	
40	-10.28	-10.25	-12.19	- 2.99	-0.070	
30	-24.59	-24.62	-19.59	- 4.28	-0.102	
20	-33.33	-33.42	-21.06	- 4.01	-0.096	
10	-11.13	-11.17	-11.03	- 2.40	-0.058	
0	0.00	0.00	0.00	0.00	0.000	

d) Difference of the spheroid and an ellipsoid with equal axes

Table 20.

$ \varphi^\circ $	$\frac{\bar{r} - E}{(m)}$	Sea level		Elevation above sea level ( $\omega = 0$ )		
		$\omega \neq 0$	$\omega = 0$	1,000 km	10,000 km	100,000 km
90	0.00	0.00	0.00	0.00	0.00	0.0000
80	- 7.56	- 6.31	- 2.33	- 0.20	-0.0065	
70	-20.74	-16.32	- 5.98	- 0.76	-0.0243	
60	-22.68	-14.43	- 7.01	- 1.52	-0.0475	
50	-24.96	-14.20	- 8.46	- 2.28	-0.0687	
40	-30.40	-19.64	-11.63	- 2.80	-0.0805	
30	-31.27	-22.94	-14.70	- 2.86	-0.0781	
20	-35.88	-31.34	-15.69	- 2.33	-0.0607	
10	-18.67	-17.40	- 8.55	- 1.28	-0.0320	
0	0.00	0.00	0.00	0.00	0.0000	
-10	- 5.09	- 3.77	+ 4.24	+ 1.41	+0.0327	
-20	+ 6.95	+11.66	+12.18	+ 2.86	+0.0630	
-30	+13.62	+22.14	+19.47	+ 3.94	+0.0824	
-40	+13.43	+24.38	+24.63	+ 4.36	+0.0861	
-50	+34.41	+45.51	+30.53	+ 3.99	+0.0743	
-60	+27.15	+35.66	+25.26	+ 2.94	+0.0518	
-70	+ 1.75	+ 6.23	+11.05	+ 1.58	+0.0267	
-80	- 1.96	- 0.71	+ 2.17	+ 0.44	+0.0073	
-90	0.00	0.00	0.00	0.00	0.0000	

\* E is the geocentric ellipsoidal radius.

The results presented in Table 20 resemble those in section IIId); however, by the differentiation process, the amplitudes increase or, in other words, the spheroidal surfaces become rougher (see  $\varphi = -10^\circ, -80^\circ$  at sea level); the maximum is reached at sea level ( $\omega = 0$ ) around the latitude  $\varphi = -50^\circ$ .

e) Difference of the spheroids ( $\omega \neq 0$ ) and ( $\omega = 0$ ) at sea level

Table 21.

$\varphi^\circ$	$\bar{r}(\omega \neq 0) - \bar{r}(\omega = 0)$ (m)
90	11,058.17
80	10,725.00
70	9,765.56
60	8,295.14
50	6,491.07
40	4,570.86
30	2,765.99
20	1,294.41
10	333.71
0	0.00
-10	333.67
-20	1,294.32
-30	2,765.96
-40	4,570.89
-50	6,491.02
-60	8,295.19
-70	9,765.82
-80	10,725.31
-90	11,058.48

The results from Table 21 are opposite to those in section IIIf), i.e., the equigravity spheroid ( $\omega \neq 0$ ) encloses the equigravitational spheroid ( $\omega = 0$ ) at sea level.

f) Mean flattening<sup>8</sup> of the spheroids ( $\omega = 0$ ) as a function of the elevation above sea level

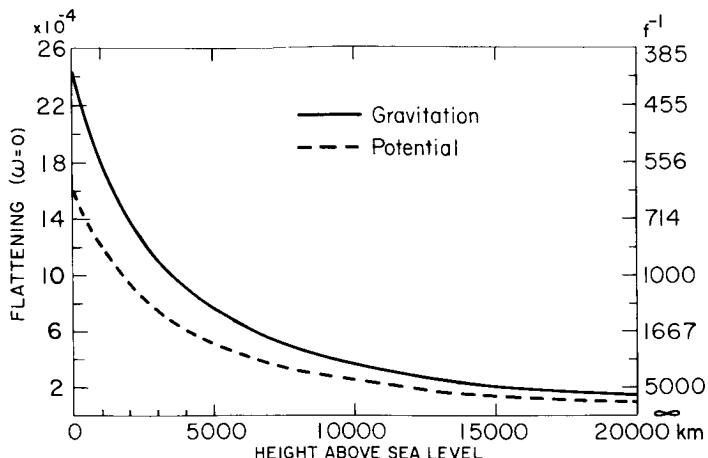


Figure 11.

The mean flattening of the equigravitational spheroids decreases asymptotically with increasing elevation above sea level. For contrast we also show in Figure 11 the mean flattening of the equipotential spheroids ( $\omega = 0$ ).

Comparison: Flattening of the equigravitational (equigravity) spheroids and equipotential spheroids:

Table 22.

$\bar{r}(\varphi = 0)$	$f^{-1}$ ( $\bar{V} = \text{const}$ )	$f^{-1}$ ( $\bar{g} = \text{const}$ )	$\bar{r}_{90^\circ}(\bar{V}) - \bar{r}_{90^\circ}(\bar{g})$ (km)	Elevation
6,378,165 m ( $\omega \neq 0$ )	298	1,412	-16.9	sea level
6,378,165 m ( $\omega = 0$ )	615	409	5.2	sea level
7,378,165 m ( $\omega = 0$ )	823	548	4.5	1,000 km above sea level
16,378,165 m ( $\omega = 0$ )	4,060	2,706	2.0	10,000 km above sea level
106,378,165 m ( $\omega = 0$ )	171,290	114,193	0.3	100,000 km above sea level

<sup>8</sup>For the definition of the mean flattening see page 12.

Including the centrifugal force ( $\omega \neq 0$ ), the equipotential spheroid stays at sea level within the equigravity spheroid. Otherwise, if no centrifugal force is included ( $\omega = 0$ ), the equigravitational spheroid is surrounded by the equipotential spheroid. At infinite elevation, both the equipotential and the equigravitational spheroid fall together (see also Figure 11). In column 4 of Table 22 we also give the polar distance between an equipotential and equigravitational spheroid assuming that the equatorial radius is identical. Note the decrease of the numbers with higher elevation.

- g) Distance of the actual equator of the spheroid from the gravity center

As in section IIh), we find the distance  $D$  of the actual equator from the gravity center:

$$D \approx a \left[ -3 c_{30} \left(1 + \frac{h}{a}\right)^{-2} + \frac{45}{8} c_{50} \left(1 + \frac{h}{a}\right)^{-4} - \frac{35}{4} c_{70} \left(1 + \frac{h}{a}\right)^{-6} \right. \\ \left. + \frac{1575}{128} c_{90} \left(1 + \frac{h}{a}\right)^{-8} - \frac{2079}{128} c_{110} \left(1 + \frac{h}{a}\right)^{-10} + \frac{21021}{1024} c_{130} \left(1 + \frac{h}{a}\right)^{-12} \dots \right] . \quad (49)$$

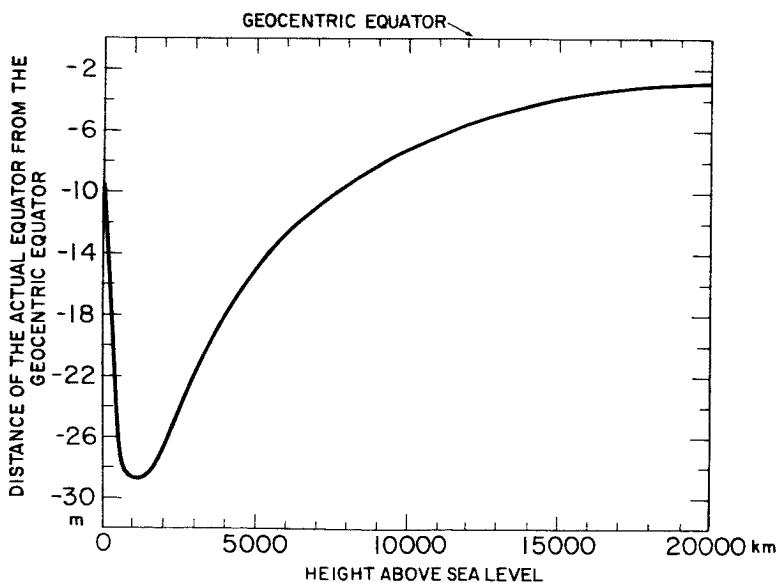


Figure 12.

Near sea level the actual equator moves at first southward and returns then asymptotically to the geocentric equator with increasing height, as is shown in Figure 12.

## 2. Oscillation of the surface normals

See section I3 and Contour Maps 7 and 8. The only basic difference here is the use of the function  $\bar{g}$  and  $g$  instead of  $\bar{V}$  and  $V$  for the determination of  $\xi$  and  $\eta$ . Both results in section I3 and II A2 are needed in the next section to obtain the plane, defined by the corresponding surface normals of  $V = \text{const}$  and  $g = \text{const}$ .

Table 23.

Elevation above sea level (km)	Range of $\xi$ (seconds of arc)	
0	15"	-15"
1,000	7.8	- 7.7
10,000	0.43	- 0.52
100,000	0.0061	- 0.0060

Table 24.

Elevation above sea level (km)	Range of $\eta$ (seconds of arc)	
0	16"	-17"
1,000	8.0	- 9.1
10,000	0.54	- 0.72
100,000	0.011	- 0.011

The surface normals of the equigravitational (equigravity) surfaces oscillate in a far larger range than those of the equipotential surfaces. Tables 23 and 24 show the magnitudes of the components  $\xi$ ,  $\eta$  at different elevations. The results at sea level were again practically the same for  $g(\omega \neq 0) = \text{const}$  and  $g(\omega = 0) = \text{const}$ , and hence only one contour map was plotted.

3. Intersection of the equigravitational (equigravity) surfaces with the equipotential surfaces

The angle by which the surfaces  $V = \text{const}$  and  $g = \text{const}$  intersect each other (see Figure 13) is obtained from the scalar product of their gradients

$$\cos \mu = \frac{\text{grad } V \cdot \text{grad } g}{|\text{grad } V| |\text{grad } g|} ; \quad (50)$$

in the case of the spheroids, this becomes

$$\cos \bar{\mu} = \frac{\text{grad } \bar{V} \cdot \text{grad } \bar{g}}{|\text{grad } \bar{V}| |\text{grad } \bar{g}|} . \quad (51)$$

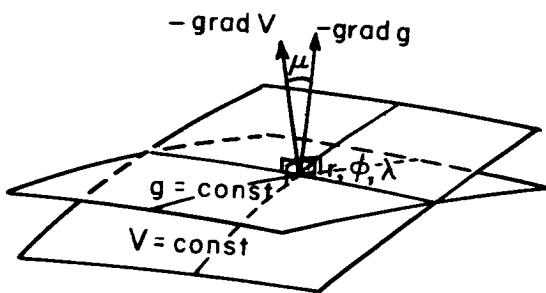


Figure 13.<sup>9</sup>

We determine  $\mu$  stepwise again by

$$\mu = \bar{\mu} + \Delta\mu , \quad (52)$$

wherein  $\Delta\mu$  is found from equation (50) with the help of equations (51) and (52).

It should be pointed out that the angles  $\mu$  are computed along the equipotential surfaces of section I.

---

<sup>9</sup>The negative sign is due to the opposite direction of  $\text{grad } V$  and  $\text{grad } g$  with respect to the surface normals.

Example: The angle  $\mu$  included by the normal of the equipotential surface at 1,000 km elevation and the normal of the equigravitational surface at  $\varphi = 20^\circ$  and  $\lambda = 40^\circ$  is obtained from sections a) and b) below.

$$\bar{\mu} = 1'20''6$$

$$\Delta\mu = -1''2$$

$$\mu = 1'19''4$$

Oscillation of the surface normals:

1. Equipotential surface (see Contour Maps 3 and 4)

$$\xi = -0''9 \quad \eta = 1''6$$

2. Equigravitational surface (see Contour Maps 7 and 8)

$$\xi = -2''1 \quad \eta = 1''6$$

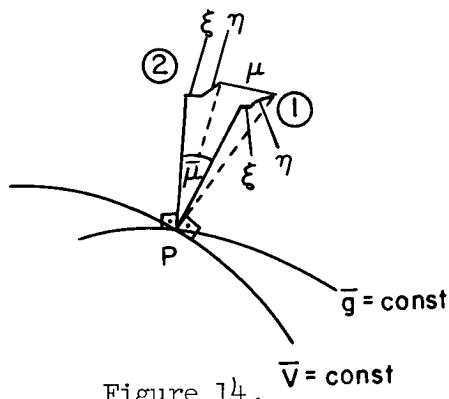


Figure 14.

For the determination of the plane in which  $\mu$  has to be taken, we describe around P a unit sphere and project on its surface the relative position of the normals of the equipotential and equigravitational surface. While the spheroidal part  $\bar{\mu}$  lies always in the meridian, the plane of  $\mu$  is obtained by  $\xi, \eta$  of the equipotential surface 1 and by  $\xi, \eta$  of the equigravitational surface 2, as shown in Figure 14. Because the  $\eta$  values in this example are incidentally both the same,  $\Delta\mu$  is immediately seen to be the difference of the  $\xi$  values 2 and 1.

a) Spheroidal (zonal) part

Table 25.

$\varphi^\circ$	$\Delta\bar{\mu}$	Sea level		Elevation above sea level ( $\omega = 0$ )		
		$\omega \neq 0$	$\omega = 0$	1,000 km	10,000 km	100,000 km
90	0' 0"0	0' 0"0	0' 0"00	0"00	0"000	
80	3' 6"7	0' 55"5	0' 42"21	8"66	0"206	
70	5' 48"6	1' 46"7	1' 20"01	16"28	0"387	
60	7' 48"2	2' 25"6	1' 48"22	21"94	0"521	
50	8' 54"5	2' 44"0	2' 2"85	24"96	0"593	
40	8' 54"4	2' 44"7	2' 2"87	24"98	0"593	
30	7' 51"1	2' 24"1	1' 48"00	21"99	0"521	
20	5' 49"7	1' 47"2	1' 20"65	16"35	0"387	
10	3' 2"3	1' 0"9	0' 43"92	8"73	0"206	
0	0' 0"29	0' 0"29	0' 0"33	0"05	0"000	
-10	3' 5"0	0' 58"2	0' 42"84	8"65	0"206	
-20	5' 51"0	1' 45"9	1' 20"20	16"31	0"387	
-30	7' 48"0	2' 27"2	1' 48"78	22"01	0"521	
-40	8' 54"8	2' 44"2	2' 3"24	25"05	0"593	
-50	8' 54"8	2' 43"7	2' 3"44	25"08	0"593	
-60	7' 44"0	2' 29"8	1' 50"10	22"07	0"522	
-70	5' 45"1	1' 50"2	1' 21"98	16"40	0"387	
-80	3' 5"2	0' 56"9	0' 43"27	8"73	0"206	
-90	0' 0"0	0' 0"0	0' 0"00	0"00	0"000	

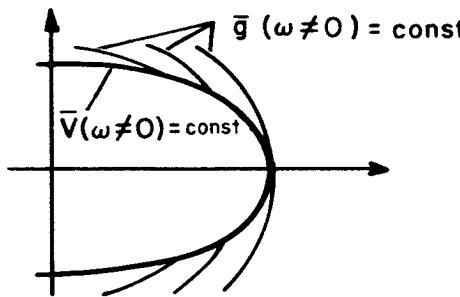


Figure 15.

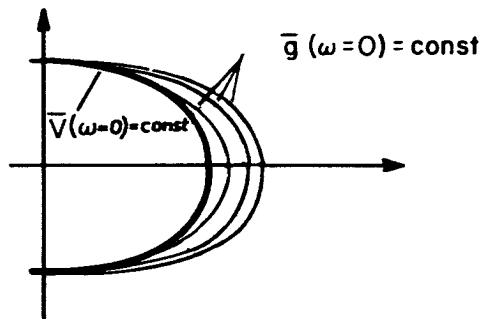


Figure 16.

At sea level the two surfaces  $\bar{g}$  ( $\omega \neq 0$ ) and  $\bar{V}$  ( $\omega \neq 0$ ) intersect each other, as shown in Figure 15. If we ignore the centrifugal force at sea level and higher elevations, Figure 16 reflects the situation for ( $\omega = 0$ ) (see Table 25).

b) Corrections,  $\Delta\mu$

See Number Map 1, page 88 . In the immediate neighborhood of the equator and the poles,  $\Delta\mu$  changes very rapidly with latitude. This is due to the facts that:

1. at the equator the  $\bar{\mu}$  values are very small compared with the  $\mu = \bar{\mu} + \Delta\mu$  values; and
2. at the poles we consider only a point value instead of values along a latitude curve.

In the number maps the geocentric longitude (first horizontal line) has to be multiplied by 10. The value  $\lambda = -5$  actually means  $\lambda = -50^\circ$ , etc. The geocentric latitude is shown in the first vertical line.

B. Direction of the Gradient Field — Orthogonal Trajectories<sup>10</sup> of the Equipotential Surfaces

In the previous sections we were concerned with the geometric structure of surfaces of constant potential or constant gravitation (gravity). Here we shall investigate briefly the characteristics of the curves orthogonally to the equipotential surfaces. Because their tangent is collinear with the gradient of  $V = \text{const}$  (see Figure 17), the differential equation of the orthogonal trajectories is immediately found to be

$$\frac{\vec{dx}}{ds} = - \frac{\text{grad } V}{g} \quad (s = \text{curve length}) . \quad (53)^{11}$$

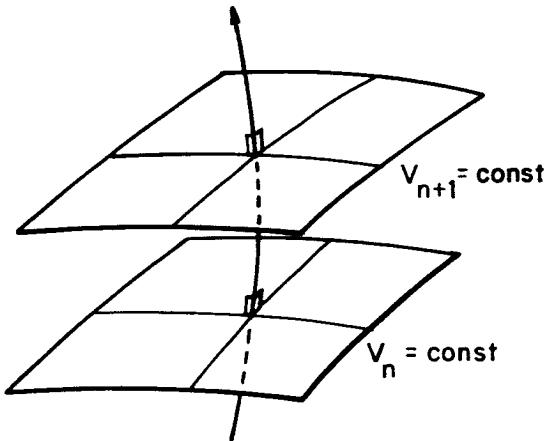


Figure 17.

<sup>10</sup> Also called plumb line, vertical, gradient curve, etc.

<sup>11</sup> The negative sign accounts for the direction of the orthogonal trajectory toward outer space.

1. The gradient (direction) field in a spherical coordinate system

With the help of equation (53) it is possible to determine the angle  $\delta$  included by the geocentric radius vector and the tangent of the orthogonal trajectory at a certain point (see Figure 18):

$$\cos \delta = - \frac{\text{grad } V \cdot \vec{r}}{g \cdot r} . \quad (54)$$

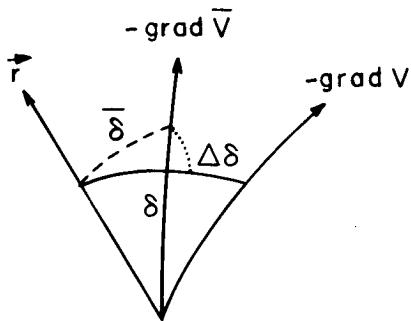


Figure 18.

This angle  $\delta$  represents the structure of the gradient field and the direction of the trajectories. The spheroidal part is obtained from

$$\bar{\delta} = \arcsin \frac{1}{rg} \left| \frac{\partial \bar{V}}{\partial \varphi} \right| , \quad (55)$$

and the correction term  $\Delta\delta = \delta - \bar{\delta}$ :

$$\Delta\delta = \arcsin \frac{1}{rg} \left[ \left( \frac{\partial V}{\partial \varphi} \right)^2 + \frac{1}{\cos^2 \varphi} \left( \frac{\partial V}{\partial \lambda} \right)^2 \right]^{\frac{1}{2}} - \arcsin \frac{1}{rg} \left| \frac{\partial \bar{V}}{\partial \varphi} \right| . \quad (56)$$

On pages 50 and 91 to 92 we give only the magnitudes of  $\bar{\delta}$  and  $\Delta\delta$ . But with the previous results from section I3 we can easily determine the direction in which  $\delta$  has to be taken (see also section II A3).

Example: The angle  $\delta$  included by the geocentric radius and the gradient  $V$  at  $\varphi = 20^\circ$  and  $\lambda = 40^\circ$  in 1,000 km elevation can be found with the help of sections a) and b) below.

$$\begin{array}{rcl} \bar{\delta} & = & 2'41".5 \\ \underline{\Delta\delta} & = & -0".9 \\ \delta & = & 2'40".6 \end{array}$$

To get the direction in which  $\delta$  has to be taken, we proceed similarly to section II A3. The oscillation of the gradient vector is obtained from section I3:

$$\xi = -0".9$$

$$\eta = 1".6 ,$$

which gives, together with  $\bar{\delta}$ , the direction in question (see Figure 19). Although we take, throughout section III, the distance above sea level along the geocentric radius  $r$  as elevations, the  $\xi$ ,  $\eta$  values (which actually refer to an equipotential surface) remain, for practical purposes, unchanged.

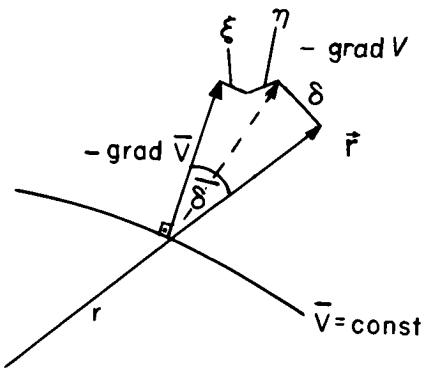


Figure 19.

a) Zonal part of the gradient field

Table 26.

$\phi^\circ$	$\delta$	Sea level		Elevation above sea level ( $\omega = 0$ )		
		$\omega \neq 0$	$\omega = 0$	1,000 km	10,000 km	100,000 km
90	0' 0":0000	0' 0":0000	0' 0":0000	0":0000	0":00000	0":00000
80	3' 55":5234	1' 54":2138	1' 25":5555	17":3829	0":41190	
70	7' 23":5148	3' 35":2924	2' 40":9575	32":6677	0":77410	
60	9' 58":4738	4' 50":5056	3' 36":9644	44":0105	1":04292	
50	11' 21":0097	5' 30":1226	4' 6":6982	50":0454	1":18595	
40	11' 21":8628	5' 30":2499	4' 6":7332	50":0480	1":18594	
30	10' 0":2821	4' 50":4767	3' 37":0957	44":0210	1":04292	
20	7' 26":4583	3' 36":1446	2' 41":4673	32":6912	0":77414	
10	3' 58":8622	1' 56":1838	1' 26":3763	17":4209	0":41200	
0	0' 0":5934	0' 0":5913	0' 0":4717	0":0460	0":00017	
-10	3' 57":3108	1' 54":6377	1' 25":4978	17":3438	0":41172	
-20	7' 25":7662	3' 35":4555	2' 41":0780	32":6546	0":77400	
-30	10' 1":1638	4' 51":3566	3' 37":4094	44":0400	1":04299	
-40	11' 22":6848	5' 31":0699	4' 7":4158	50":1225	1":18622	
-50	11' 22":3416	5' 31":4534	4' 7":7750	50":1608	1":18637	
-60	10' 1":1708	4' 53":2003	3' 38":5202	44":1406	1":04339	
-70	7' 26":1276	3' 37":9031	2' 42":4809	32":7801	0":77450	
-80	3' 56":9709	1' 55":6602	1' 26":4569	17":4481	0":41212	
-90	0' 0":0000	0' 0":0000	0' 0":0000	0":0000	0":00000	

The values at sea level from Table 26 are referred to an equipotential surface  $\bar{V} = \text{const}$  ( $\omega \neq 0$ ) with  $\bar{r} = 6,378,165 \text{ m}$  for  $\varphi = 0$ . Adding the elevation to the geocentric radius, we obtain the angle  $\bar{\delta}$  at higher elevations, and analogously, for section b) we obtain  $\Delta\delta$  by considering  $V = \text{const}$  ( $\omega \neq 0$ ). The angle  $\bar{\delta}$  is, of course, zero at the poles; however, at the equator, we get nonzero results, since the actual equator of an equipotential spheroid does not coincide with the geocentric equator of our coordinate system (see section IIh)). The maximum values of  $\bar{\delta}$  appear as expected in the latitude between  $|40^\circ|$  and  $|50^\circ|$ .

### b) Corrections, $\Delta\delta$

See Number Map 2, page 91. The rapid change of  $\Delta\delta$  near the equator and near the poles is due to the same fact already mentioned in section IIIA3b). In Table 27 we give the range of the  $\Delta\delta$  for different elevations; the polar and equatorial areas are, however, excluded because of the above reasons. Note that the range of the  $\Delta\delta$  values agrees closely with that of the  $\xi$  values of the equipotential surfaces. For explanation see Figure 19.

Table 27.

Elevation above sea level (km)	Range of $\Delta\delta$ (seconds of arc)	
0	5".5	-5".8
1,000	3".0	-3".1
10,000	0".29	-0".23
100,000	0".0039	-0".0037

2. Curvature and torsion of the orthogonal trajectories

The second and third derivatives of equation (53) with respect to s lead to the radii of curvature and of torsion of the orthogonal trajectories: radius of curvature,

$$\rho = \sqrt{\frac{1}{\left(\frac{d^2\vec{x}}{ds^2}\right)^2 + \left(\frac{d^2\vec{x}}{ds^2}\right)^2}}, \quad (57)$$

radius of torsion (attached with a sign),

$$\tau = \frac{1}{\rho^2} \frac{1}{\left(\frac{d\vec{x}}{ds} \frac{d^2\vec{x}}{ds^2} \frac{d^3\vec{x}}{ds^3}\right)}. \quad (58)$$

If we define a local coordinate system  $y^1$ , which coincides with the moving trihedral of the trajectory, then the canonical representation of an orthogonal trajectory can be expressed (see also Kreyszig, 1959):

$$\begin{aligned} y^1 &= s - \frac{1}{6\rho^2} s^3 + \frac{\rho'}{8\rho^3} s^4 + \frac{s^5}{120\rho^2} \left[ \frac{1}{\rho^2} + \frac{1}{\tau^2} + \frac{4\rho''}{\rho} - \frac{11(\rho')^2}{\rho^2} \right] + \dots, \\ y^2 &= \frac{s^2}{2\rho} - \frac{\rho'}{6\rho^2} s^3 + \frac{s^4}{24\rho} \left[ \frac{2(\rho')^2 - 1}{\rho^2} - \frac{1}{\tau^2} - \frac{\rho''}{\rho} \right] \\ &\quad + \frac{s^5}{120\rho} \left[ \frac{3\tau'}{\tau^3} + \frac{3\rho'}{\rho\tau^2} + \frac{6}{\rho^3} \rho' - \frac{6(\rho')^3}{\rho^3} + \frac{6\rho'\rho''}{\rho^2} - \frac{\rho'''}{\rho} \right] + \dots, \end{aligned} \quad (59)$$

$$y^3 = \frac{s^3}{6\rho\tau} - \frac{s^4}{24\rho\tau} \left[ \frac{2\rho'}{\rho} + \frac{\tau'}{\tau} \right] \\ + \frac{s^5}{120\rho\tau} \left[ -\frac{1}{\rho^2} - \frac{1}{\tau^2} + \frac{6(\rho')^2}{\rho^2} + \frac{3\rho'\tau'}{\rho\tau} + \frac{2(\tau')^2}{\tau^2} - \frac{3\rho''}{\rho} - \frac{\tau''}{\tau} \right] + \dots,$$

wherein

$$\frac{d\rho}{ds} \equiv \rho', \quad \frac{d^2\rho}{ds^2} \equiv \rho'' \text{ etc., and similar for } \tau .$$

With a positive  $\rho$  the trajectory under consideration behaves for  $\tau > 0$  like a right-hand screw and for  $\tau < 0$  like a left-hand screw. The tangent  $\vec{t}$ , the normal  $\vec{n}$ , and the binormal  $\vec{b}$  have in this case the following components in the above-defined local cartesian coordinate system:

$$\vec{t} = (1, 0, 0)$$

$$\vec{n} = (0, 1, 0)$$

$$\vec{b} = (0, 0, 1) .$$

The derivatives of  $\vec{x}$  with respect to the curve length  $s$  were computed partly from analytical expressions and partly by numerical differentiation of the integrated trajectory (see the next section) in a spherical coordinate system.

a) Radius of curvature of the orthogonal trajectories

1) Zonal part, i.e., radius of curvature of the orthogonal trajectories of the equipotential spheroids  $\bar{V} = \text{const}$

Table 28.

$\varphi^\circ$	Sea level		Elevation above sea level					
	$\omega \neq 0$	$\omega = 0$	10 km ( $\omega \neq 0$ )	10 km ( $\omega = 0$ )	100 km ( $\omega = 0$ )	1,000 km ( $\omega = 0$ )	10,000 km ( $\omega = 0$ )	100,000 km ( $\omega = 0$ )
90	$\infty$	$\infty$	$\infty$	$\infty$	$\infty$	$\infty$	$\infty$	$\infty$
80	3520 E 3*	1182 E 4	3501 E 3	1188 E 4	1236 E 4	1797 E 4	1945 E 5	5328 E 7
70	1883 E 3	6145 E 3	1873 E 3	6174 E 3	6434 E 3	9476 E 3	1035 E 5	2835 E 7
60	1399 E 3	4500 E 3	1392 E 3	4522 E 3	4720 E 3	7006 E 3	7680 E 4	2104 E 7
50	1222 E 3	3996 E 3	1216 E 3	4015 E 3	4185 E 3	6172 E 3	6753 E 4	1850 E 7
40	1219 E 3	3979 E 3	1213 E 3	3998 E 3	4171 E 3	6172 E 3	6751 E 4	1850 E 7
30	1379 E 3	4548 E 3	1372 E 3	4569 E 3	4764 E 3	7022 E 3	7673 E 4	2104 E 7
20	1854 E 3	6113 E 3	1844 E 3	6141 E 3	6396 E 3	9405 E 3	1032 E 5	2834 E 7
10	3551 E 3	1076 E 4	3530 E 3	1082 E 4	1135 E 4	1727 E 4	1934 E 5	5324 E 7
0	2230 E 6	2269 E 6	2369 E 6	2413 E 6	4844 E 6	2303 E 6	3705 E 7	6467 E 10
-10	3501 E 3	1125 E 4	3481 E 3	1131 E 4	1182 E 4	1771 E 4	1951 E 5	5331 E 7
-20	1848 E 3	6188 E 3	1838 E 3	6215 E 3	6466 E 3	9458 E 3	1035 E 5	2835 E 7
-30	1389 E 3	4451 E 3	1381 E 3	4473 E 3	4674 E 3	6973 E 3	7666 E 4	2104 E 7
-40	1218 E 3	3991 E 3	1212 E 3	4010 E 3	4179 E 3	6153 E 3	6732 E 4	1849 E 7
-50	1222 E 3	4005 E 3	1215 E 3	4023 E 3	4189 E 3	6143 E 3	6722 E 4	1849 E 7
-60	1412 E 3	4371 E 3	1404 E 3	4393 E 3	4594 E 3	6885 E 3	7634 E 4	2102 E 7
-70	1902 E 3	5949 E 3	1892 E 3	5977 E 3	6233 E 3	9245 E 3	1027 E 5	2832 E 7
-80	3548 E 3	1152 E 4	3529 E 3	1157 E 4	1204 E 4	1752 E 4	1929 E 5	5322 E 7
-90	$\infty$	$\infty$	$\infty$	$\infty$	$\infty$	$\infty$	$\infty$	$\infty$

\* E 3 means  $\times 10^3$ , etc.

The radius of curvature  $\bar{p}$  at the equator increases from sea level ( $\omega = 0$ ) up to 100 km ( $\omega = 0$ ), then drops at 1000 km height ( $\omega = 0$ ) and starts increasing again (refer to Table 28). This behavior is due to the fact that the actual equator of the equipotential spheroids moves at first southward and finally touches the geocentric equator asymptotically with higher elevation (see section IIh)). The values  $\bar{p}$  at sea level were computed on the spheroidal part of the Earth's surface. Elevation 10 km, for example, means 10 km above sea level along the geocentric radius.

- 2) Zonal and tesseral (sectorial) part, i.e., radius of curvature of the orthogonal trajectories of the equipotential surfaces  $V = \text{const}$

In Number Map 3, the radius of curvature is given as a function of the latitude (first vertical line) and longitude (first horizontal line) at a certain elevation above sea level. As easily recognized, each geocentric latitude and longitude number has to be multiplied by 10. Further,  $\pm 90^\circ$  in latitude actually means  $\pm 89^\circ 875$ , which had to be taken because the computer program does not work at  $\pm 90^\circ$ , for mathematical reasons, since spherical coordinates were used; a reprogramming hardly seemed worthwhile. The first four digits are significant numbers of the curvature radius, and the last number is a power to the ten, which has to be added to the scaling power to get the radius in kilometers.

Example: The radius of curvature is at sea level for  $\omega = 0$  and  $\varphi = 20^\circ$ ,  $\lambda = 40^\circ$

$$r = 6\ 229 \times 10^3 \text{ km} .$$

b) Radius of torsion of the orthogonal trajectories

See Number Map 4. Similar to the radius of curvature, we give the radius of torsion for different sections of outer space as a function of the geocentric latitude and longitude. Again the latitude  $\pm 90^\circ$  actually stands for  $\pm 89^\circ 875$ , as stated previously. The first two digits are significant numbers of the radius of torsion, and the number behind the exponent E has to be added to the scaling power in order to get the radius of torsion in kilometers.

Example: The radius of torsion at 1,000 km elevation for  $\omega = 0$  and  $\varphi = 20^\circ$ ,  $\lambda = 40^\circ$

$$\tau = - 11 \times 10^4 \text{ km} .$$

The sign was attached to the radius for reasons of simplicity. Actually it means that the torsion is negative and has a radius of  $11 \times 10^4$  km.

The range of the radius of torsion can go up to infinity. In the computer programs, however, only a relatively small spectrum could be recorded. If we have at a certain latitude, for example, a change in the exponent of magnitude three or larger, then the numerical results of the absolute larger exponent cannot be guaranteed, because of rounding errors.

• 3. Integration of the differential equation; orthogonal trajectories

The orthogonal trajectories can be obtained pointwise by integrating the differential equation system (53) by a Runge-Kutta procedure. If we apply a curve fitting, we get the trajectory as a function of the arc length  $s$  starting from an initial point on the Earth's surface.

a) Gravity field ( $\bar{V} + \bar{Z}$ ); zonal part

We consider for the moment only the zonal part of the geopotential. By using a step length of 2.5 km up to an elevation of 10 km, we represent the orthogonal trajectory as a function of the arc length:

$$\begin{aligned}\bar{r} &= \bar{r}(s) , \\ \varphi &= \varphi(s) ;\end{aligned}\tag{60}$$

or, applying an interpolation formula,

$$\begin{aligned}\bar{r} &= \bar{r}_0 + \sum_{q=1}^p \binom{p}{q} \bar{r}_q , \\ \varphi &= \varphi_0 + \sum_{q=1}^p \binom{p}{q} \bar{\varphi}_q ,\end{aligned}\tag{61}$$

wherein

$\bar{r}_0, \varphi_0$  = initial points of the numerical integration (not exactly coinciding with sea level because of the way the computer programs were set up. See also p. 59 and p. 62).

$p = \frac{s}{\Delta}$ , with  $\Delta$  the step length of the integration process. The values  $s$  and  $\Delta$  must have the same length unit.

Table 29.<sup>12</sup>

Initial points		Coefficients *					
$\varphi_0^\circ$	$\bar{r}_0$ (m)	$\bar{r}_1$ (m)	$\bar{r}_2$ (m)	$\bar{\varphi}_1$	$\bar{\varphi}_2$	$\bar{\varphi}_3$	
90	--	--	--	--	--	--	--
80	6 357 432.285 872 9	2 499.998 369 9	-0.7 E-6	0.448 984 3 E-6	-0.734 E-10	0.5 E-12	
70	6 359 277.419 177 4	2 499.994 219 4	-0.26 E-5	0.845 236 1 E-6	-0.141 8 E-9	0.9 E-12	
60	6 362 110.511 335 3	2 499.989 474 4	-0.47 E-5	0.114 004 26 E-5	-0.192 8 E-9	0.12 E-11	
50	6 365 590.836 372 7	2 499.986 371 0	-0.62 E-5	0.129 655 931 E-5	-0.213 98 E-9	0.136 E-11	
40	6 369 299.878 656 8	2 499.986 336 8	-0.62 E-5	0.129 742 848 E-5	-0.212 49 E-9	0.137 E-11	
30	6 372 791.523 500 6	2 499.989 410 6	-0.49 E-5	0.114 157 545 E-5	-0.183 59 E-9	0.120 E-11	
20	6 375 642.854 393 7	2 499.994 142 4	-0.27 E-5	0.848 664 87 E-6	-0.136 09 E-9	0.89 E-12	
10	6 377 511.144 032 6	2 499.998 323 3	-0.7 E-6	0.453 912 94 E-6	-0.794 4 E-10	0.51 E-12	
0	6 378 164.510 543 6	2 500 0		0.112 732 E-8	-0.45 E-12	0	
-10	6 377 518.386 571 7	2 499.998 345 0	-0.8 E-6	-0.450 967 25 E-6	0.731 6 E-10	-0.49 E-12	
-20	6 375 657.063 778 4	2 499.994 160 5	-0.27 E-5	-0.847 348 73 E-6	0.133 12 E-9	-0.88 E-12	
-30	6 372 804.700 009 8	2 499.989 379 6	-0.48 E-5	-0.114 324 701 E-5	0.189 60 E-9	-0.123 E-11	
-40	6 369 307.994 930 5	2 499.986 303 9	-0.63 E-5	-0.129 899 059 E-5	0.212 97 E-9	-0.136 E-11	
-50	6 365 593.939 136 6	2 499.986 317 6	-0.62 E-5	-0.129 909 398 E-5	0.215 43 E-9	-0.135 E-11	
-60	6 362 102.623 656 6	2 499.989 379 4	-0.46 E-5	-0.114 517 66 E-5	0.203 3 E-9	-0.12 E-11	
-70	6 359 254.691 583 3	2 499.994 151 1	-0.25 E-5	-0.850 214 4 E-6	0.151 0 E-9	-0.9 E-12	
-80	6 357 398.595 474 8	2 499.998 349 8	-0.7 E-6	-0.451 744 1 E-6	0.778 E-10	-0.5 E-12	
-90	--	--	--	--	--	--	

\*  
 $\bar{\varphi}_1$ ,  $\bar{\varphi}_2$ ,  $\bar{\varphi}_3$  are in radians.

The coefficients  $\bar{r}_q$  and  $\bar{\varphi}_q$  of Table 29 are valid from sea level (spheroidal part) up to 10 km elevation. Within this range the trajectory can be presented with a numerical accuracy of about 13-14 digits. A short analysis shows that the orthogonal trajectory is curved toward the poles within the above range  $L = 10$  km. The opposite result is obtained if  $\omega = 0$  (see page 62).

b) Gravity field ( $V + Z$ ); zonal and tesseral part

See Number Map 5 pages 99-104. Including the tesseral (sectorial) harmonics

<sup>12</sup>The steplength  $\Delta$  is 2.5 km. For a numerical example see next section.

our integration process, the orthogonal trajectory has the equations

$$r = r(s) ,$$

$$\varphi = \varphi(s) ,$$

$$\lambda = \lambda(s) .$$

We fit the point array obtained by the Runge-Kutta procedure again with an interpolation formula:

$$r = r_0 + \sum_{q=1}^p \binom{p}{q} r_q ,$$

$$\varphi = \varphi_0 + \sum_{q=1}^p \binom{p}{q} \varphi_q ,$$

$$\lambda = \lambda_0 + \sum_{q=1}^p \binom{p}{q} \lambda_q ,$$

wherein

$r_0, \varphi_0, \lambda_0$  are initial points of the numerical integration with  $p$  defined as in equation (61).

Hence with the tables on page 99-104 the trajectory is fully determined in its whole run from sea level up to 10 km elevation with an accuracy of 13 significant digits.

Explanation of the number maps: The trajectory was computed  $10 \times 10^\circ$  in  $\varphi$  and  $\lambda$  all over the Earth's surface. At the North and South Poles,  $89^\circ 875$  in  $\varphi$  was taken instead of  $90^\circ$ ; the step length of the numerical integration was  $\Delta = 2.5$  km.

$r_q$  tables: The first column and the first row are  $\varphi/10$  and  $\lambda/10$  of the trajectory under consideration. The set of two or three numbers to be found expresses  $r_0, r_1, r_2$ , whereby in front of  $r_0$  (larger number) the digits 63 had to be omitted because of space problems in the printout. If we like meters as length units in  $r$ , then  $r_0, r_1, r_2$  have to be multiplied by  $10^{-6}$ .

$\varphi_q, \lambda_q$  tables: The three differences  $\varphi_1, \varphi_2, \varphi_3$  and  $\lambda_1, \lambda_2, \lambda_3$ , respectively, have to be multiplied by  $10^{-13}$  to get radians.

Example: The equation of the plumb line at  $\varphi_0 = 20^\circ$ ,  $\lambda_0 = 40^\circ$  starting from sea level ( $V = \text{const}, \omega \neq 0$ ):

$$r = r_0 + p r_1 + \frac{p(p-1)}{2} r_2 = 6,375,657.361\ 403$$

$$+ 2,499.994\ 183\ p - 0.000\ 003\ \frac{p(p-1)}{2} ,$$

$$\varphi = \varphi_0 + p \varphi_1 + \frac{p(p-1)}{2} \varphi_2 + \frac{p(p-1)(p-2)}{6} \varphi_3$$

$$= 20^\circ + p^\circ \left( 0.000\ 000\ 845\ 730\ 2\ p \right.$$

$$- 0.000\ 000\ 000\ 130\ 8\ \frac{p(p-1)}{2} ,$$

$$\left. + 0.000\ 000\ 000\ 000\ 9\ \frac{p(p-1)(p-2)}{6} \right)$$

$$\lambda = \lambda_0 + p \lambda_1 + \frac{p(p-1)}{2} \lambda_2$$

$$= 40^\circ + p^\circ \left( 0.000\ 000\ 003\ 274\ 4\ p \right.$$

$$\left. - 0.000\ 000\ 000\ 000\ 4\ \frac{p(p-1)}{2} \right)$$

---

$13^\circ$  is the conversion factor from radians to degrees.

At 7 km above sea level the spherical coordinates  $r$ ,  $\varphi$ ,  $\lambda$  of the trajectory assume the following values:

With

$$p = \frac{7}{2.5} = 2.8 ,$$

$$\frac{p(p-1)}{2} = 2.52 ,$$

$$\frac{p(p-1)(p-2)}{6} = 0.672 ,$$

and

$$\rho^\circ = 57.295\ 779\ 51 ,$$

we get

$$\begin{aligned} r_7 \text{ km} &= 6,375,657.361\ 403 \\ &\quad + 6,999.983\ 712 \\ &\quad - .000\ 008 \\ &= \underline{\underline{6,382,657.345\ 107}} \text{ m} , \end{aligned}$$

$$\begin{aligned} \varphi_7 \text{ km} &= 20^\circ + \rho^\circ (0.000\ 002\ 368\ 044\ 6 \\ &\quad - .000\ 000\ 000\ 329\ 6 \\ &\quad + .000\ 000\ 000\ 000\ 6) \\ &= 20^\circ 00' 135\ 660\ 111\ 0 \\ &= 20^\circ 00' 00'' 488\ 376\ 40 , \end{aligned}$$

$$\begin{aligned} \lambda_7 \text{ km} &= 40^\circ + \rho^\circ (0.000\ 000\ 009\ 168\ 3 \\ &\quad - .000\ 000\ 000\ 001\ 0) \end{aligned}$$

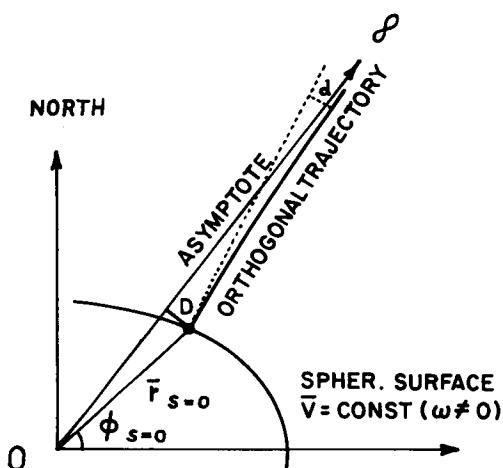
$$\begin{aligned} &= 40^\circ 00' 000\ 525\ 247\ 6 \\ &= 40^\circ 00' 00'' 001\ 890\ 9 . \end{aligned}$$

c) Gravitational field ( $\bar{V}, \bar{Z} = 0$ ); zonal part

Quite similarly the zonal part of the gravitational field is treated. Starting from the Earth's spheroidal surface, the differential equations (53) are integrated up to 100,000 km elevation; the geocentric radius  $\bar{r}$ , the geocentric latitude  $\varphi$ , and the distance  $d$  of the trajectory from the tangent of the starting point are given pointwise as functions of the curve length (see Tables 31, 32, and 33). Each trajectory approaches asymptotically a certain radius vector. The distance  $D$  of the starting point from this line is shown in Table 30.

Table 30.

$\varphi^\circ$	D (km)
90	0
80	1.76
70	3.31
60	4.47
50	5.08
40	5.08
30	4.47
20	3.33
10	1.78
0	0.00755
-10	1.76
-20	3.32
-30	4.48
-40	5.09
-50	5.10
-60	4.49
-70	3.34
-80	1.78
-90	0



Note that the orthogonal trajectory refers to  $\bar{V} = \text{const}$  ( $\omega = 0$ )

Figure 20.

Compared to a straight line (Figure 20), the orthogonal trajectories vary most within the range from the Earth's surface up to about 10,000 km elevation, as can be seen easily from Table 32 on page 66.

Example: At  $\varphi = 40^\circ$  the geocentric latitude changes from sea level up to 10,000 km by  $2'20\text{"}2$ , while for the nine-times-longer range from 10,000 km up to 100,000 km, the change is only  $24\text{"}4$ , and so forth.

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Table 31.--The geocentric radius  $\bar{r}$  of the trajectories, given in meters (see page 62).

$s \text{ km} \times 10^3$	$\Psi = 90^\circ$	$\Psi = 80^\circ$	$\Psi = 70^\circ$	$\Psi = 60^\circ$	$\Psi = 50^\circ$	$\Psi = 40^\circ$	$\Psi = 30^\circ$
0	6 357 432.285 9	6 359 277.419 2	6 362 110.511 3	6 365 590.836 4	6 369 299.878 7	6 372 791.523 5	
1	7 357 452.170 2	7 359 276.009 1	7 362 109.765 6	7 365 589.872 5	7 369 298.914 2	7 372 790.777 0	
4	10 357 452.034 7	10 362 109.765 6	10 365 589.716 0	10 369 297.866 9	10 372 789.538 8	10 372 790.777 0	
10	16 357 451.978 0	16 359 276.329 5	16 362 108.984 8	16 365 588.275 2	16 369 297.315 6	16 372 789.444 1	
20	26 357 451.963 3	26 359 276.277 5	26 362 108.536 6	26 365 588.153 1	26 369 297.193 3	26 372 789.425 7	
30	36 357 451.960 4	36 362 108.418 2	46 362 108.412 4	46 365 588.121 7	46 369 297.161 9	46 372 789.417 4	
40	46 357 451.959 5	46 359 276.264 2	56 362 108.410 0	56 365 588.118 6	56 369 297.158 8	56 372 789.417 4	
50	56 357 451.959 1	56 359 276.262 9	56 362 108.410 0	56 365 588.118 6	56 369 297.158 8	56 372 789.417 4	
60	66 357 451.958 9	66 359 276.262 2	66 362 108.408 8	66 365 588.117 1	66 369 297.157 3	66 372 789.416 2	
70	76 357 451.958 8	76 359 276.261 9	76 362 108.408 1	76 365 588.116 2	76 369 297.156 4	76 372 789.415 6	
80	86 357 451.958 8	86 359 276.261 7	86 362 108.407 8	86 365 588.115 7	86 369 297.156 0	86 372 789.415 2	
90	96 357 451.958 8	96 359 276.261 5	96 362 108.407 5	96 365 588.115 4	96 369 297.155 6	96 372 789.415 0	
100	106 357 451.958 7	106 359 276.261 4	106 362 108.407 4	106 365 588.115 2	106 369 297.155 4	106 372 789.414 8	
$s \text{ km} \times 10^3$	$\Psi = 20^\circ$	$\Psi = 10^\circ$	$\Psi = 0^\circ$	$\Psi = -10^\circ$	$\Psi = -20^\circ$	$\Psi = -30^\circ$	$\Psi = -40^\circ$
0	6 375 602.854 4	6 377 511.144 0	6 378 164.510 5401	6 377 518.386 6	6 375 657.063 8	6 372 804.700 0	6 369 307.994 9
1	7 375 602.441 1	7 377 511.025 3	7 378 164.510 5401	7 377 518.270 6	7 375 656.692 8	7 372 803.950 3	7 369 307.025 3
4	10 375 641.958 4	10 377 510.887 5	10 378 164.510 537	10 377 518.135 2	10 375 656.172 1	10 372 803.075 4	10 369 305.892 1
10	16 375 641.756 9	16 377 510.830 2	16 378 164.510 536	16 377 518.078 5	16 375 655.971 2	16 372 802.710 0	16 369 305.419 0
20	26 375 641.704 7	26 377 510.815 4	26 378 164.510 536	26 377 518.063 8	26 375 655.919 0	26 372 802.615 2	26 369 305.296 4
30	36 375 641.694 5	36 377 510.812 5	36 378 164.510 536	36 377 518.060 9	36 375 655.908 9	36 372 802.596 8	36 369 305.272 6
40	46 375 641.694 3	46 377 510.811 5	46 378 164.510 536	46 377 518.060 0	46 375 655.905 6	46 372 802.590 9	46 369 305.265 0
50	56 375 641.689 9	56 377 510.811 2	56 378 164.510 536	56 377 518.059 6	56 375 655.904 3	56 372 802.588 5	56 369 305.261 9
60	66 375 641.689 3	66 377 510.811 0	66 378 164.510 536	66 377 518.059 4	66 375 655.903 6	65 372 802.587 3	66 369 305.260 3
70	76 375 641.688 9	76 377 510.810 9	76 378 164.510 535	76 377 518.059 3	76 375 655.903 3	76 372 802.586 6	76 369 305.259 5
80	86 375 641.688 7	86 377 510.810 8	86 378 164.510 535	86 377 518.059 3	86 375 655.903 1	85 372 802.586 3	86 369 305.259 0
90	96 375 641.688 2	96 377 510.810 8	96 378 164.510 535	96 377 518.059 2	96 375 655.902 9	95 372 802.586 0	96 369 305.259 7
100	106 375 641.688 5	106 377 510.810 8	106 378 164.510 535	106 377 518.059 2	106 375 655.902 9	106 372 802.585 9	106 369 305.258 5
$s \text{ km} \times 10^3$	$\Psi = -50^\circ$	$\Psi = -60^\circ$	$\Psi = -70^\circ$	$\Psi = -80^\circ$	$\Psi = -90^\circ$		
0	6 365 599.999 1	6 362 102.623 7	6 359 254.691 6	6 357 398.595 5			
1	7 365 592.967 0	7 362 101.865 6	7 359 254.272 6	7 357 398.477 1			
4	10 365 591.891 3	10 362 100.984 3	10 359 253.785 9	10 357 398.339 2			
10	16 365 591.357 6	16 362 100.617 6	16 359 253.593 7	16 357 398.281 9			
20	26 365 591.225 0	26 362 100.522 8	26 359 253.521 4	26 357 398.267 1			
30	36 365 591.211 2	36 362 100.504 3	36 359 253.521 3	36 357 398.264 2			
40	46 365 591.203 6	46 362 100.498 5	46 359 253.518 0	46 357 398.263 3			
50	56 365 591.200 4	56 362 100.496 0	56 359 253.516 7	56 357 398.263 0			
60	66 365 591.198 9	66 362 100.494 9	66 359 253.516 0	66 357 398.262 8			
70	76 365 591.198 1	76 362 100.494 2	76 359 253.515 7	76 357 398.262 7			
80	86 365 591.197 6	86 362 100.493 8	86 359 253.515 5	86 357 398.262 6			
90	96 365 591.197 3	96 362 100.493 6	96 359 253.515 3	96 357 398.262 6			
100	106 365 591.197 1	106 362 100.493 5	106 359 253.515 3	106 357 398.262 5			

$\Delta r = \Delta r$

Table 32. --The geocentric latitude  $\Phi$  of the trajectories, given in radians.

		$\Phi = 90^\circ$	$\Phi = 80^\circ$	$\Phi = 70^\circ$	$\Phi = 60^\circ$	$\Phi = 50^\circ$	$\Phi = 40^\circ$	$\Phi = 30^\circ$
$\frac{\pi}{2}$	0	1.396 263 40	1.221 730 48	1.047 197 55	0.872 664 63	0.698 151 70	0.523 598 78	
	1	1.396 333 66	1.221 862 73	1.047 375 84	0.872 867 22	0.698 324 24	0.523 776 88	
	4	1.396 436 49	1.222 056 05	1.047 636 21	0.873 163 35	0.698 630 36	0.524 037 38	
	10	1.396 499 35	1.222 774 16	1.047 795 44	0.873 344 53	0.698 811 38	0.524 196 64	
	20	1.396 525 21	1.222 222 81	1.047 861 00	0.873 418 90	0.698 885 97	0.524 262 26	
	30	1.396 532 91	1.222 237 27	1.047 880 50	0.873 441 08	0.698 908 16	0.524 281 79	
	40	1.396 536 19	1.222 243 44	1.047 888 81	0.873 450 53	0.698 917 62	0.524 290 11	
	50	1.396 537 89	1.222 246 63	1.047 892 11	0.873 455 42	0.698 922 52	0.524 294 42	
	60	1.396 538 88	1.222 248 49	1.047 895 62	0.873 458 28	0.698 925 37	0.524 296 93	
	70	1.396 539 50	1.222 249 67	1.047 897 21	0.873 460 08	0.698 927 18	0.524 298 52	
80	80	1.396 539 92	1.222 250 46	1.047 898 27	0.873 461 30	0.698 928 39	0.524 299 59	
	90	1.396 540 22	1.222 251 02	1.047 899 05	0.873 462 16	0.698 929 25	0.524 300 34	
	100	1.396 540 44	1.222 251 43	1.047 899 58	0.873 462 78	0.698 929 88	0.524 300 89	
		$\Phi = 20^\circ$	$\Phi = 10^\circ$	$\Phi = 0^\circ$	$\Phi = -10^\circ$	$\Phi = -20^\circ$	$\Phi = -30^\circ$	$\Phi = -40^\circ$
$\frac{\pi}{2}$	0	0.349 065 85	0.174 532 925	0.0	-0.174 532 925	-0.349 065 85	-0.523 598 78	-0.698 131 70
	1	0.349 198 31	0.174 603 913	0.384 1 E-6	-0.174 603 089	-0.349 197 94	-0.523 776 26	-0.698 324 78
	4	0.349 391 95	0.174 707 315	0.8917 E-6	-0.174 705 609	-0.349 391 16	-0.524 038 01	-0.698 631 67
	10	0.349 510 24	0.174 770 400	1.1089 E-6	-0.174 768 330	-0.349 509 30	-0.524 197 36	-0.698 813 03
	20	0.349 558 98	0.174 796 373	1.1657 E-6	-0.174 794 207	-0.349 528 00	-0.524 263 01	-0.698 887 12
	30	0.349 573 48	0.174 804 094	1.1768 E-6	-0.174 801 910	-0.349 572 19	-0.524 282 54	-0.698 909 92
	40	0.349 579 66	0.174 807 387	1.1804 E-6	-0.174 805 197	-0.349 578 67	-0.524 290 86	-0.698 919 39
	50	0.349 582 86	0.174 809 089	1.1819 E-6	-0.174 806 897	-0.349 581 86	-0.524 295 17	-0.698 924 29
	60	0.349 584 72	0.174 810 082	1.1826 E-6	-0.174 807 888	-0.349 583 73	-0.524 297 68	-0.698 927 14
	70	0.349 585 90	0.174 810 710	1.1830 E-6	-0.174 808 516	-0.349 584 91	-0.524 299 27	-0.698 928 95
80	80	0.349 586 70	0.174 811 154	1.1832 E-6	-0.174 808 939	-0.349 585 70	-0.524 300 34	-0.698 930 17
	90	0.349 587 26	0.174 811 432	1.1834 E-6	-0.174 809 237	-0.349 586 26	-0.524 301 10	-0.698 931 03
	100	0.349 587 67	0.174 811 650	1.1835 E-6	-0.174 809 455	-0.349 586 67	-0.524 301 65	-0.698 931 65
		$\Phi = -56^\circ$	$\Phi = -60^\circ$	$\Phi = -64^\circ$	$\Phi = -70^\circ$	$\Phi = -80^\circ$	$\Phi = -90^\circ$	
$\frac{\pi}{2}$	0	-0.872 664 63	-1.047 197 55	-1.221 730 48	-1.396 263 40	-1.396 334 48	-1.396 403 48	
	1	-0.872 868 08	-1.047 377 30	-1.221 864 16	-1.396 334 48	-1.396 403 48	-1.396 473 19	
	4	-0.872 165 42	-1.047 639 52	-1.222 058 95	-1.396 501 35	-1.396 501 35	-1.396 571 31	
	10	-0.873 346 94	-1.047 799 08	-1.222 177 61	-1.396 527 31	-1.396 527 31	-1.396 595 19	
	20	-0.873 121 65	-1.047 864 79	-1.222 226 39	-1.396 535 02	-1.396 535 02	-1.396 603 19	
	30	-0.873 443 86	-1.047 884 32	-1.222 240 89	-1.396 538 31	-1.396 538 31	-1.396 603 19	
	40	-0.873 453 33	-1.047 892 64	-1.222 247 06	-1.396 540 01	-1.396 540 01	-1.396 603 19	
	50	-0.873 458 22	-1.047 896 95	-1.222 250 26	-1.396 542 01	-1.396 542 01	-1.396 603 19	
	60	-0.873 161 07	-1.047 899 45	-1.222 252 12	-1.396 544 01	-1.396 544 01	-1.396 603 19	
	70	-0.873 462 88	-1.047 901 04	-1.222 253 30	-1.396 546 01	-1.396 546 01	-1.396 603 19	
$\frac{1}{2}$	80	-0.873 464 10	-1.047 902 11	-1.222 254 09	-1.396 548 01	-1.396 548 01	-1.396 603 19	
	90	-0.873 464 96	-1.047 902 87	-1.222 254 65	-1.396 550 01	-1.396 550 01	-1.396 603 19	
	100	-0.873 465 58	-1.047 903 42	-1.222 255 06	-1.396 552 57	-1.396 552 57	-1.396 603 19	

Table 33. --The distance  $d$  of the trajectory from the tangent in the initial point, given in km.

$s \text{ km}$ $\times 10^3$	$\varphi = 90^\circ$	$\varphi = 80^\circ$	$\varphi = 70^\circ$	$\varphi = 60^\circ$	$\varphi = 50^\circ$	$\varphi = 40^\circ$	$\varphi = 30^\circ$	$\varphi = 20^\circ$
0	0.000 0	0.000 0	0.000 0	0.000 0	0.000 0	0.000 0	0.000 0	0.000 0
1	0.036 8	0.070 4	0.095 8	0.108 3	0.108 5	0.095 1	0.070 9	0.070 9
4	0.422 2	0.862 4	1.087 2	1.232 4	1.233 7	1.083 5	0.863 5	0.863 5
10	1.678 1	3.179 4	4.301 3	4.881 1	4.885 1	4.294 0	3.201 9	3.201 9
20	0	6.173 8	7.897 8	10.678 3	12.122 8	12.132 3	10.667 3	10.667 3
30	0	6.813 1	12.886 2	17.418 9	19.778 3	19.793 6	17.105 1	17.105 1
40	9.503 2	17.969 8	24.288 9	27.580 2	27.601	24.272	18.088	18.088
50	12.216 9	23.098 1	31.217 3	35.150	35.477	31.198	23.248	23.248
60	14.943 6	28.250 8	38.179	43.358	43.391	38.158	28.434	28.434
70	17.678 3	33.418	45.162	51.288	51.327	45.138	33.624	33.624
80	20.418 0	38.596	52.157	59.233	59.278	52.131	38.843	38.843
90	23.161 4	43.779	59.161	67.189	67.239	59.132	44.060	44.060
100	25.907 2	48.968	66.172	75.151	75.208	66.141	49.281	49.281

$s \text{ km}$ $\times 10^3$	$\varphi = 10^\circ$	$\varphi = 0^\circ$	$\varphi = -10^\circ$	$\varphi = -20^\circ$	$\varphi = -30^\circ$	$\varphi = -40^\circ$	$\varphi = -50^\circ$	$\varphi = -60^\circ$
0	0.000 0	0.000 0	0.000 0	0.000 0	0.000 0	0.000 0	0.000 0	0.000 0
1	0.039 6	0.070 3	0.098 1	0.070 3	0.096 6	0.108 5	0.108 4	0.098 2
4	0.443 4	0.002 21	0.431 1	0.802 9	1.094 0	1.236 0	1.236 7	1.168 3
10	1.743 5	0.010 51	1.702 4	3.183 8	4.344 9	4.897 8	4.902 8	4.372 5
20	4.316 4	0.026 59	4.223 6	7.910 4	10.733 0	12.165 8	12.179 2	10.839 6
30	7.033 8	0.043 19	6.888 4	12.907 4	17.505 8	19.848 7	19.870 7	17.671 9
40	9.002 1	0.059 92	9.603 9	18.000	24.107	27.678	27.709	21.633 0
50	12.594 3	0.076 71	12.343 1	23.137	31.369	35.576	35.615	31.654 4
60	15.399 5	0.093 51	15.095 4	26.299	38.364	43.511	43.559	38.709
70	18.212 7	0.110 31	17.855 6	33.476	45.379	51.470	51.526	45.783
80	21.031 0	0.127 13	20.620 9	38.662	52.107	59.143	59.508	52.873
90	23.852 9	0.143 95	23.389 8	43.855	59.144	67.426	67.999	59.967
100	26.677 4	0.160 78	26.161 3	49.053	66.187	75.417	75.199	67.070

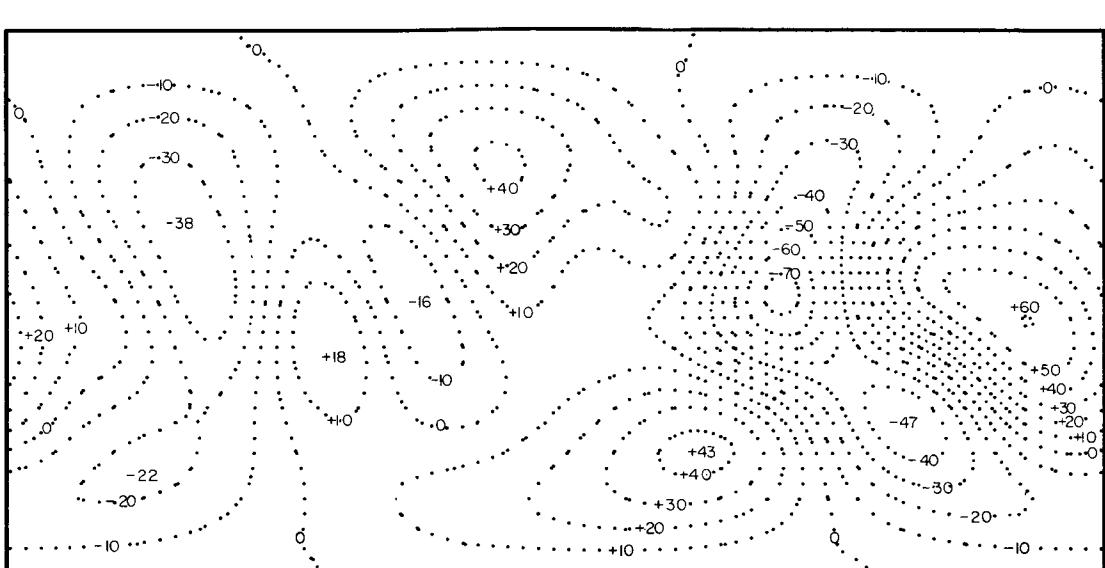
$s \text{ km}$ $\times 10^3$	$\varphi = 70^\circ$	$\varphi = -80^\circ$	$\varphi = -90^\circ$
0	0.000 0	0.000 0	0
1	0.072 6	0.037 8	
4	0.823 0	0.432 6	
10	3.249 5	1.715 1	
20	8.056 5	4.258 7	
30	13.134 6	6.946 6	
40	18.308 3	9.685 3	
50	23.226 7	12.447 9	
60	28.769 6	15.223 4	
70	34.027	18.006 9	
80	39.295	20.795 5	
90	44.569	23.587 7	
100	49.848	26.382 4	

**III. APPENDICES**

**APPENDIX A. Contour Maps**

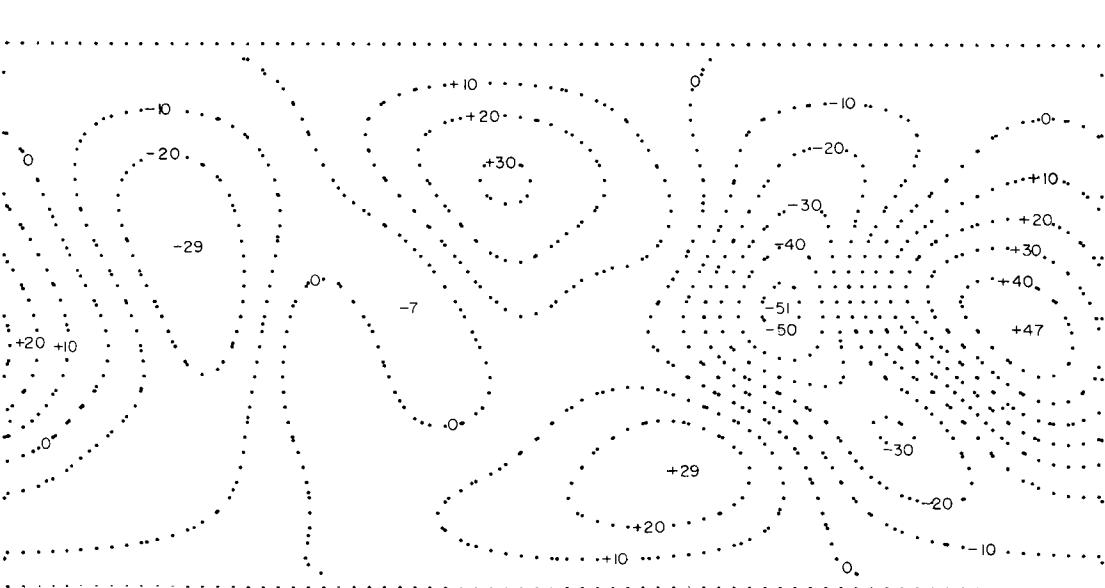
Contour Map 1.--Shape of the equipotential surfaces.

$\Delta r$  at sea level ( $\omega \neq 0$ ,  $\omega = 0$ )



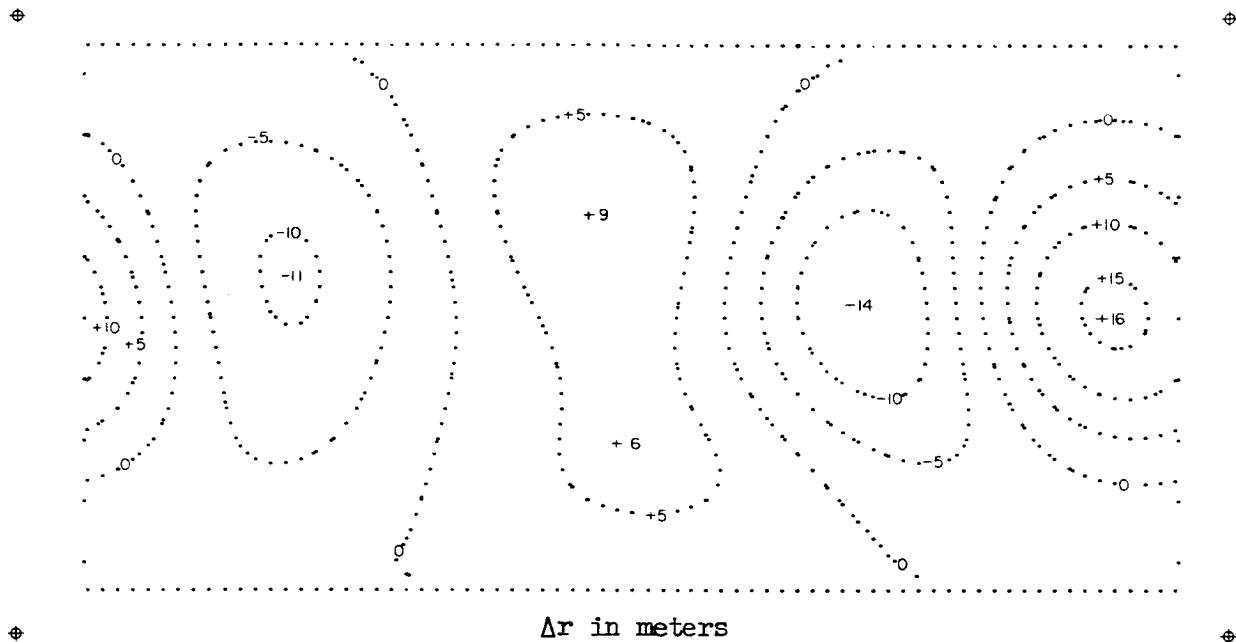
$\Delta r$  in meters

$\Delta r$  at 1,000 km elevation ( $\omega = 0$ )

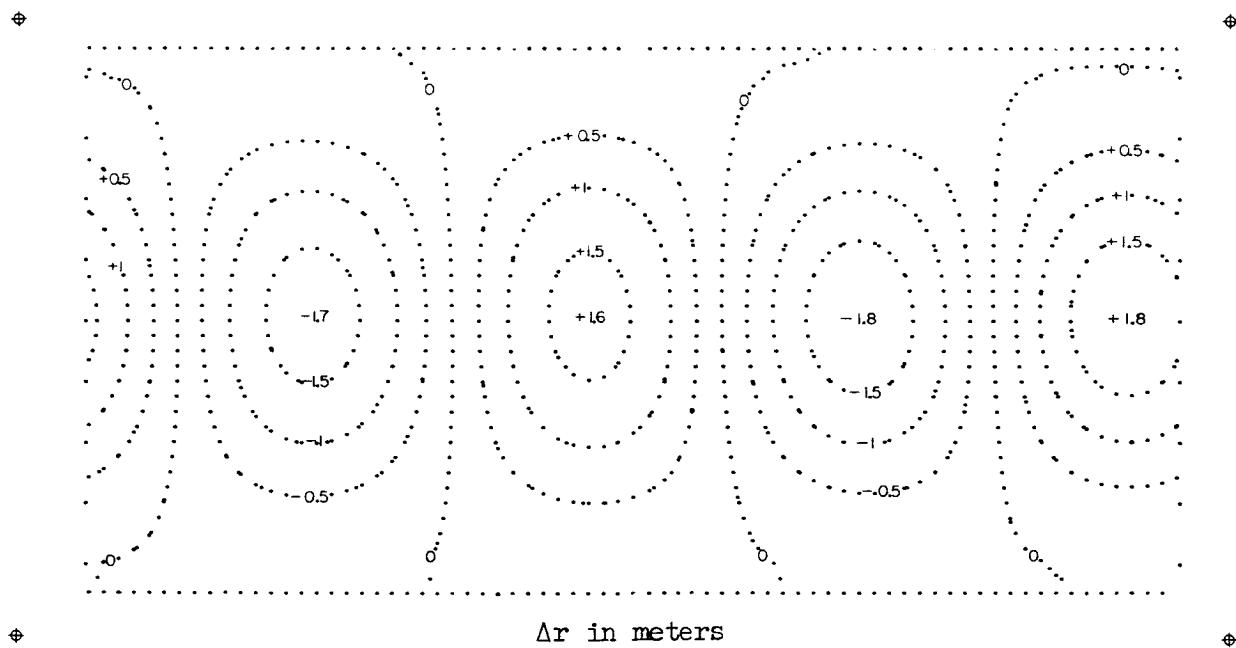


$\Delta r$  in meters

$\Delta r$  at 10,000 km elevation ( $\omega = 0$ )

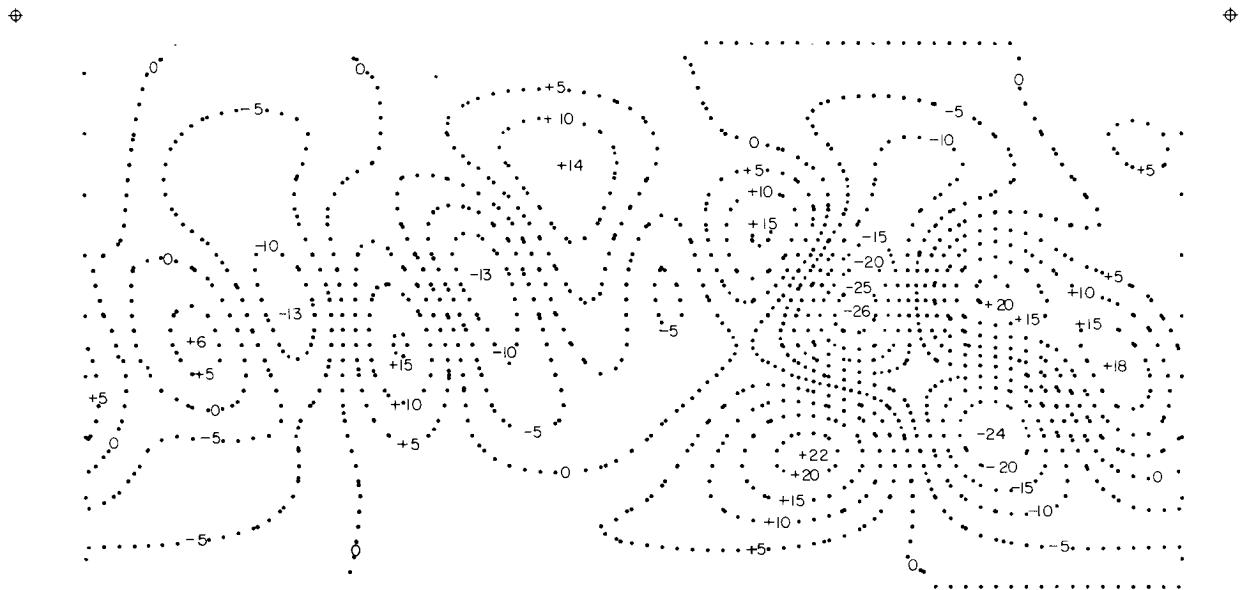


$\Delta r$  at 100,000 km elevation ( $\omega = 0$ )



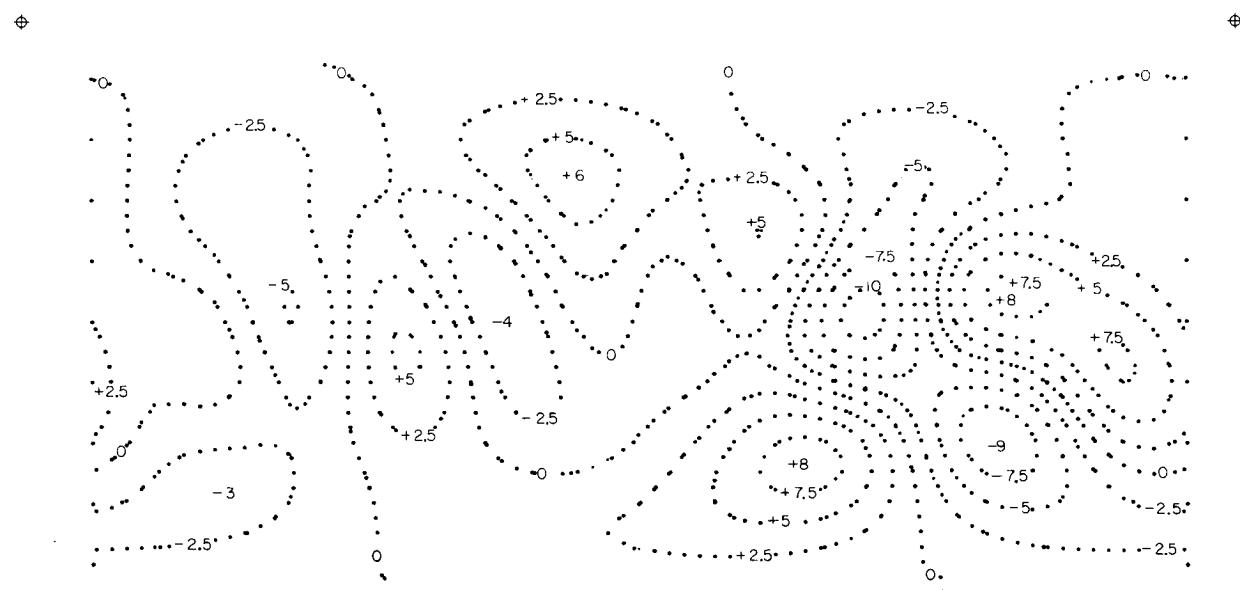
Contour Map 2.--Gravitation (gravity) along an equipotential surface.

$\Delta g$  at sea level ( $\omega \neq 0$ ,  $\omega = 0$ )



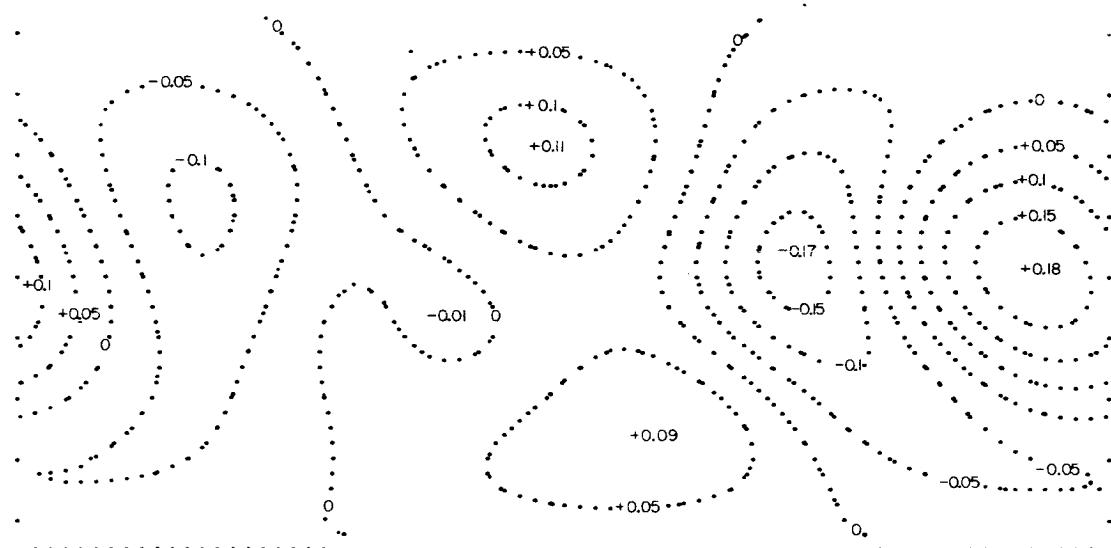
$\Delta g$  in  $10^{-3}$  cm/sec $^2$

$\Delta g$  at 1,000 km elevation ( $\omega = 0$ )



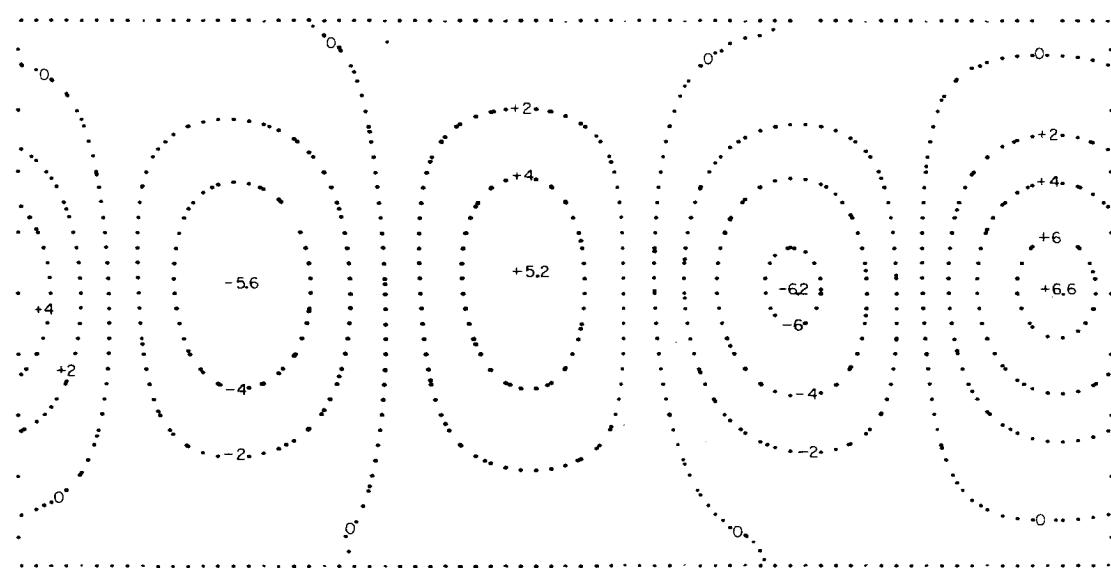
$\Delta g$  in  $10^{-3}$  cm/sec $^2$

$\Delta g$  at 10,000 km elevation ( $\omega = 0$ )



$\Delta g$  in  $10^{-3} \text{ cm/sec}^2$

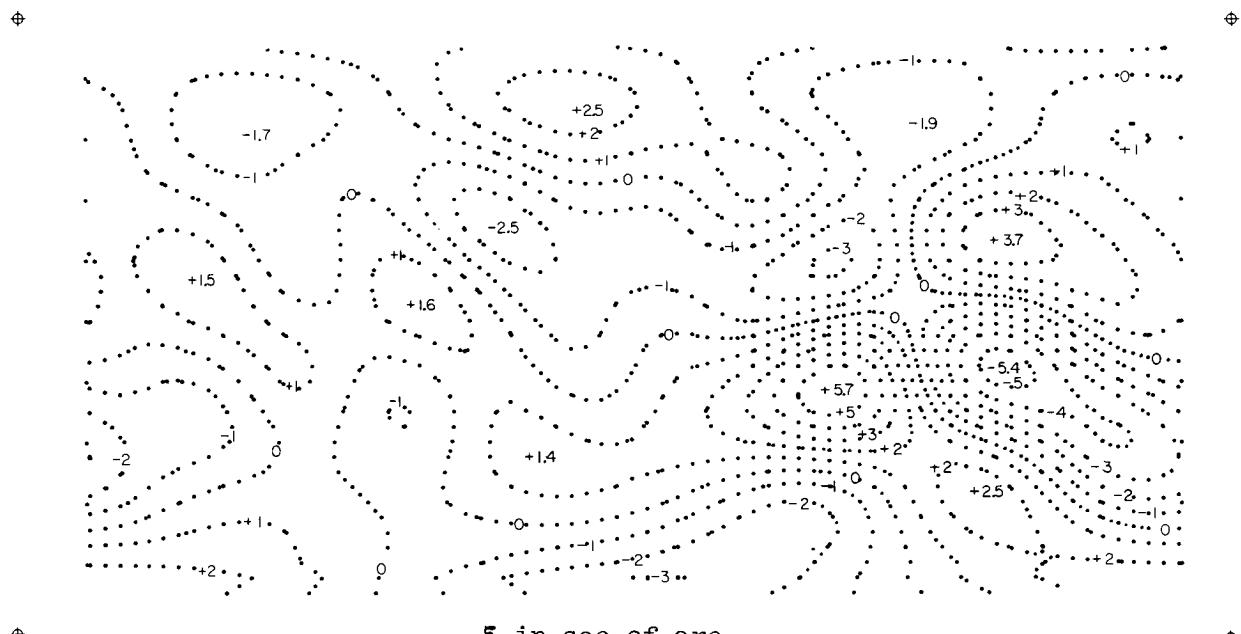
$\Delta g$  at 100,000 km elevation ( $\omega = 0$ )



$\Delta g$  in  $10^{-8} \text{ cm/sec}^2$

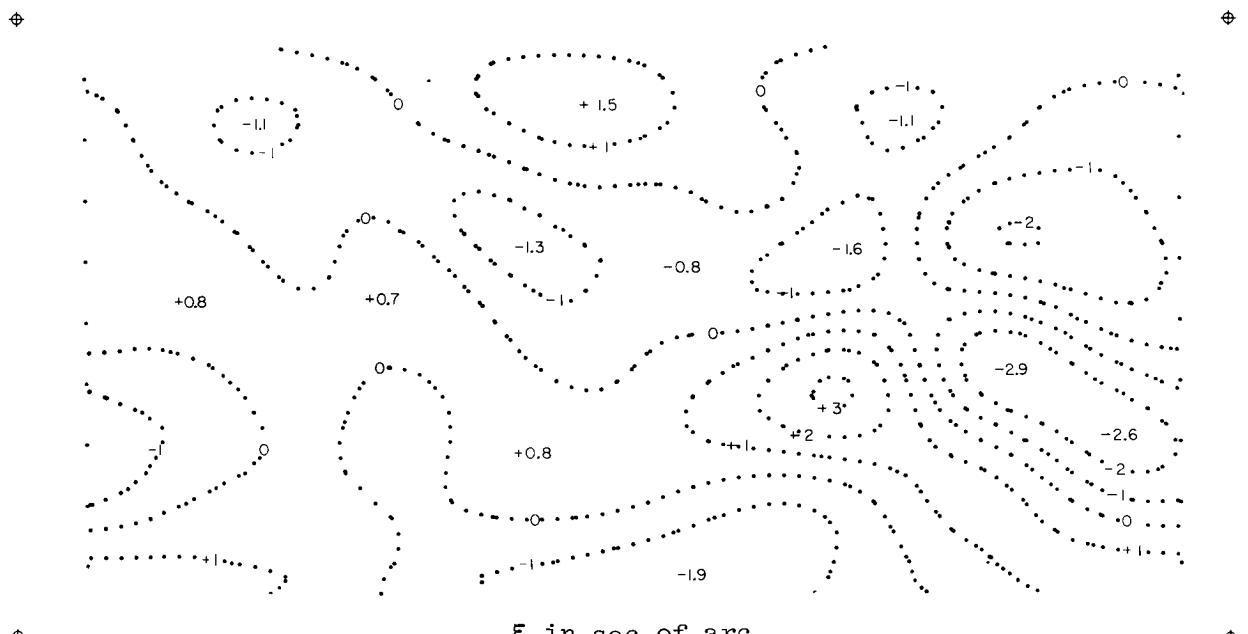
Contour Map 3---Oscillation of the surface normals; latitude curves  
 (equipotential surfaces).

$\xi$  at sea level ( $\omega \neq 0$ ,  $\omega = 0$ )



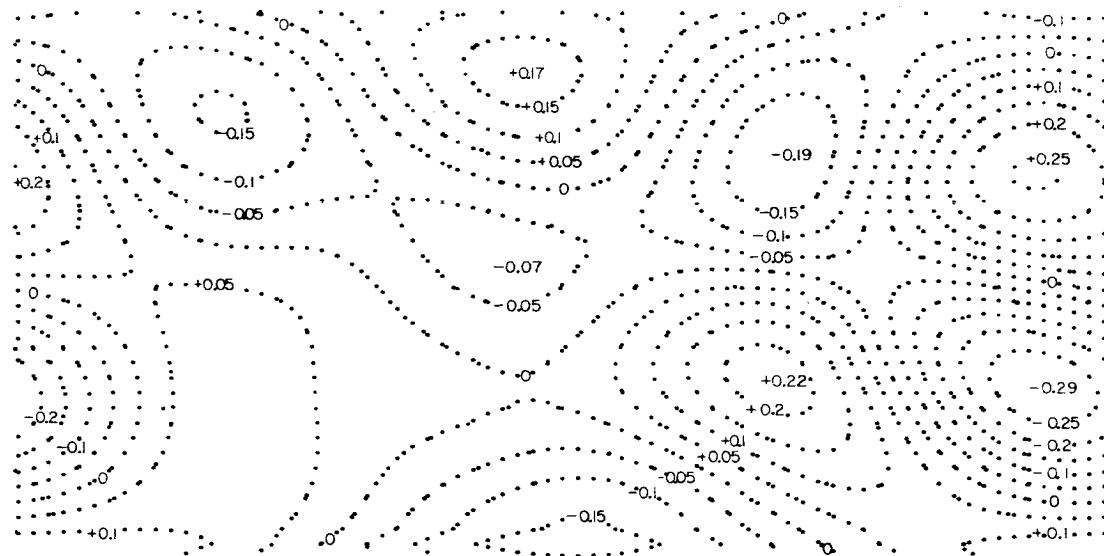
$\xi$  in sec of arc

$\xi$  at 1,000 km elevation ( $\omega = 0$ )



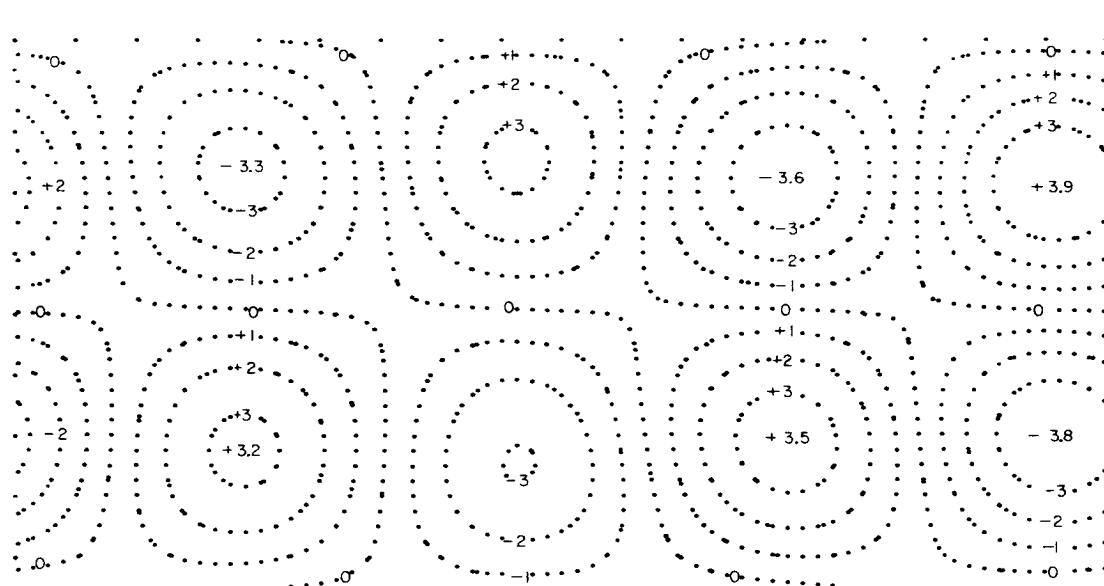
$\xi$  in sec of arc

$\xi$  at 10,000 km elevation ( $\omega = 0$ )



$\xi$  in sec of arc

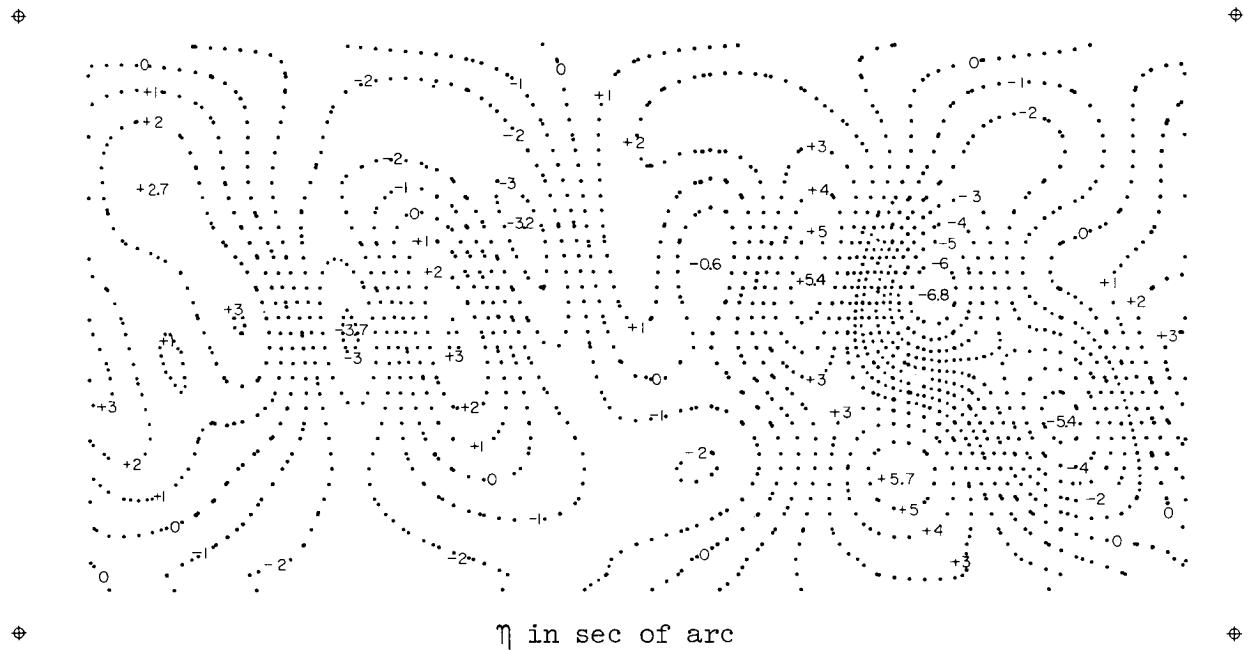
$\xi$  at 100,000 km elevation ( $\omega = 0$ )



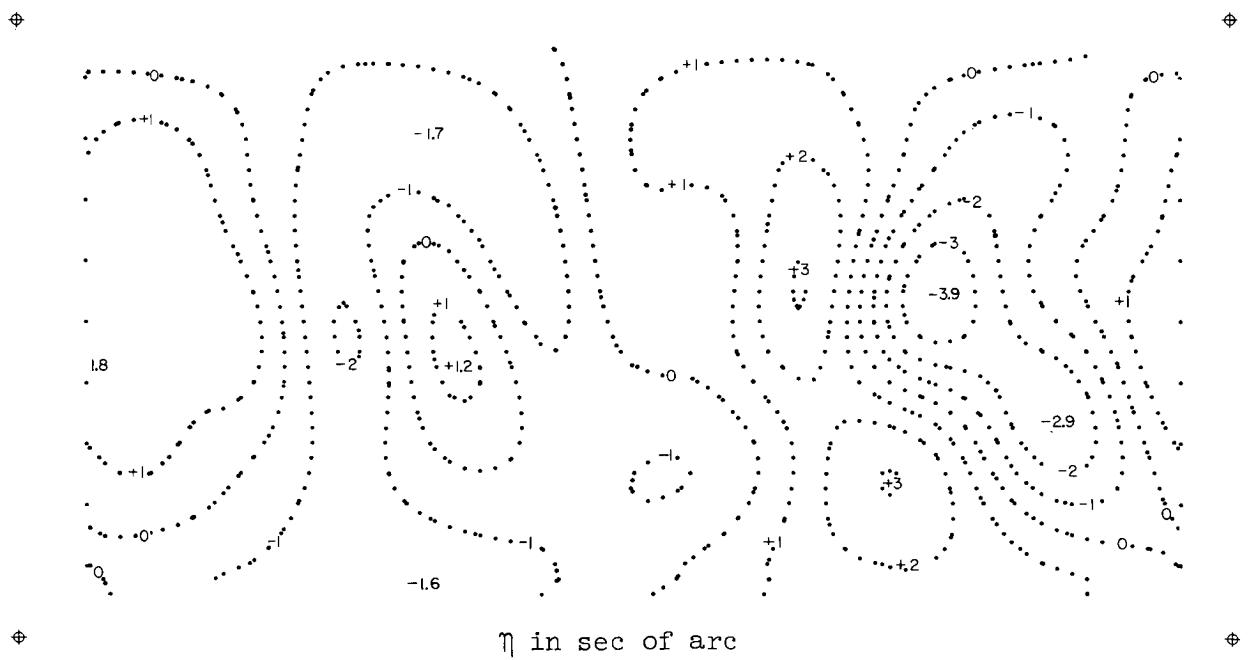
$\xi$  in  $10^{-3}$  sec of arc

Contour Map 4.--Oscillation of the surface normals; longitude curves (equipotential surfaces).

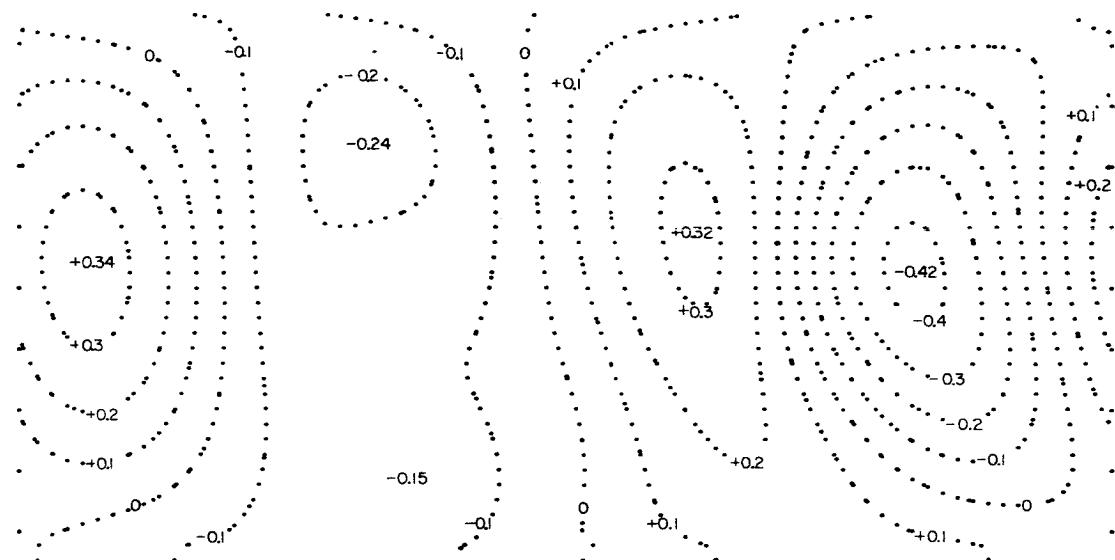
$\eta$  at sea level ( $\omega \neq 0$ ,  $\omega = 0$ )



$\eta$  at 1,000 km elevation ( $\omega = 0$ )

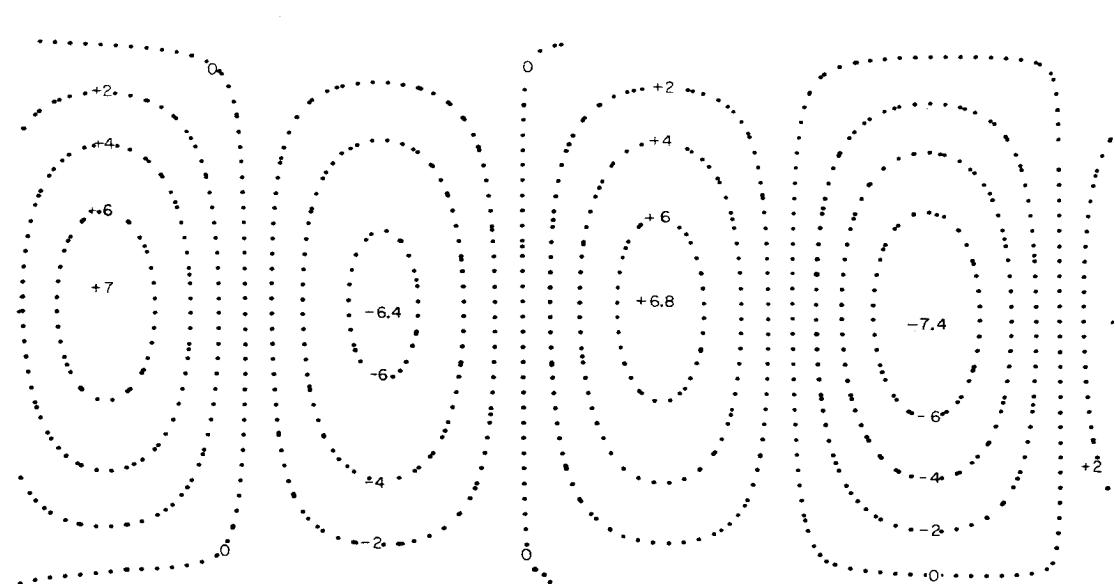


$\eta$  at 10,000 km elevation ( $\omega = 0$ )



$\eta$  in sec of arc

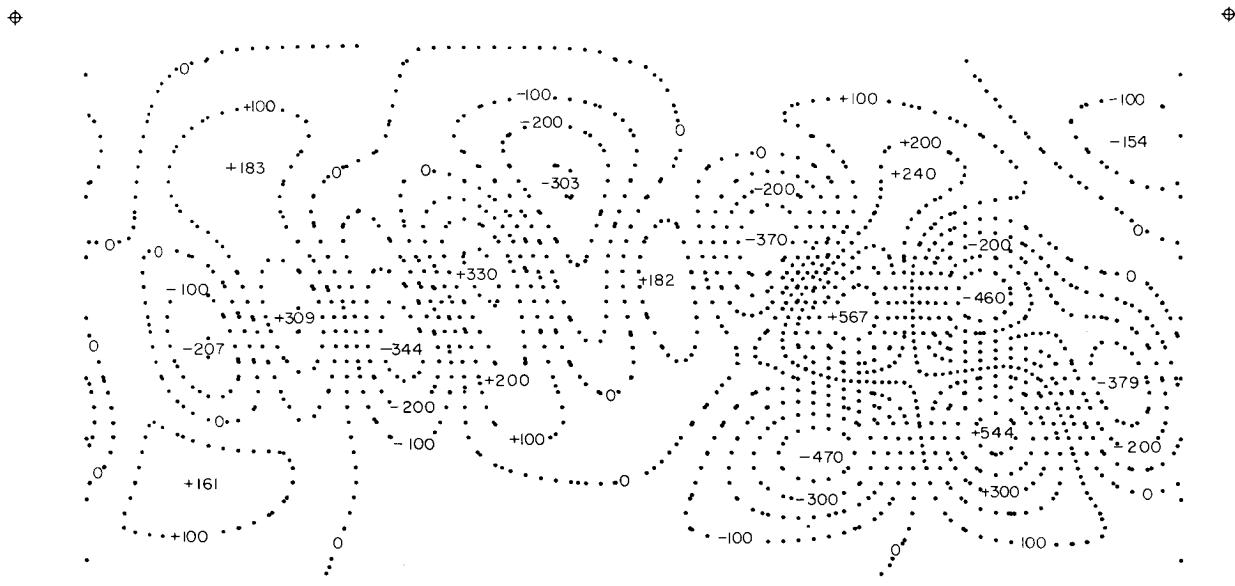
$\eta$  at 100,000 km elevation ( $\omega = 0$ )



$\eta$  in  $10^{-3}$  sec of arc

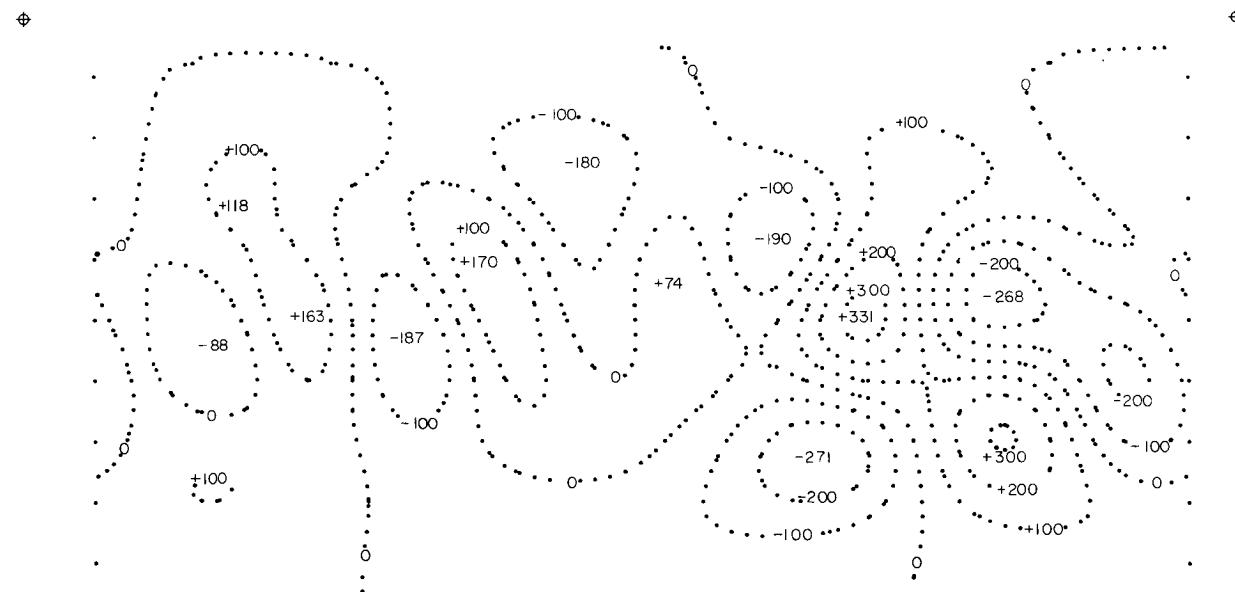
Contour Map 5.--Gaussian and mean curvature.

$\Delta \frac{1}{\sqrt{K}}$ ,  $\Delta \frac{1}{H}$  at sea level ( $\omega \neq 0$ ,  $\omega = 0$ )



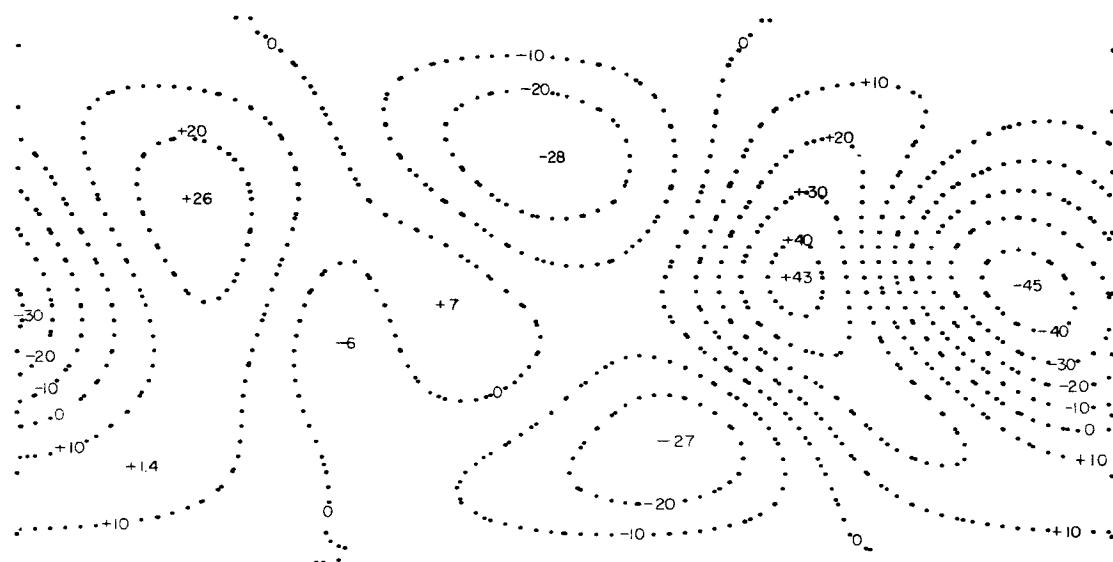
$$\Delta \frac{1}{\sqrt{K}}, \Delta \frac{1}{H} \text{ in meters}$$

$\Delta \frac{1}{\sqrt{K}}$ ,  $\Delta \frac{1}{H}$  at 1,000 km elevation ( $\omega = 0$ )

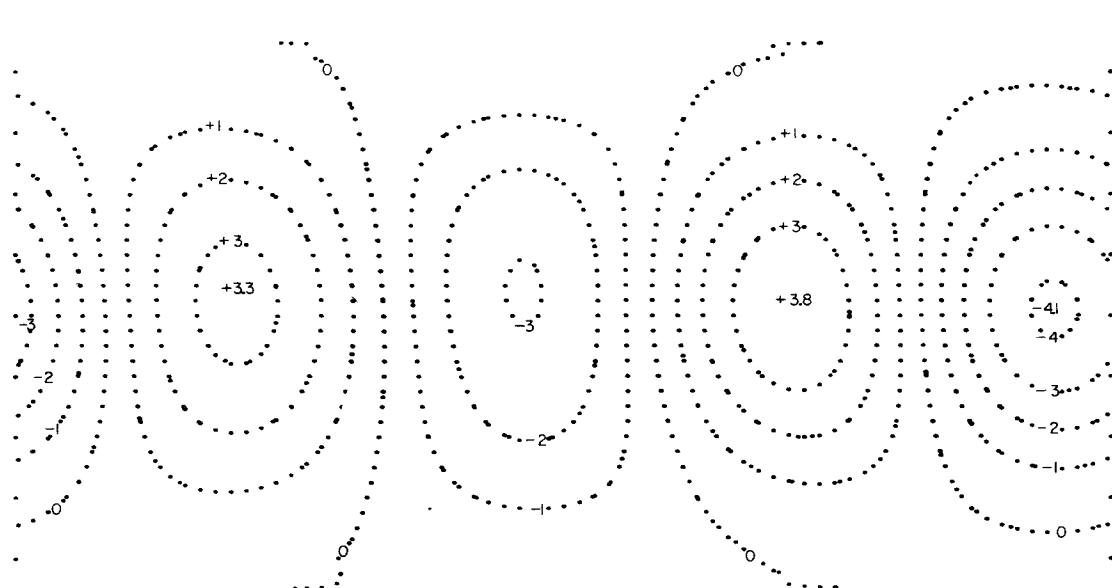


$$\Delta \frac{1}{\sqrt{K}}, \Delta \frac{1}{H} \text{ in meters}$$

$\Delta \frac{1}{\sqrt{K}}$ ,  $\Delta \frac{1}{H}$  at 10,000 km elevation ( $\omega = 0$ )



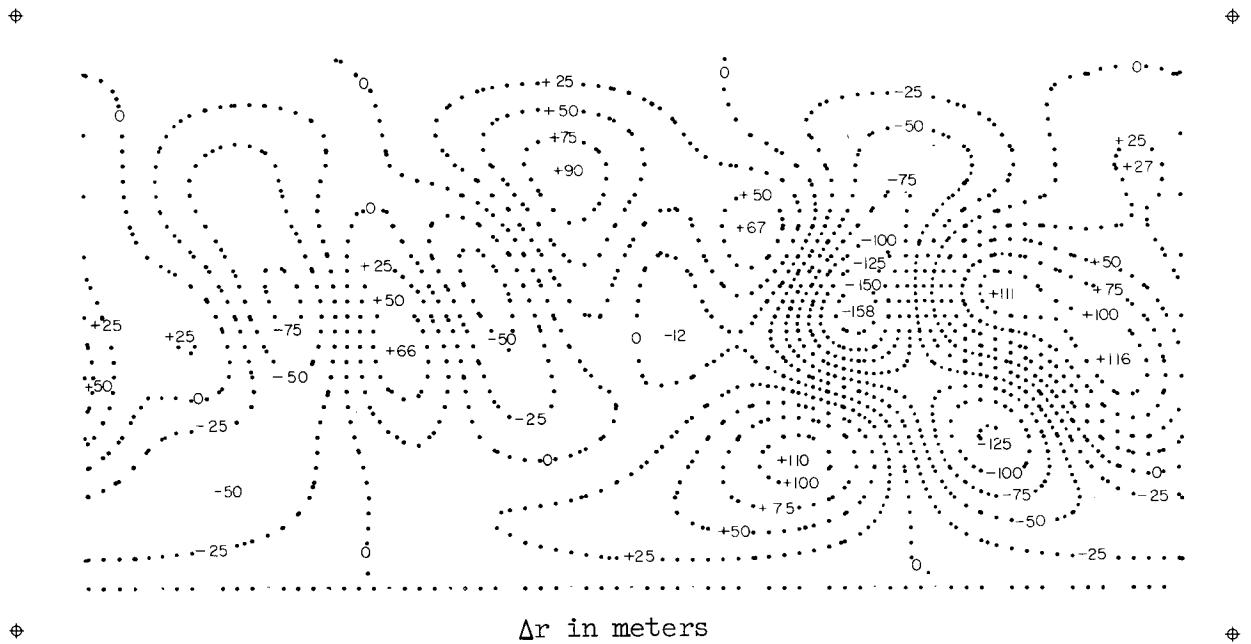
$\Delta \frac{1}{\sqrt{K}}$ ,  $\Delta \frac{1}{H}$  in meters  
 $\Delta \frac{1}{\sqrt{K}}$ ,  $\Delta \frac{1}{H}$  at 100,000 km elevation ( $\omega = 0$ )



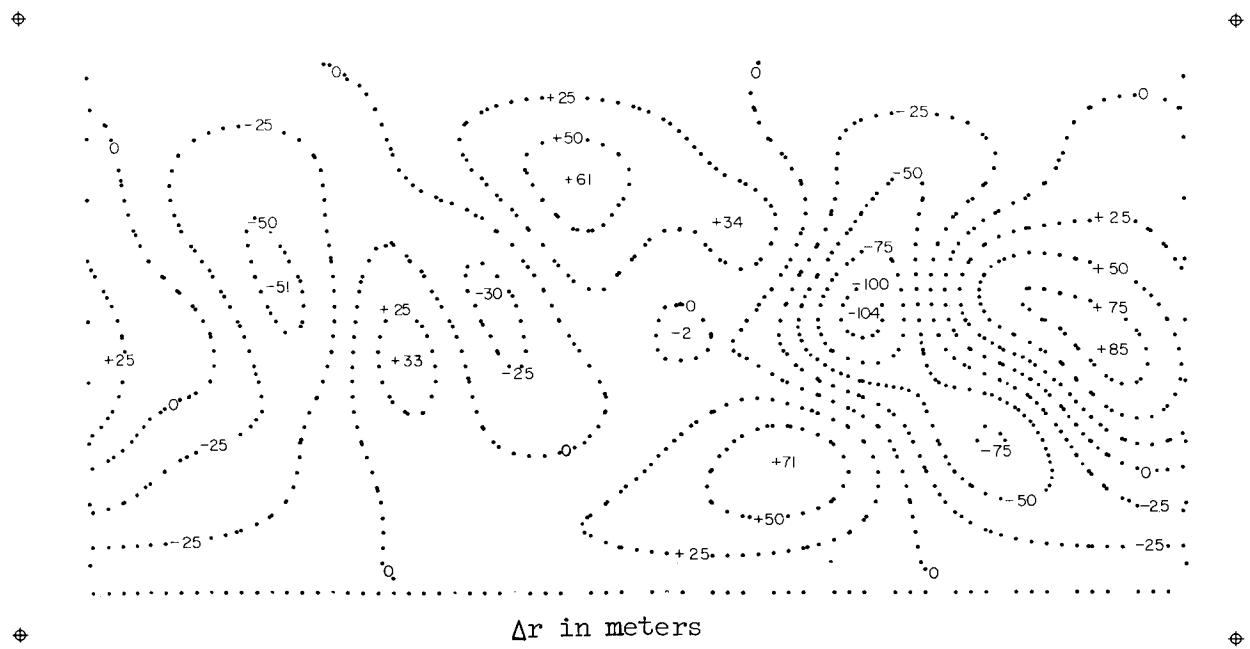
$\Delta \frac{1}{\sqrt{K}}$ ,  $\Delta \frac{1}{H}$  in meters

Contour Map 6.--Shape of the equigravitational (equigravity) sur

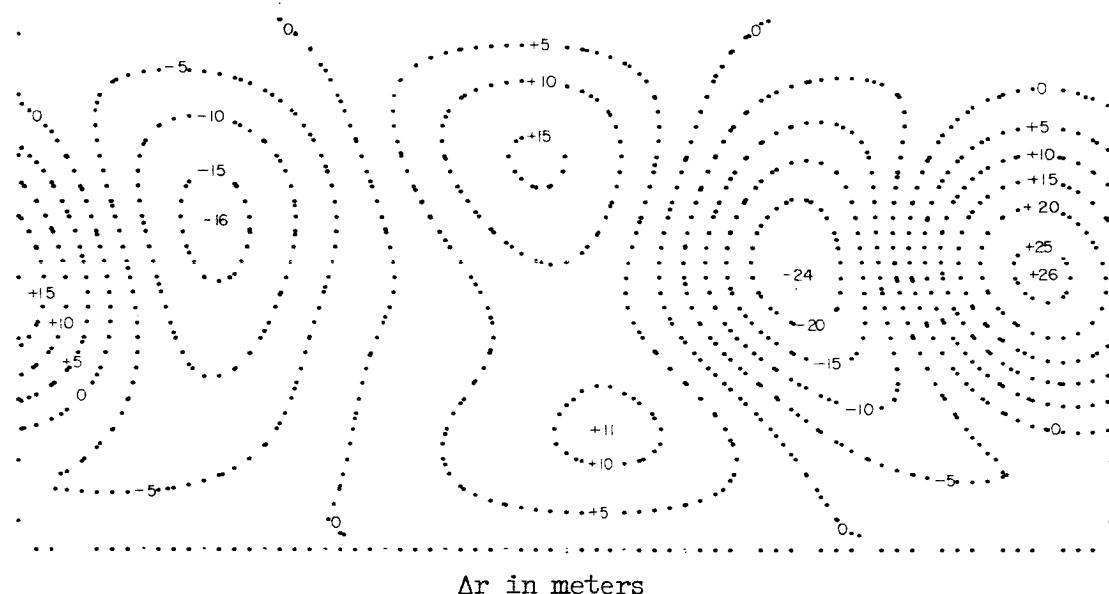
$\Delta r$  at sea level ( $\omega \neq 0$ ,  $\omega = 0$ )



$\Delta r$  at 1,000 km elevation ( $\omega = 0$ )

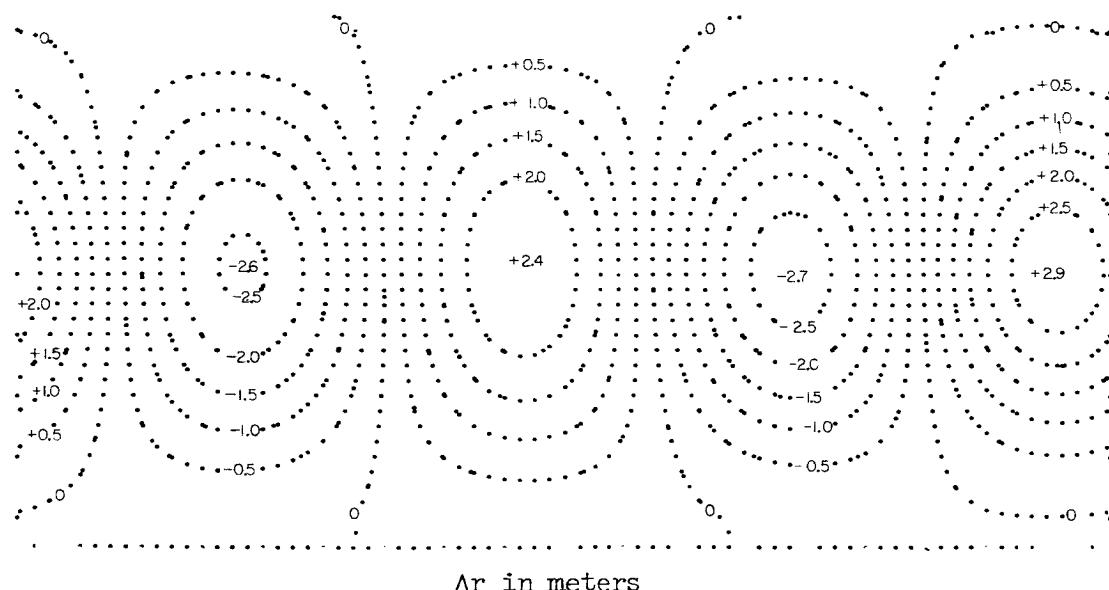


$\Delta r$  at 10,000 km elevation ( $\omega = 0$ )



$\Delta r$  in meters

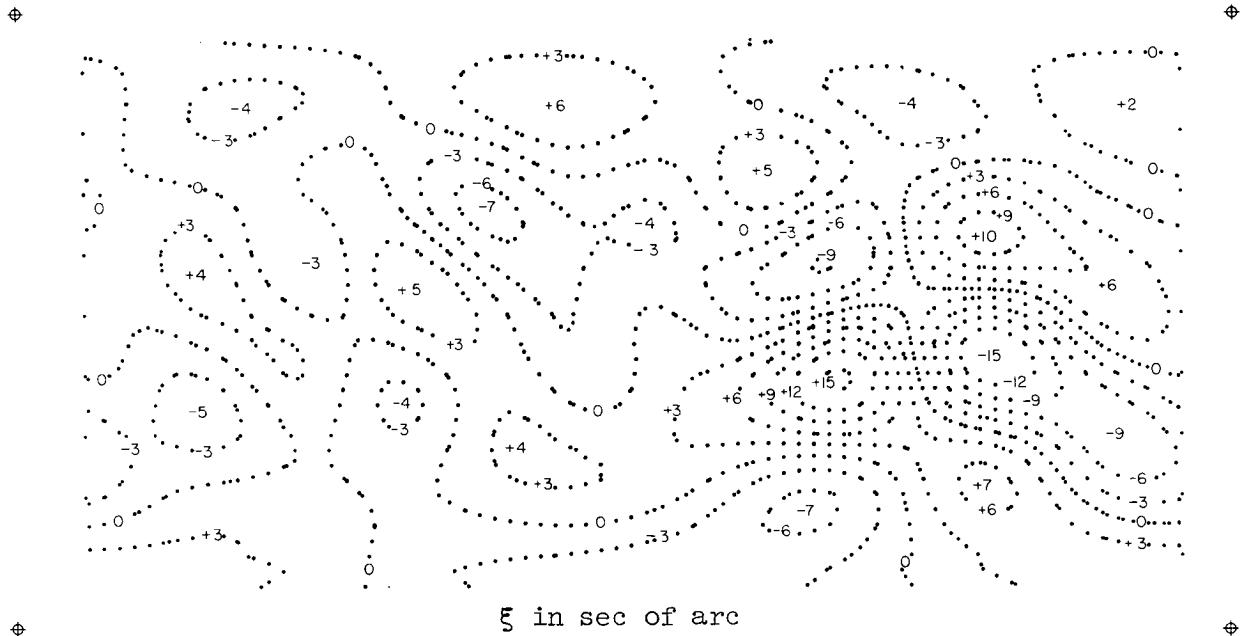
$\Delta r$  at 100,000 km elevation ( $\omega = 0$ )



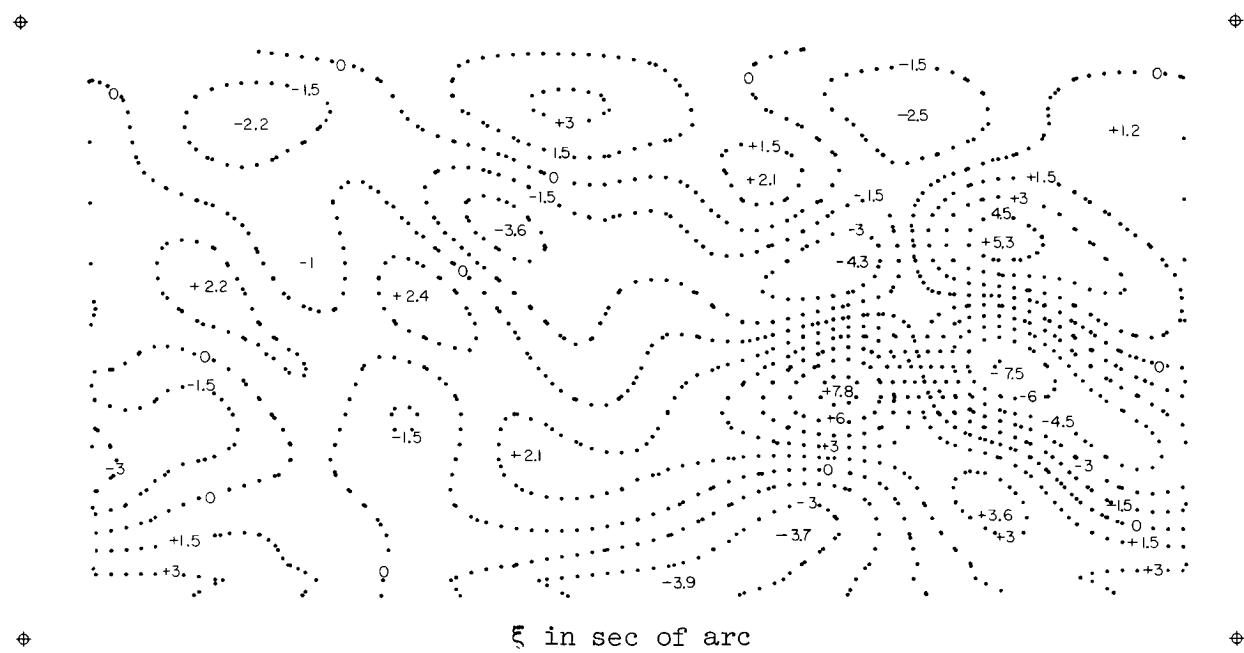
$\Delta r$  in meters

Contour Map 7.--Oscillation of the surface normals (equigravitational, equigravity surfaces).

$\xi$  at sea level ( $\omega \neq 0$ ,  $\omega = 0$ )



$\xi$  at 1,000 km elevation ( $\omega = 0$ )

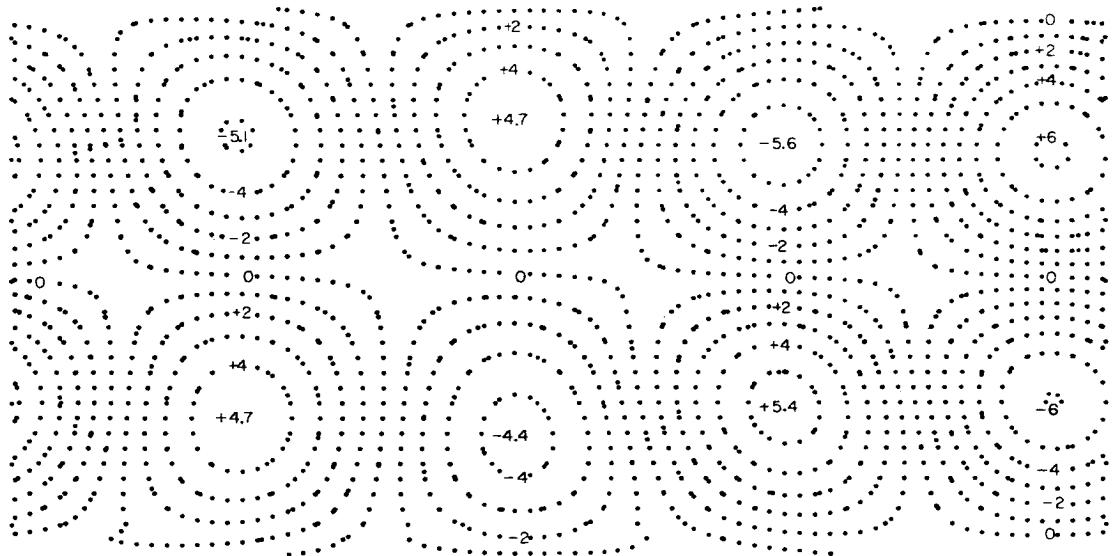


$\xi$  at 10,000 km elevation ( $\omega = 0$ )



$\xi$  in sec of arc

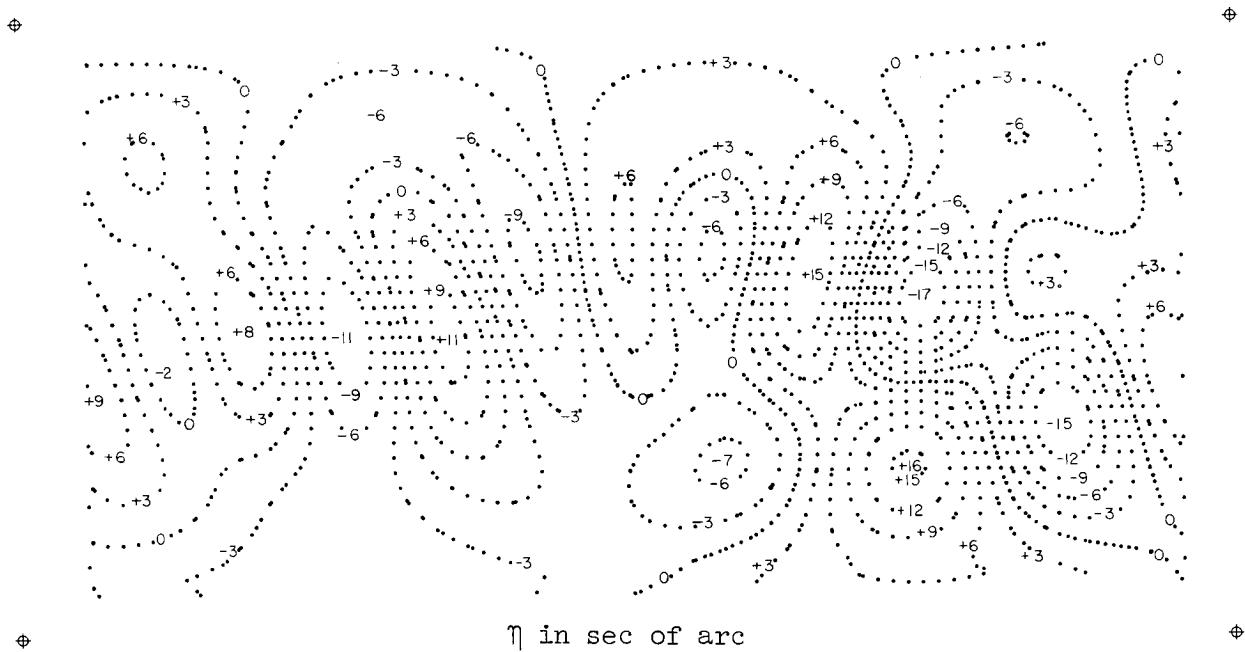
$\xi$  at 100,000 km elevation ( $\omega = 0$ )



$\xi$  in  $10^{-3}$  sec of arc

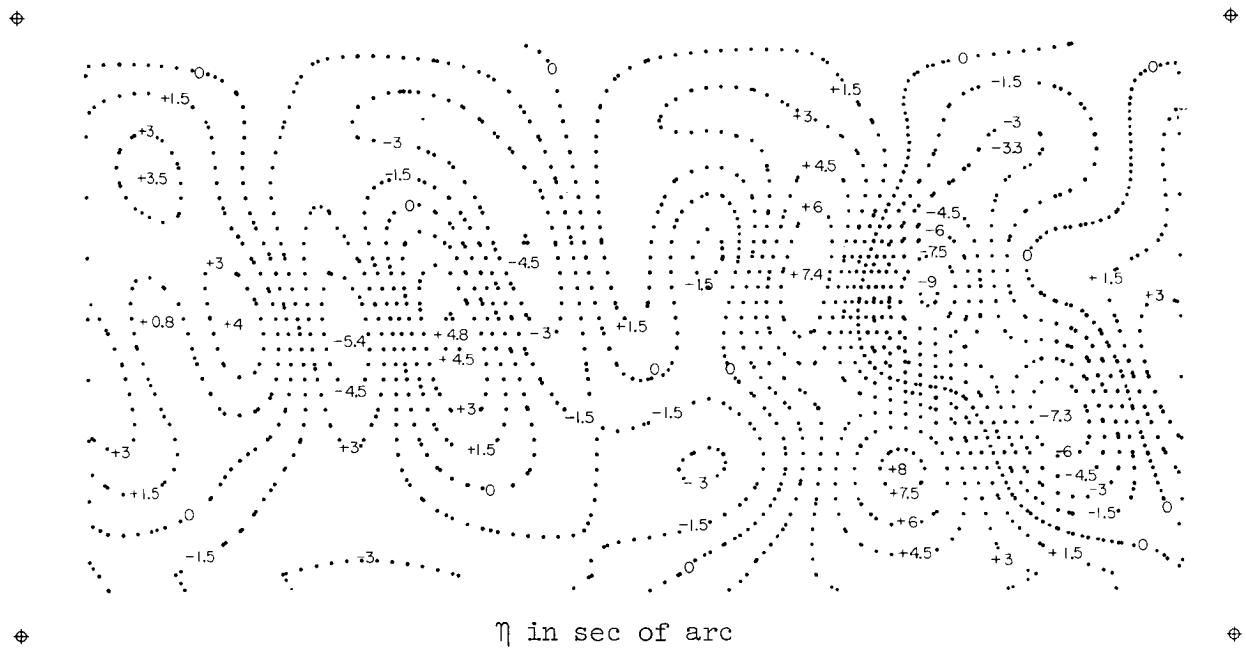
Contour Map 8.--Oscillation of the surface normals (equigravitational, equigravity surfaces).

$\eta$  at sea level ( $\omega \neq 0$ ,  $\omega = 0$ )



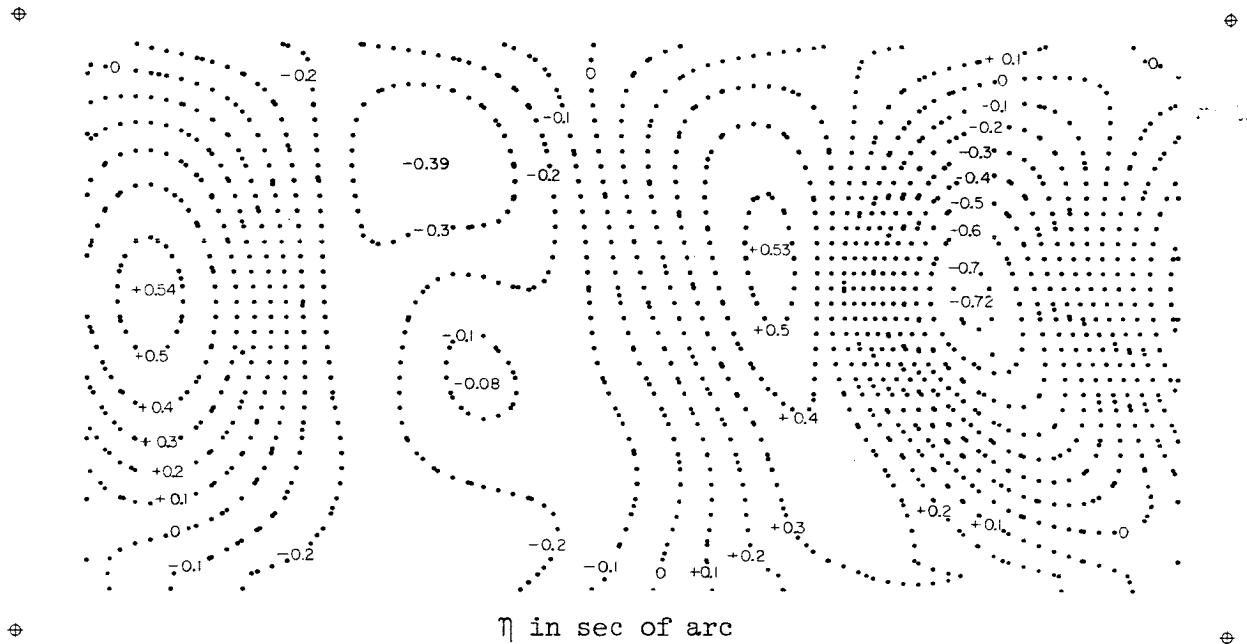
$\eta$  in sec of arc

$\eta$  at 1,000 km elevation ( $\omega = 0$ )

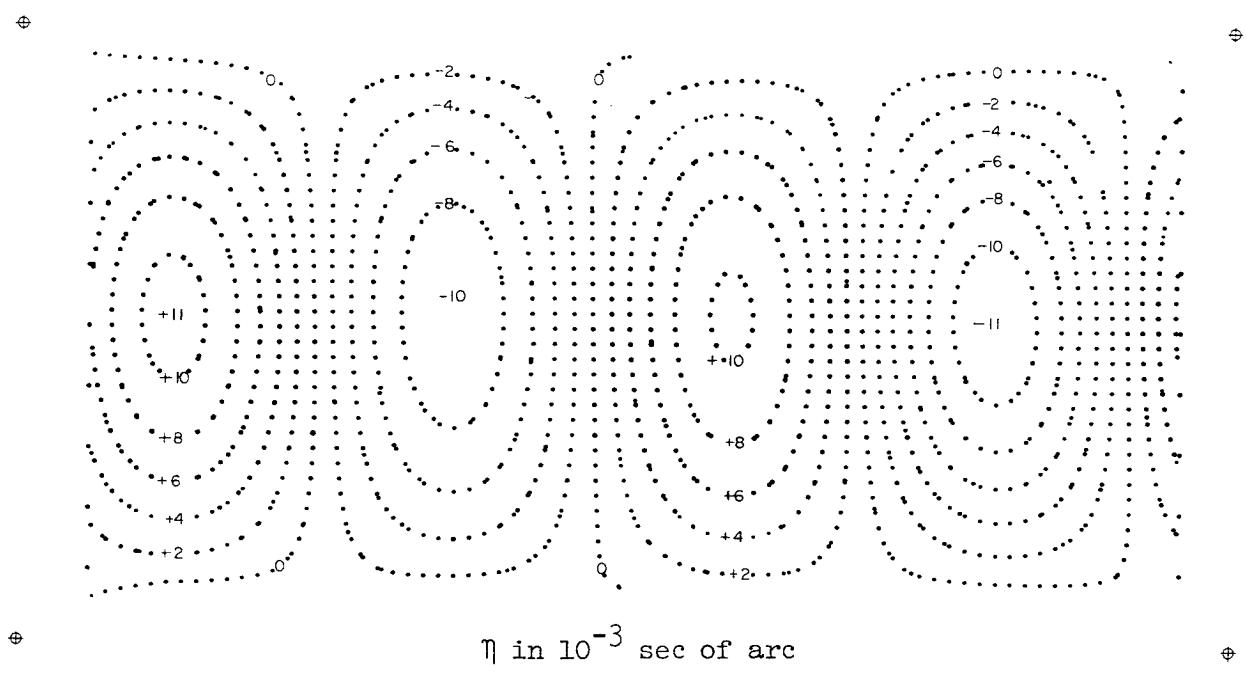


$\eta$  in sec of arc

$\eta$  at 10,000 km elevation ( $\omega = 0$ )



$\eta$  at 100,000 km elevation ( $\omega = 0$ )



**APPENDIX B. Number Maps**

Number Map 1.--Intersection of the equigravitational (equigravity) surfaces  
with the equipotential surfaces.

$\Delta\mu$  at sea level ( $\omega \neq 0$ )

-18	-17	-16	-15	-14	-13	-12	-11	-10	-9	-8	-7	-6	-5	-4	-3	-2	-1	C+	1+	2+	3+	4+	5+	6+	7+	8+	9+10+11+12+13+14+15+16+17+18													
90	5	5	5	5	5	5	5	5	5	5	5	5	5	5	5	5	5	5	5	5	5	5	5	5	5	5														
85	-5	-3	-0	2	5	6	6	4	1	-2	-5	-8	-11	-13	-13	-11	-8	-4	-C	4	7	10	12	13	13	11	9	6	3	-1	-4	-6	-7	-6	-5					
80	-9	-5	1	7	11	14	15	13	10	5	-0	-7	-13	-18	-21	-23	-22	-15	-14	-E	2	4	10	14	17	19	18	14	10	4	-2	-8	-11	-13	-12	-9				
75	-11	-5	3	10	16	20	20	17	13	7	-0	-7	-15	-22	-27	-29	-26	-24	-18	-11	-4	3	8	13	17	21	23	23	20	15	8	0	-7	-13	-16	-15	-11			
70	-11	-4	5	13	19	21	20	17	11	6	-0	-6	-13	-21	-28	-32	-31	-27	-26	-12	-6	1	2	6	11	17	22	25	24	20	13	5	-4	-11	-15	-15	-11			
65	-9	-2	6	14	19	20	17	12	7	2	-0	-3	-8	-15	-23	-29	-26	-18	-11	-7	-7	-5	1	9	18	24	26	23	18	10	2	-5	-10	-12	-9					
60	-5	-0	7	13	17	17	12	5	0	-2	-0	2	2	-4	-14	-23	-25	-21	-13	-7	-7	-12	-18	-15	-12	-8	12	20	24	22	19	14	8	2	-3	-6	-5			
55	-1	1	6	11	13	12	6	-1	-6	-6	-0	9	14	10	-1	-6	-18	-29	-32	-24	-9	6	15	17	16	14	13	12	9	4	0	-1								
50	3	3	4	6	8	7	2	-5	-11	-10	-0	4	14	25	15	14	-1	-9	-6	2	5	-3	-21	-37	-41	-31	-13	2	8	6	3	3	7	12	13	11	7	3		
45	7	3	1	1	2	2	-6	-12	-12	-2	17	33	37	27	11	1	2	10	12	0	-22	-41	-44	-35	-10	2	2	-7	-14	-13	-4	7	14	15	11	7				
40	9	4	-2	-5	-6	-1	1	-2	-10	-13	-5	15	36	45	37	31	9	9	16	18	4	-19	-37	-38	-21	-1	7	-2	-20	-33	-32	-19	-2	10	15	13	9			
35	10	4	-5	-11	-10	-3	4	4	-4	-12	-9	9	32	46	43	28	15	14	20	21	8	-13	-28	-24	-5	14	15	-4	-32	-51	-50	-34	-13	3	11	13	10			
30	9	3	-7	-16	-16	-5	8	12	3	-11	-4	2	23	41	43	31	18	15	20	22	12	-4	-14	-5	16	31	25	-4	-39	-62	-63	-47	-25	-8	3	9	-9			
25	7	3	-9	-20	-21	-8	10	19	10	-9	-20	-13	1C	31	39	32	26	15	18	2C	14	5	3	17	37	47	33	-1	-40	-65	-68	-53	-34	-19	-7	2	7			
20	4	3	-9	-22	-25	-11	10	23	16	-6	-25	-24	-6	18	31	29	20	13	15	15	14	20	36	54	57	38	2	-35	-58	-61	-51	-39	-28	-18	-6	4				
15	-1	3	-8	-23	-28	-15	7	23	18	-4	-27	-33	-19	4	22	16	11	8	9	14	21	33	45	62	58	36	6	-23	-40	-43	-39	-37	-35	-28	-14	-1				
10	-5	2	-5	-20	-27	-18	1	18	17	-2	-27	-37	-29	-8	13	22	16	9	3	11	25	40	53	58	47	28	8	-7	-14	-16	-20	-28	-36	-35	-22	-5				
5	-10	1	-3	-16	-25	-20	-6	9	14	0	-23	-35	-32	-17	5	19	18	8	-0	-2	8	25	39	48	44	27	12	10	13	16	6	-15	-33	-38	-27	-10				
-10	29	0	-2	1	5	25	55	41	10	65	68	18	55	78	39	20	53	24	16	33	11	22	33	33	71	79	28	45	97	93	61	48	36	6	19	35	41	29		
-5	17	3	-3	2	12	21	20	11	3	2	6	20	30	22	-7	-10	-17	-11	3	8	2	-11	-17	-5	18	34	33	9	-32	-67	-75	-56	-20	13	31	30	17			
-10	17	4	-5	-7	-1	14	21	18	8	-2	-4	6	19	2C	8	-8	-16	-13	0	9	1	20	25	17	1	20	25	17	1	20	25	17	1	20	25	17	1			
-15	15	5	-8	-14	-10	5	19	21	11	-5	-13	-7	8	16	9	-6	-17	-14	-7	5	11	10	17	41	71	86	68	20	-41	-85	-98	-80	-47	-14	8	17	15	15		
-20	11	4	-9	-20	-20	-5	13	20	11	-7	-20	-17	-C	12	11	-2	-14	-14	-4	8	15	19	31	57	85	96	75	25	-35	-77	-90	-77	-51	-26	-7	6	11	11	11	
-25	4	2	-11	-25	-27	-14	6	16	10	-9	-25	-23	-6	1C	14	3	-5	-12	-4	8	18	26	4C	64	88	95	74	28	-24	-60	-73	-66	-50	-35	-21	-6	4			
-30	-5	-2	-11	-26	-32	-21	-2	11	6	-11	-25	-24	-8	10	16	9	-3	-8	-2	5	2C	28	60	78	82	63	28	-11	-37	-48	-47	-44	-40	-32	-18	-5				
-35	-14	-6	-12	-25	-32	-25	-8	5	3	-11	-23	-22	-6	12	20	15	4	-1	2	11	20	2C	27	36	47	58	59	47	25	2	-13	-20	-26	-34	-41	-30	-14			
-40	-4C	-22	-11	-12	-21	-28	-24	-11	0	0	-9	-18	-17	-3	13	22	20	12	6	7	14	19	22	25	26	32	32	32	27	27	19	13	10	6	-4	-21	-39	-46	-38	-22
-45	-28	-15	-11	-16	-21	-18	-10	-1	-1	-7	-13	-11	C	15	24	23	18	13	13	15	17	11	7	5	5	8	13	21	27	26	13	-9	-33	-46	-42	-28				
-50	-30	-16	-9	-9	-11	-10	-5	-0	-0	-4	-8	-6	3	15	22	24	21	18	16	16	13	6	-4	-14	-19	-18	-9	6	24	37	29	26	1	-25	-41	-42	-30			
-55	-55	-28	-15	-6	-2	-2	-1	3	1	-2	-5	-3	4	12	19	22	21	20	17	14	8	-3	-17	-29	-36	-34	-21	0	24	41	45	33	10	-16	-33	-37	-28			
-60	-22	-11	-2	4	5	7	7	6	4	0	-3	-2	2	8	14	17	18	17	15	1C	11	-2	-6	-11	-17	-22	-25	-24	-20	-13	-4	6	-6	-22	-27	-22				
-65	-65	-12	-5	3	9	12	13	12	9	5	1	-2	-3	1	3	7	10	12	12	9	4	-6	-18	-31	-42	-46	-41	-26	-5	16	33	39	34	20	3	-10	-15	-12		
-70	-70	-1	3	9	14	16	16	14	11	6	1	-3	-5	-5	-2	1	2	3	3	1	4	-14	-32	-22	-30	-43	-25	-22	-5	12	26	32	30	22	11	2	-2	-1		
-75	10	12	15	17	18	17	15	11	6	1	-4	-7	-8	-9	-8	-7	-6	-7	-9	-13	-18	-25	-30	-33	-32	-27	-17	-4	9	19	25	26	23	18	13	10	10			
-80	20	20	20	20	19	19	17	13	9	5	-0	-5	-8	-11	-13	-14	-15	-16	-18	-2C	-23	-25	-26	-23	-18	-11	-2	7	14	20	23	23	21	20	20	20				
-85	26	25	24	22	19	16	13	8	4	-1	-5	-9	-13	-16	-18	-20	-22	-23	-24	-25	-25	-23	-2C	-17	-12	-6	0	6	12	17	20	23	25	26	26	26				
-90	27	27	27	27	27	27	27	27	27	27	27	27	27	27	27	27	27	27	27	27	27	27	27	27	27	27	27	27	27	27	27	27	27	27	27					

$\Delta\mu$  in  $10^{-1}$  sec of arc

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$\Delta\mu$  at 1,000 km elevation ( $\omega = 0$ )

$\Delta\mu$  in  $10^{-1}$  sec of arc

$\Delta\mu$  at 10,000 km elevation ( $\omega = 0$ )

$\Delta\mu$  in  $10^{-2}$  sec of arc

$\Delta\mu$  at 100,000 km elevation ( $\omega = 0$ )

$\Delta\mu$  in  $10^{-4}$  sec of arc

Number Map 2.--The gradient (direction) field in a spherical coordinate system.

$\Delta\delta$  at sea level ( $\omega \neq 0$ ,  $\omega = 0$ )

$\Delta\delta$  in  $10^{-1}$  sec of arc

$\Delta\delta$  at 1,000 km elevation ( $\omega = 0$ )

$\Delta\delta$  in  $10^{-1}$  sec of arc

$\Delta S$  at 10,000 km elevation ( $\omega = 0$ )

$\Delta\delta$  in  $10^{-2}$  sec of arc

$\Delta\delta$  at 100,000 km elevation ( $\omega = 0$ )

$\Delta\delta$  in  $10^{-4}$  sec of arc

Number Map 3.--Radius of curvature of the orthogonal trajectories  
 (length unit: km). (See page 55 for explanation of  
 first column.)

$\rho$  at sea level ( $w \neq 0$ )

-18	-17	-16	-15	-14	-13	-12	-11	-10	-9	-8	-7	-6	-5	-4	-3	-2	-1	0	
9	26192	26952	27782	28662	29582	30522	31452	32342	33112	33732	34112	34222	34312	33512	32742	31772	32682	29562	28422
8	35340	35290	35180	35070	34940	34930	34940	34950	35010	35100	35210	35330	35450	35550	35610	35640	35630	35560	35470
7	14890	18850	18800	18760	18730	18710	18720	18740	18770	18800	18830	18870	18900	18950	18980	19010	19000	18980	18940
6	14010	13990	13970	13950	13940	13940	13960	13980	13990	14000	13990	13990	13990	14000	14030	14060	14070	14060	14030
5	12220	12220	12220	12210	12210	12210	12220	12240	12250	12250	12250	12250	12250	12250	12220	12220	12240	12240	12220
4	12170	12190	12200	12210	12200	12200	12190	12200	12220	12220	12220	12220	12220	12220	12170	12170	12170	12170	12160
3	13770	13780	13820	13840	13840	13810	13770	13760	13790	13830	13840	13800	13730	13680	13670	13700	13740	13750	13740
2	18520	18530	18590	18660	18680	18600	18490	18430	18460	18450	18670	18670	18570	18450	18430	18440	18470	18470	18470
1	35610	35470	35610	35900	36050	35870	35490	35180	35180	35550	36030	36240	36070	35680	35270	35090	35170	35340	35460
0	20642	26353	27262	37942	36362	11412	15112	49582	97441	93461	31862	11472	81941	15622	28892	11852	17912	34452	18372
-1	34680	34930	35112	35130	34980	34750	34620	34680	34870	35040	35080	34890	34660	34640	34870	35170	35340	35250	35000
-2	18420	18450	18520	18580	18580	18500	18410	18370	18420	18510	18580	18560	18480	18410	18420	18480	18550	18550	18500
-3	13900	13890	13920	13960	13980	13950	13890	13860	13870	13920	13960	13960	13910	13860	13860	13900	13910	13890	
-4	12230	12210	12210	12230	12250	12240	12210	12180	12200	12220	12220	12220	12190	12150	12130	12140	12160	12170	12170
-5	12290	12250	12240	12240	12240	12240	12230	12220	12220	12240	12240	12240	12210	12180	12160	12160	12180	12180	12180
-6	14190	14160	14130	14110	14100	14100	14110	14120	14130	14130	14120	14100	14080	14070	14070	14070	14070	14080	14080
-7	19030	19010	18970	18950	18930	18930	18960	18960	18990	19020	19050	19040	19050	19030	19010	19000	19000	19020	19020
-8	35110	35110	35100	35100	35120	35160	35220	35300	35390	35490	35570	35640	35700	35750	35770	35780	35800	35830	35830
-9	12772	12832	13002	13272	13662	14182	14862	15732	16842	18272	20122	22572	25912	30542	37772	49452	71102	11633	15583

0	1	2	3	4	5	6	7	8	9	10	11	12	13	14	15	16	17	18	
9	28422	27352	26392	25532	24802	24202	23722	23362	23122	22992	22982	23052	23252	23522	23892	24352	24882	25502	26192
8	35470	35350	35230	35120	35010	34930	34870	34830	34860	34930	35020	35130	35250	35350	35420	35460	35440	35380	
7	18940	18900	18860	18840	18820	18800	18770	18740	18710	18690	18700	18760	18800	18850	18890	18910	18910	18890	
6	14030	14010	14010	14030	14050	14050	14030	14000	13960	13930	13920	13930	13940	13950	13970	13990	14000	14010	14010
5	12220	12210	12230	12270	12310	12320	12300	12250	12220	12210	12210	12220	12220	12220	12190	12200	12210	12220	12220
4	12160	12150	12180	12240	12280	12280	12240	12200	12180	12200	12240	12270	12270	12240	12200	12170	12160	12160	12170
3	13740	13730	13760	13810	13830	13810	13750	13700	13720	13810	13910	13980	13980	13930	13870	13820	13780	13770	13770
2	18470	18470	18470	18470	18460	18460	18270	18250	18350	18350	18730	18850	18870	18810	18750	18690	18640	18570	18520
1	35460	35460	35300	35040	34760	34520	34440	34630	34990	35360	35640	35780	35820	35890	36060	36220	36190	35930	35610
0	13872	48432	26922	18092	18122	8921	81001	21432	13652	66565	66511	10352	12872	16992	70722	29752	17432	15012	20642
-1	35000	34840	34880	35020	35010	34670	34160	33870	34050	34760	35770	36600	36810	36400	35670	35010	34620	34540	34680
-2	18500	18430	18400	18380	18310	18190	18040	17990	18090	18350	18660	18890	18960	18960	18750	18610	18510	18440	18420
-3	13890	13860	13830	13800	13770	13710	13660	13650	13700	13810	13920	14000	14030	14030	14020	14000	13980	13940	13900
-4	12170	12150	12140	12130	12120	12120	12110	12120	12140	12150	12160	12170	12190	12230	12270	12270	12230	12230	12230
-5	12180	12180	12200	12220	12250	12260	12260	12240	12220	12160	12130	12130	12120	12120	12270	12310	12310	12290	12290
-6	14080	14090	14120	14160	14200	14240	14260	14250	14200	14130	14060	14010	13990	14020	14080	14140	14190	14210	14190
-7	19020	19050	19090	19150	19200	19240	19250	19220	19150	19050	18960	18880	18850	18860	18900	18960	19010	19040	19030
-8	35830	35880	35930	35980	36010	36000	35940	35860	35520	35550	35350	35200	35100	35040	35030	35050	35070	35090	35110
-9	15583	99582	63092	45272	35272	28992	24742	21702	19662	17752	16432	15432	13982	13532	13162	12932	12802	12772	12772

scaling:  $\times 10^3$

$\rho$  at sea level ( $w = 0$ )

-18	-17	-16	-15	-14	-13	-12	-11	-10	-9	-8	-7	-6	-5	-4	-3	-2	-1	0	
9	66732	74972	74572	70212	66672	63792	61452	59602	58172	57152	56522	56282	55642	57352	58142	59772	62062	65112	69132
8	11631	11731	11851	11971	12071	12141	12161	12131	12051	11941	11811	11681	11551	11451	11381	11361	11431	11531	
7	60790	61190	61700	62210	62580	62740	62670	62430	62110	61670	61640	61050	60650	60230	59870	59660	59670	59910	
6	44830	44970	45100	45400	45530	45370	45160	44990	44930	44990	45070	45050	44870	44570	44310	44220	44340	44580	
5	40050	40090	39880	39450	39060	38960	39200	39630	39550	39480	39680	40160	40080	40300	40300	40010	40130	40050	
4	40180	40230	39900	39330	38910	38890	39250	39730	39940	39730	39900	40160	40250	40350	40350	40050	40160	40020	
3	46110	46170	45840	45300	45060	45310	45910	46420	46260	45340	44240	43580	43570	44050	44050	44710	45250	45590	
2	61830	61990	61910	61840	62290	63150	64030	64640	63530	61080	59870	57870	56730	58350	58990	59570	60150	60820	
1	10781	10811	10941	11171	11511	11681	11611	11241	10711	10461	10471	10491	10411	10271	10151	10181	10391	10671	
0	18402	44852	27012	18152	18152	98281	81061	21422	13662	66091	68511	10342	12882	17032	71292	29742	17492	14992	
-1	11251	11381	11231	11251	11571	12141	12611	12341	11421	10681	98670	97190	10011	10601	11241	11661	11721	11551	
-2	61660	62370	62760	63000	63770	65360	67240	68050	66600	63340	59870	57670	57010	57630	58890	60320	61520	62210	
-3	44450	44790	45100	45370	45780	46400	46990	47110	46440	45320	44140	43380	43100	43070	43120	43260	43550	43950	
-4	40100	40240	40380	40460	40520	40620	40720	40710	40540	40310	40160	40040	39760	39310	38930	38820	39010	39370	
-5	40450	40440	40360	40180	39940	39													

$\rho$  at 1,000 km elevation ( $\omega = 0$ )

	-18	-17	-16	-15	-14	-13	-12	-11	-10	-9	-8	-7	-6	-5	-4	-3	-2	-1	0
9	17513	14853	12963	11573	10513	96802	90302	85152	81092	77922	75522	73822	72722	72132	72262	72932	74232	76232	79312
8	17851	17941	18041	18151	18231	18291	18291	18251	18161	18051	17911	17771	17631	17521	17441	17401	17411	17461	17551
7	94190	94590	95100	95610	96000	96180	96140	95910	95570	95190	94780	94360	93900	93450	93050	92800	92770	92950	93290
6	69850	70030	70280	70520	70690	70730	70620	70430	70240	70140	70130	70150	70090	69890	69600	69340	69230	69300	69490
5	61730	61750	61820	61910	61990	62010	61920	61760	61610	61610	61790	62080	62310	62060	61750	61560	61580	61710	
4	61830	61740	61650	61620	61670	61780	61820	61730	61560	61480	61660	62100	62580	62810	62660	62310	62040	62120	
3	70320	70190	69930	69710	69750	70510	70240	69880	69810	69490	69270	69280	69390	95260	96000	95960	95550	95210	95160
2	93840	93860	93340	92740	92650	93340	94390	94940	94510	93940	92700	92820	93900	95260	96000	95960	95550	95210	95160
1	17041	17171	17071	16861	16761	16891	17221	17471	17401	17101	16771	16611	16771	17191	17511	17741	17731	17591	17471
0	37802	29272	99832	83122	41002	27282	39582	72612	23672	23402	55052	26192	19672	34312	85092	32812	47072	14613	58472

	0	1	2	3	4	5	6	7	8	9	10	11	12	13	14	15	16	17	18
9	79012	82722	87562	93812	10193	11253	12653	14583	17343	21553	28533	42853	82933	20574	75213	40973	28123	21513	17513
8	17551	17671	17811	17951	18091	18221	18331	18401	18441	18431	18371	18211	18161	18031	17921	17841	17801	17801	17851
7	93290	94060	94060	94390	94600	95040	95450	95880	96240	96420	96360	95640	95120	94620	94220	93990	93980	94190	
6	69490	69640	69660	69550	69450	69510	69770	70170	70550	70780	70830	70720	70540	70350	70160	69980	69850	69800	69850
5	61710	61750	61560	61200	60900	60870	61130	61540	61880	61980	61870	61710	61640	61680	61760	61810	61800	61760	61730
4	62120	62120	62180	62200	62280	62390	62490	62500	62360	62120	61970	61240	60860	60800	61030	61380	61680	61840	61830
3	70970	70990	70690	70230	70200	70260	70820	71260	71070	70170	69670	68330	68200	68560	69130	69650	70020	70250	70320
2	95160	95250	95170	95100	95430	96140	96930	97190	96230	94190	92180	91000	90750	91070	91600	92120	92690	93340	93840
1	17471	17481	17591	17791	18051	18191	18211	18111	17781	17321	17051	17011	17001	16901	16731	16591	16601	16781	17041
0	58472	13713	82662	69770	56172	22192	19732	44852	34182	16122	16522	22422	34702	57072	48143	56712	34462	29942	37802

scaling:  $\times 10^3$

$\rho$  at 10,000 km elevation ( $\omega = 0$ )

	-18	-17	-16	-15	-14	-13	-12	-11	-10	-9	-8	-7	-6	-5	-4	-3	-2	-1	0
9	50643	33673	25163	20163	16903	14633	12983	11743	10783	10453	94512	90302	88532	84332	82162	81912	80512	80632	81402
8	19551	19571	19581	19601	19611	19601	19591	19561	19521	19471	19421	19361	19301	19251	19201	19171	19151	19171	
7	10341	10361	10371	10391	10401	10411	10401	10391	10381	10361	10341	10321	10301	10291	10271	10271	10261	10271	
6	76610	76720	76850	76970	77061	77130	77110	77100	77080	77010	76930	76770	76690	76530	76460	76460			
5	67290	67390	67490	67590	67680	67760	67760	67720	67690	67650	67630	67610	67580	67540	67490	67450	67410	67410	
4	67210	67300	67390	67490	67570	67640	67670	67660	67640	67610	67610	67620	67650	67670	67650	67610	67580	67560	
3	76300	76390	76490	76580	76680	76760	76810	76780	76760	76760	76760	76760	76760	76760	76760	76760	76760	76760	
2	10251	10261	10271	10291	10301	10311	10321	10321	10321	10311	10311	10331	10351	10371	10381	10391	10391	10381	
1	19151	19171	19181	19201	19221	19241	19261	19271	19271	19271	19271	19271	19271	19271	19271	19271	19271	19271	
0	93532	82512	80972	85192	76762	12863	1753	15273	11893	12073	15883	22183	27153	45993	68373	30733	26063	34443	89673

	0	1	2	3	4	5	6	7	8	9	10	11	12	13	14	15	16	17	18
9	81402	82822	84962	87912	91802	96822	10323	11153	12213	13603	15483	18133	21963	29113	39173	63383	12064	92773	50643
8	19171	19201	19251	19301	19361	19431	19481	19531	19571	19591	19601	19601	19591	19571	19561	19541	19541	19541	19551
7	10271	10281	10291	10301	10311	10331	10361	10381	10401	10401	10401	10401	10381	10371	10351	10341	10331	10331	10341
6	76460	76500	76560	76560	76770	76900	77020	77110	77150	77120	77040	76910	76770	76640	76550	76500	76490	76530	76610
5	67410	67420	67440	67490	67560	67650	67740	67800	67820	67710	67650	67630	67730	67720	67620	67610	67710	67720	67730
4	67560	67550	67550	67560	67610	67680	67760	67760	67720	67710	67650	67630	67730	67720	67620	67610	67600	67610	67610
3	76960	76920	76900	76930	77000	77080	77130	77100	77090	77060	77040	77020	77000	76930	76910	76910	76910	76910	76910
2	10381	10371	10361	10361	10371	10381	10381	10381	10361	10331	10301	10271	10271	10251	10241	10231	10241	10251	
1	19831	19801	19871	19841	19821	19821	19821	19821	19821	19821	19821	19821	19821	19821	19821	19821	19821	19821	
0	89673	89673	26283	13833	92322	76902	83142	13023	20653	97682	65432	57872	61572	77152	12053	30623	28403	12963	93532

scaling:  $\times 10^4$

$\rho$  at 100,000 km elevation ( $\omega = 0$ )

	-18	-17	-16	-15	-14	-13	-12	-11	-10	-9	-8	-7	-6	-5	-4	-3	-2	-1	0
9	50162	49932	49312	48432	47332	46082	44752	43412	42102	40862	39712	38722	37342	37132	36512	36072	35772	35622	35612
8	53240	53280	53320	53370	53420	53460	53470	53470	53450	53410	53350	53280	53210	53140	53080	53040	53010	53040	53040
7	28290	28310	28340	28370	28400	28420	28440	28440	28430	28420	28390	28360	28330	28300	28280	28260	28250	28250	28260
6	20990	21000	21030	21050	21070	21090	21100	21100	21090	21080	21060	21040	21020	21000	20990	20980	20980	20980	20990
5	18450	18470	18480	18500	18520	18540	18550	18560	18550	18550	18540	18520	18500	18490	18470	18460	18460	18460	18460
4	18450	18460	18480	18500	18520	18540	18550	18550	18550	18550	18550	18520	18510	18490	18480	18470	18460	18460	18470
3	20970	20990	21010	21030	21050	21070	21080	21090	21090	21070	21060	21050	21030	21020	21010	21000	21010	21010	21010
2	28250	28270	28290	28320	28350	28380	28400	28400	28400	28390	28370	28350	28340	28320	28310	28300	28300	28300	28300
1	53060	53090	53130	53180	53230	53270	53300	53330	53330	53330	53320	53320	53250	53230	53210	53200	53190	53190	53190
0	50182	53462	50612	50342	34172	44822	75902	24173	12623	61892	44372	37982	36622	39272	47572	69712	16723	29833	78452
-1	53110	53170	53260	53310	53380	53430	53470	53490	53490	53480	53450	53410	53360	53310	53260	53220	53190	53190	53210
-2	28250	28260	28310	28340	28380	28400	28420	28430	28430	28420	28380	28360	28330	28310	28290	28260	28300	28300	28300
-3	20970	20980	21010	21030	21050	21070	21080	21090	21080	21070	21060	21040	21030	21010	21000	21000	21000	21000	21000
-4	18440	18450	18470	18490	18510	18520	18540	18550	18550	18550	18550	18520	18510	18490	18480	18470	18460	18460	18460
-5	18440	18450	18470	18490	18510	18520	18530	18540	18540	18530	18520	18510	18500	18480	18470	18460	18450	18450	18450
-6	20970	20990	21000	21020	21040	21060	21080	21100	21110	21100	21090	21060	21040	21010	20990	20980	20970	20970	20970
-7	28270	28290	28310	28340	28360	28380	28400	28400	28400	28390	28380	28360	28340	28310	28290	28260	28240	28230	28230
-8	53200	53230	53260	53300	53340	53370	53390	53390	53370	53340	53290	53230	53170	53110	53060	53010	52990	52990	52990
-9	50782	50632	50102	49422	48112	46802	45382	43942	42532	41192	39962	38852	37932	37342	36422	35922	35542	35522	35522
0	0	1	2	3	4	5	6	7	8	9	10	11	12	13	14	15	16	17	18
9	35612	35742	36022	36442	36992	37682	38512	39462	40532	41702	42392	44242	45562	46932	47942	48912	49642	50072	50162
8	53040	53080	53130	53200	53270	53330	53380	53420	53430	53430	53400	53370	53320	53240	53210	53200	53210	53240	53240
7	28260	28280	28300	28340	28370	28400	28420	28430	28440	28430	28410	28380	28350	28330	28300	28280	28260	28290	28290
6	20990	21000	21020	21040	21060	21080	21100	21110	21110	21100	21090	21060	21040	21010	20990	20980	20970	20980	20990
5	18460	18470	18490	18510	18530	18540	18560	18560	18560	18550	18550	18520	18500	18480	18460	18440	18440	18450	18450
4	18470	18480	18490	18510	18530	18540	18550	18560	18560	18560	18550	18540	18520	18500	18480	18460	18440	18430	18450
3	21010	21010	21030	21050	21070	21080	21100	21110	21110	21100	21080	21060	21030	21000	20980	20970	20960	20960	20970
2	28300	28310	28330	28350	28380	28400	28420	28430	28430	28420	28390	28360	28330	28300	28270	28250	28230	28240	28250
1	53190	53200	53210	53240	53280	53320	53350	53370	53380	53370	53330	53280	53230	53170	53120	53080	53060	53050	53060
0	78452	47222	36272	32292	32622	37922	54262	12633	17183	57612	36492	29362	27582	27912	37552	61782	26243	11383	50182
-1	53210	53250	53300	53360	53430	53480	53520	53530	53520	53480	53410	53320	53240	53160	53100	53060	53050	53070	53110
-2	28300	28320	28340	28370	28400	28430	28450	28450	28450	28430	28400	28370	28330	28290	28260	28240	28230	28250	28250
-3	21000	21010	21030	21050	21070	21090	21100	21110	21110	21090	21080	21050	21020	21000	20980	20960	20950	20950	20970
-4	18460	18470	18480	18500	18520	18530	18550	18550	18550	18540	18540	18510	18490	18470	18450	18430	18430	18440	18440
-5	18450	18460	18470	18490	18510	18520	18540	18540	18540	18540	18540	18530	18510	18490	18470	18450	18440	18430	18440
-6	20970	20980	21000	21020	21040	21050	21070	21080	21080	21060	21050	21030	21000	20990	20970	20960	20960	20970	20970
-7	28230	28250	28270	28300	28320	28350	28370	28390	28390	28380	28360	28330	28310	28290	28270	28260	28260	28270	28270
-8	52990	53010	53060	53110	53170	53230	53280	53320	53340	53350	53340	53310	53280	53250	53210	53190	53180	53180	53200
-9	35252	35332	35552	35922	36432	37092	37892	38832	39902	41092	42382	43752	45152	46552	47862	49312	49932	50542	52782

scaling:  $\times 10^7$

Number Map 4.--Radius of torsion of the orthogonal trajectories  
(length unit: km). (See page 55 for explanation  
of first column.)

$\tau$  at sea level ( $\omega \neq 0$ )

-18	-17	-16	-15	-14	-13	-12	-11	-10	-9	-8	-7	-6	-5	-4	-3	-2	-1	0	
9	73E1	88E1	12E2	18E2	51E2	-55E2	-17E2	-10E2	-73E1	-58E1	-49E1	-44E1	-42E1	-44E1	-49E1	-56E1	-69E1	-90E1	
8	-73E2	-71E2	-66E2	-71E2	-91E2	-15E3	-71E3	27E3	12E3	87E2	75E2	74E2	83E2	11E3	24E3	-81E3	-15E3	-8AE2	-69E2
7	-99E2	-73E2	-69E2	-79E2	-11E3	-27E3	78E3	18E3	11E3	95E2	85E2	81E2	85E2	10E3	17E3	28E4	-17E3	-94E2	-75E2
6	-12E3	-84E2	-78E2	-95E2	-16E3	-14E4	25E3	15E3	13E3	14E3	15E3	13E3	10E3	98E2	12E3	37E3	-21E3	-94E2	-78E2
5	-19E3	-11E3	-96E2	-11E3	-24E3	70E3	16E3	13E3	13E3	39E3	25E5	98E3	18E3	97E2	89E2	16E3	-32E3	-92E2	-79E2
4	-44E3	-16E3	-12E3	-14E3	-27E3	69E3	15E3	11E3	15E3	24E4	-16E3	-15E3	63E4	10F3	65E2	92E2	-88E3	-83E2	-73E2
3	-65E4	-37E3	-24E3	-18E3	-19E3	-49E3	21E3	93E2	95E2	56E3	-98E2	-65E2	-13E3	51E2	59E2	64E3	-71E2	-61E2	
2	-23E4	90E3	53E3	-42E3	-10E3	-87E3	-46E3	78E2	48E2	89E2	-95E2	-38E2	-46E2	32E3	40E2	37E2	16E3	-57E2	-45E2
1	-10E3	17E3	56E2	13E3	-47E2	-25E2	-41E2	56E2	19E2	22E2	-30E3	-18E2	-17E2	-72E2	26E2	19E2	50E2	-37E2	-24E2
0	18E1	14E0	-13E1	-84E0	31E1	39E1	25E1	-71E1	-35E1	-36E1	-13E1	42E1	62E1	39E1	-18E1	-51E1	-37E1	58E1	47E1
-1	27E2	15E3	-37E2	-35E2	15E3	23E2	23E2	54E2	-20E2	-15E2	-39E2	31E2	16E2	25E2	-16E3	-27E2	-36E2	19E3	40E2
-2	41E2	10E3	-10E3	-61E3	61E2	51E2	29E3	-52E2	-74E2	85E2	37E2	47E2	43E3	-78E2	-78E2	-65E3	13E3		
-3	53E2	87E2	-69E2	-12E3	-31E3	16E3	11E3	57E3	-10E3	-67E2	-12E3	17E3	65E2	73E2	26E3	-19E3	-14E3	-38E3	69E3
-4	71E2	86E2	26E3	-45E3	-40E3	13E4	47E3	-94E3	-11E3	-20E3	43E3	12E3	26E3	-91E3	-33E3	-49E3	-77E3		
-5	10E3	97E2	15E3	65E3	-85E3	-53E3	-41E3	-24E3	-15E3	-14E3	-22E3	-32E4	30E3	22E3	84E3	-37E4	-90E3	-39E3	
-6	17E3	12E3	16E3	13E3	-13E4	-28E3	-17E3	-13E3	-11E3	-12E3	-16E3	-31E3	-55E4	53E5	40E3	47E3	89E3	-20E4	-33E3
-7	55E3	29E3	36E3	20E4	-37E3	16E3	-11E3	93E2	-86E2	-90E2	-10E3	-16E3	-25E3	-70E3	28E4	88E3	11E4	-23E4	-40E3
-8	-49E3	-30E3	-19E3	-13E3	-93E2	-72E2	-60E2	-54E2	-52E2	-53E2	-58E2	-67E2	-85E2	-11E3	-17E3	-27E3	-47E3	-86E3	-17E4
-9	-13E2	-77E1	-55E1	-42E1	-34E1	-28E1	-20E1	-17E1	-15E1	-13E1	-11E1	-89E0	-73E0	-58E0	-44E0	-32E0	-24E0	38E1	

scaling:  $\times 10$   
 $\tau$  at sea level ( $\omega = 0$ )

-18	-17	-16	-15	-14	-13	-12	-11	-10	-9	-8	-7	-6	-5	-4	-3	-2	-1	0	
9	-90E1	-13E2	-25E2	-17E3	38E2	18E2	12E2	92E1	77E1	67E1	61E1	57E1	54F1	53F1	54E1	55E1	58E1	64E1	73E1
8	-69E2	-63E2	-64E2	-73E2	-93E2	-14E3	-40E3	45E3	14E3	90E2	69E2	60E2	59E2	65E2	86E2	15E3	23E4	-17E3	-93E2
7	-75E2	-73E2	-82E2	-10E3	-61E3	-16E3	-22E3	-47E3	91E3	11E3	81E3	71E2	72E2	87E2	13E3	69E3	-21E3	-99E2	
6	-78E2	-92E2	-14E3	-27E3	-33E3	-21E3	-14E3	-31E3	-21E3	-40E4	21E3	11E3	91E2	84E2	92E2	12E3	30E3	-39E3	-12E3
5	-79E2	-13E3	-86E3	22E3	37E3	-26E3	-98E2	-82E2	-11E3	-37E3	37E3	16E3	13E3	12E3	13E3	21E3	10E5	-19E3	
4	-73E2	-20E3	13E3	77E2	13E3	-19E3	-63E2	-58E2	-99E2	-10E3	-25E4	18E3	17E3	32E3	55E3	31E3	19E3	44E3	-44E3
3	-61E2	-33E3	68E2	47E2	97E2	-99E2	-41E2	-46E2	-15E3	10E3	68E2	61E3	63E3	12E3	13E3	21E3	10E5	-19E3	
2	-45E2	-31E3	45E2	34E2	10E3	-47E2	-26E2	-38E2	27E3	33E2	29E2	63E2	-14E3	-57E2	-74E2	-30E3	24E3	34E3	-65E4
1	-24E2	-11E3	27E2	22E2	21E3	-19E2	-13E2	-26E2	32E2	12E2	12E2	28E2	-62E2	-21E2	-36E2	-86E2	-11E3	-75E2	-10E3
0	47E1	16E1	-26E1	-54E1	27E1	36E1	41E1	17E1	-32E1	-79E1	-68E1	-30E1	-29E1	-66E1	12E1	-10E2	19E2	38E1	18E1
-1	40E2	73E2	-97E2	-60E2	14E3	26E2	23E2	77E2	-74E2	-18E2	-23E2	-90E2	16E3	-16E3	-42E2	-48E2	14E3	28E2	27E2
-2	13E3	20E3	-25E3	-12E3	-36E3	14E3	91E2	16E3	-66E3	-33E3	33E3	13E3	85E4	-59E2	-34E2	-42E2	51E2	41E2	
-3	69E3	10E5	-18E3	-10E3	-10E3	-24E3	51E3	14E3	86E2	60E2	51E2	71E2	-12E4	-51E2	-33E2	-39E2	-14E3	83E2	53E2
-4	-77E3	-29E3	-12E3	-23E2	-78E2	-11E3	-11E4	11E3	54E2	39E2	61E2	-43E5	-58E2	-37E2	-43E2	-10E3	16E3	71E2	
-5	-39E3	-17E3	10E5	-75E2	-74E2	-11E3	-31E4	92E2	47E2	37E2	39E2	62E2	62E3	-79E2	-48E2	-52E2	-10E3	47E3	10E3
-6	-13E3	-11E3	-10E3	-82E3	-89E2	16E3	41E3	78E2	46E2	42E2	64E2	64E2	62E2	-73E2	-73E2	-12E3	-24E4	17E3	
-7	-40E3	-20E3	-15E3	-14E3	-20E3	-98E4	14E3	70F2	50F2	44E2	48E2	66E2	44E2	51E2	67E2	10E3	17E3	41E3	55E3
-8	-17E4	97E4	74E3	28E3	14E3	88E2	62E2	49E2	42E2	41E2	44E2	11E2	22E2	11E3	18E4	-34E3	-18E3	-14E2	-13E2
-9	38E1	13E0	41E0	55E0	72E0	91E0	11E1	14E1	20E1	21E1	30E1	38E1	51E1	73E1	12E2	37E2	-39E2	-13E2	

-18	-17	-16	-15	-14	-13	-12	-11	-10	-9	-8	-7	-6	-5	-4	-3	-2	-1	0	
9	74E1	12E2	38E2	-19E2	-61E1	-31E1	-18E1	-11E1	-67E0	-48E0	-49E0	-86E0	-27E1	14E2	33E1	24E1	20E1	19E1	19E1
8	72E2	70E2	79E2	10E3	21E3	-14E4	-15E3	-83E2	-58E2	-42E2	-41E2	-45E2	-55E2	-83E2	-20E3	34E3	96E2	59E2	
7	71E2	75E2	98E2	13E3	32E3	12E4	-11E4	-36E3	-18E3	-11E3	-81E2	-66E2	-62E2	-67E2	-86E2	-15E3	75E4	14E3	78E2
6	71E2	92E2	21E3	-22E4	-47E3	-19E4	43E3	26E3	36E3	36E3	-20E3	-11E3	-91E2	-85E2	-92E2	-12E3	-33E3	32E3	11E3
5	62E2	11E3	-43E3	-10E3	-12E3	-10E4	14E3	92E2	11E3	27E3	-66E3	-22E3	-16E3	-13E3	-11E3	-11E3	-11E4	19E3	
4	55E2	17E3	-92E2	-52E2	-73E2	29E4	75E2	58E2	89E2	50E3	-33E3	-44E3	30E4	24E4	-33E3	-14E3	-12E3	-22E3	10E4
3	44E2	24E3	-49E2	-33E2	-55E2	16E3	40E2	40E2	11E3	-12E3	-89E2	-34E3	13E3	95E2	19E3	-29E3	-12E3	-15E3	-42E3
2	32E2	21E3	-32E2	-23E2	-50E2	47E2	21E2	29E2	-25E3	-12E3	-34E2	-14E3	61E2	41E2	64E2	66E3	-17E3	-20E3	-32E3
1	22E2	99E2	-25E2	-17E2	-56E2	20E2	13E2	25E2	-32E2	-14E2	-18E2	-11E3	29E2	23E2	42E2	32E3	-58E3	25E3	43E3
0	47E1	16E1	-26E1	-54E1	26E1	36E1	41E1	17E1	-32E1	-78E1	-68E1	-30E1	-28E1	-64E1	12E1	-10E2	19E2	38E1	18E1
-1	-30E2	-49E2	95E2	47E2	-15E4	-29E2	-22E2	-10E3	23E2	18E2	40E2	-99E2	-44E2	-12E3	55E2	42E2	-15E3	-24E2	-25E2
-2	-78E2	-98E2	38E3	94E2	17E3	-17E3	-92E2	-30E3	15E3	23E3	-12E3	-64E2	-12E3	81E2	33E2	36E2	69E3	-37E2	-32E2
-3	-19E3	-23E3	36E3	10E3	87E2	12E3	32E3	-42E3	-11E3	-61E2	-47E2	-56E2	-30E3	62E2	35E2	42E2	20E3	-67E2	-47E2
-4	-45E3	-59E4	16E3	77E2	61E2	72E2	15E3	-20E3	-59E2	-38E2	-35E2	-50E2	-73E3	54E2	35E2	42E2	13E3	-10E3	-57E2
-5	-94E3	41E3	12E3	70E2	60E2	72E2	17E3	-16E3	-54E2	-37E2	-36E2	-54E2	-84E3	64E2	42E2	49E2	12E3	-20E3	-78E2
-6	-57E3	52E3	15E3	95E2	85E2	10E3	29E3	-20E3	-76E2	-34E2	-34E2	-83E2	-88E3	10E3	70E2	79E2	16E3	-52E3	-13E3
-7	-35E3	19E3	25E3	15E3	23E3	-37E4	-16E3	-88E2	-69E2	-72E2	-10E3	-41E3	23E3	12E3	21E3	67E4	-26E3		
-8	-15E3	-18E3	-23E3	-27E3	-24E3	-17E3	-12E3	-92E2	-77E2	-71E2	-76E2	-96E2	-15E3	-43E3	98E3	34E3	30E3	35E3	42E3
-9	-97E1	-91E1	-87E1	-82E1	-78E1	-73E1	-69E1	-62E1	-59E1	-54E1	-52E1	-50E1	-49E1	-51E1	-54E1	-59E1	-10E2	-11E2	-10E2

scaling:  $\times 10$

$\tau$  at 1,000 km elevation ( $\omega = 0$ )

-18	-17	-16	-15	-14	-13	-12	-11	-10	-9	-8	-7	-6	-5	-4	-3	-2	-1	0	
9	30E1	33E1	37E1	40E1	44E1	48E1	51E1	55E1	60E1	64E1	67E1	73E1	78E1	84E1	90E1	97E1	11E2	12E2	13E2
8	14E3	11E3	9E3	10E3	11E3	16E3	32E3	64E4	-42E3	-23E3	-19E3	-19E3	-22E3	-35E3	-15E4	54E3	23E3	16E3	13E3
7	17E3	12E3	11E3	13E3	18E3	35E3	92E4	-47E3	-27E3	-22E3	-19E3	-18E3	-19E3	-25E3	-49E3	14E4	27E3	16E3	14E3
6	24E3	19E3	14E3	17E3	29E3	18E4	-54E3	-32E3	-32E3	-38E3	-41E3	-33E3	-24E3	-21E3	-28E3	-15E4	32E3	16E3	14E3
5	43E3	21E3	18E3	23E3	51E3	10E4	-31E3	-26E3	-31E4	81E3	16E4	-45E3	-19E3	-17E3	-35E3	41E3	14E3	13E3	12E3
4	36E4	40E3	29E3	32E3	69E3	-89E3	-27E3	-21E3	-34E3	93E3	21E3	19E3	12E4	-10E3	-11E3	-17E3	74E3	13E3	12E3
3	-72E3	85E5	10E4	56E3	46E3	12E4	-30E3	-26E3	-18E3	10E5	14E3	10E3	18E3	-22E3	-85E2	-10E3	-49E4	10E3	97E2
2	-52E3	-33E3	-29E3	-14E4	24E3	16E3	62E3	-14E3	-88E2	-18E3	14E3	63E2	79E2	-62E3	-65E2	-63E2	-34E3	86E2	71E2
1	11E5	-12E3	-70E2	-12E3	13E3	52E2	72E2	-12E3	-37E2	-46E2	38E2	35E2	15E3	-47E2	-34E2	-10E3	59E2	42E2	0
0	77E1	-23E2	-87E0	-13E1	32E2	79E1	37E1	-37E1	-59E1	-60E1	93E1	15E2	15E2	-33E1	-15E2	-11E2	-29E1	45E2	-1
-1	-52E2	14E5	48E2	48E2	-25E4	-40E2	-36E2	-29E3	30E2	29E2	72E2	-55E2	-28E2	-44E2	20E3	48E2	71E2	-20E3	-61E2
-2	-72E2	-25E5	12E3	90E2	39E3	-10E3	-79E2	-31E3	95E2	64E2	13E3	-14E3	-63E2	-80E2	-13E4	12E3	14E3	-92E4	-17E3
-3	-96E2	-18E3	44E3	17E3	35E3	-32E3	-17E3	-52E3	21E3	13E3	24E3	-32E3	-12E3	-13E3	-54E3	32E3	27E3	16E4	-51E3
-4	-12E3	-16E3	-72E3	46E3	49E3	-35E4	-62E3	-13E5	34E3	22E3	40E3	-67E3	-21E3	-46E3	20E4	80E3	29E4	-36E4	-5
-5	-18E3	-17E3	-30E3	-20E4	10E4	97E3	97E3	55E3	33E3	30E3	47E3	-54E4	-65E3	-34E3	-44E3	-84E3	-7E4	-35E4	19E4
-6	-30E3	-23E3	-29E3	-63E3	38E4	62E3	39E3	30E3	27E3	28E3	39E3	10E4	-17E4	-59E3	-46E3	-47E3	-56E3	-10E4	19E4
-7	-80E3	-45E3	-51E3	-12E4	11E4	39E3	25E3	20E3	19E3	21E3	27E3	44E3	14E4	-13E4	-57E3	-45E3	-48E3	-80E3	75E4
-8	77E3	63E3	43E3	29E3	20E3	15E3	13E3	12E3	13E3	15E3	20E3	31E3	69E3	88E5	-10E4	-60E3	-67E3	-85E3	-9
-9	-10E2	-13E2	-17E2	-23E2	-34E2	-50E2	-70E2	-88E2	-93E2	-85E2	-7E2	-60E2	-50E2	-42E2	-35E2	-31E2	-27E2	-24E2	-22E

0	1	2	3	4	5	6	7	8	9	10	11	12	13	14	15	16	17	18		
9	13E2	15E2	17E2	23E2	35E2	11E3	-61E2	-20E2	-10E2	-57E1	-32E1	-15E2	-31E3	15E1	18E1	22E1	26E1	35E1		
8	13E3	13E3	14E3	17E3	28E3	92E3	-61E3	-22E3	-14E3	-10E3	-91E2	-85E2	-88E2	-10E3	-14E3	-26E3	-14E5	27E3	14E3	
7	14E3	14E3	18E3	26E3	42E3	72E3	17E4	-22E4	51E3	-25E3	-17E3	-13E3	-12E3	-12E3	-15E3	-25E3	-12E4	37E3	17E3	
6	14E3	18E3	38E3	34E4	-31E4	15E4	47E3	36E3	51E3	-25E5	-40E3	-22E3	-17E3	-15E3	-16E3	-22E3	-50E3	90E3	24E3	
5	13E3	25E3	-90E3	-23E3	-29E3	37E4	23E3	16E3	20E3	51E3	-12E4	-40E3	-31E3	-26E3	-23E3	-23E3	-32E3	-15E4	43E3	
4	12E3	41E3	-18E3	-10E3	-15E3	10E4	13E3	10E3	17E3	10E4	-54E3	-62E3	-48E4	85E4	-78E3	-32E3	-27E3	-45E3	36E4	
3	97E2	68E3	-10E3	-69E2	-12E3	26E3	77E2	79E2	23E3	-23E3	-15E3	-43E3	18E3	31E3	-12E4	-31E3	-33E3	-72E3	-	
2	71E2	60E3	-67E2	-49E2	-11E3	94E2	45E2	62E2	-52E2	-63E2	-64E2	-21E3	13E3	82E2	11E3	43E3	-65E3	-50E3	-52E3	
1	42E2	20E3	-46E2	-33E2	-13E3	36E2	23E2	46E2	-58E2	-24E2	-28E2	-10E3	62E2	41E2	62E2	10E3	52E3	30E3	11E4	
0	45E2	53E1	-58E1	-13E2	20E1	52E1	66E1	45E1	-61E1	-19E2	-16E2	-53E1	-22E1	-19E1	10E1	-32E2	-34E2	20E2	77E1	
-1	-61E2	-10E3	16E3	87E2	-26E4	-54E2	-41E2	-15E3	49E2	34E2	61E2	-43E3	-97E2	-29E3	11E3	95E2	-24E3	-48E2	-52E2	-
-2	-17E3	-24E3	48E3	17E3	31E3	-38E3	-18E3	-39E3	52E3	64E3	-29E3	-14E3	-31E3	15E3	69E2	78E2	10E4	-86E2	-72E2	-
-3	-51E3	-91E3	38E3	16E3	15E3	24E3	16E4	-37E3	-16E3	-10E3	-87E2	-11E3	-84E3	11E3	66E2	77E2	30E3	-14E3	-96E2	-
-4	-36E4	91E3	23E3	13E3	11E3	15E3	45E3	-27E3	-10E3	-70E2	-67E2	-10E3	-14E4	10E3	69E2	81E2	21E3	-26E3	-12E3	-
-5	19E4	40E3	18E3	12E3	11E3	15E3	52E3	-22E3	-94E2	-68E2	-68E2	-10E3	-96E3	13E3	85E2	96E2	20E3	-64E3	-18E3	-
-6	19E4	38E3	20E3	15E3	14E3	21E3	12E4	-23E3	-11E3	-83E2	-86E2	-12E3	-61E3	23E3	13E3	14E3	25E3	-67E3	-30E3	-
-7	75E4	55E3	30E3	24E3	27E3	53E3	-91E3	-21E3	-13E3	-10E3	-11E3	-15E3	-37E3	89E3	28E3	36E3	36E3	15E4	-80E3	-
-8	-85E3	-12E4	-17E4	-11E4	-58E3	-31E3	-19E3	-14E3	-11E3	-10E3	-13E3	-17E3	-31E3	-88E3	34E4	10E4	82E3	77E3	-	-
-9	-22E2	-20E2	-18E2	-16E2	-15E2	-14E2	-13E2	-12E2	-11E2	-10E2	-95E1	-89E1	-81E1	-79E1	-79E1	-82E1	-90E1	-10E1	-	-

scaling:  $\times 10$

$\tau$  at 10,000 km elevation ( $\omega = 0$ )

omitted, because some of the results were already affected by rounding errors of the computer.

$\tau$  at 100,000 km elevation ( $\omega = 0$ )

omitted, because some of the results were already affected by rounding errors of the computer.

scaling:

Number Map 5.--Orthogonal trajectories, coefficients; geocentric radius.  
See p. 59 for explanation of first column.

$r_q$

-18	-17	-16	-15	-14	-13	-12	-11	-10	-9
56792513522 56792523858 56792529796 56792540178 56792551705 56792564043 56792576829 56792589682 56792602214 56792614045	2550000000 2550000000 2550000000 2550000000 2500000000 2500000000 2500000000 2500000000 2500000000 2500000000								
8 57429653382 57429733168 574298891618 5742988946510 57429897314 57429855094 57429807055 57429889135 57431263515 5743284754	249998375 249998376 249998377 249998378 249998379 249998380 249998381 249998382 249998383 249998384								
-1 5742976231334 592739344254 59271115725 59268464168 59266416646 59265159826 59266044962 59267930351 59271000673 59274974663	-1 2499984235 2499984236 2499984237 2499984238 2499984239 2499984240 2499984241 2499984242 2499984243 2499984244								
5 62112933223 62117333431 62117152241 62095768872 62091275323 62091275324 62091275325 62091275326 62091275327 62091275328	2499984466 2499984467 2499984468 2499984469 2499984470 2499984471 2499984472 2499984473 2499984474 2499984475								
-5 65597381512 65597301878 65583874976 65571818167 65565042644 655619204574 65563050629 65567854217 65574719175 6558188545	2499985342 2499985352 2499985362 2499985372 2499985382 2499985392 2499985402 2499985412 2499985422 2499985432								
4 692116521 692117143 69211721843 69211721844 69211721845 69211721846 69211721847 69211721848 69211721849 69211721850	2499986173 2499986174 2499986175 2499986176 2499986177 2499986178 2499986179 2499986180 2499986181 2499986182								
3 72765921337 72771715533 72776471198 727755281937 727764427902 7277557330563 72761938157 72772924441 7278304880	2499987381 2499987382 2499987383 2499987384 2499987385 2499987386 2499987387 2499987388 2499987389 2499987390								
2 75613616447 75553233136 75543521867 75635247179 75632153762 75632122979 75609355432 75622511615 75636666361	2499987412 2499987413 2499987414 2499987415 2499987416 2499987417 2499987418 2499987419 2499987420 2499987421								
1 77516637593 77525285619 775177125252 774752528491 77475374149 77473321812 77474691207 77489139643 7750787446	2499987507 2499987508 2499987509 2499987510 2499987511 2499987512 2499987513 2499987514 2499987515 2499987516								
-1 78146117113 78146125445 78176296811 78169223529 78159117463 78143742762 78139145717 78128564131 78146267637 7816188654	2500000000 2500000000 2500000000 2500000000 2500000000 2500000000 2500000000 2500000000 2500000000 2500000000								
-1 77553451554 77533275729 77533224149 77525191497 77512707479 77504466463 77491322365 77490138052 77520128878	2499987643 2499987644 2499987645 2499987646 2499987647 2499987648 2499987649 2499987650 2499987651 2499987652								
-1 75561194026 75575343524 75566152770 7566274661 75654713529 7564472773 7563447703 756311955056 75641648424 75659142835	2499987712 2499987713 2499987714 2499987715 2499987716 2499987717 2499987718 2499987719 2499987720 2499987721								
-2 724171219 724177413 724177414 724177415 724177416 724177417 724177418 724177419 724177420 724177421	2499987722 2499987723 2499987724 2499987725 2499987726 2499987727 2499987728 2499987729 2499987730 2499987731								
-3 72323225397 72319381436 72323764132 72321252393 72326324998 72326246111 72324787209 72324951111 72292252107 72805291172	2499987731 2499987732 2499987733 2499987734 2499987735 2499987736 2499987737 2499987738 2499987739 2499987740								
-4 5433212171 54312235739 54312259111 54325621256 543291961116 54329058989 54327314789 54329048107 69296753623 69302020701	2499987741 2499987742 2499987743 2499987744 2499987745 2499987746 2499987747 2499987748 2499987749 2499987750								
-5 55592274511 55535533912 55535533913 55535533914 55535533915 55535533916 55535533917 55535533918 55535533919 55535533920	2499987751 2499987752 2499987753 2499987754 2499987755 2499987756 2499987757 2499987758 2499987759 2499987760								
-6 55592274511 55535533912 55535533913 55535533914 55535533915 55535533916 55535533917 55535533918 55535533919 55535533920	2499987761 2499987762 2499987763 2499987764 2499987765 2499987766 2499987767 2499987768 2499987769 2499987770								
-7 57316221219 57345121113 57345116561 57345116562 57345116563 57345116564 57345116565 57345116566 57345116567 57345116568	2499987771 2499987772 2499987773 2499987774 2499987775 2499987776 2499987777 2499987778 2499987779 2499987780								
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# Coefficients, geocentric latitude

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# Coefficients, geocentric longitude

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	143	232	145	31	-41	-49	-29	-20	-28	-36
	1	1	1	1	1	1	1	1	1	1
1	-111,32	-131439	-109534	-51234	-15293	1301	10176	23832	38215	42198
	23	255	154	29	-40	-45	-32	-40	-58	-55
	1	1	1	1	1	1	1	1	1	1
0	-105457	-172534	-92913	-54786	-30786	-19592	-4576	21621	46493	52015
	23	232	137	38	-2	4	-1	-42	-88	-85
	1	1	1	1	1	1	1	1	1	1
-1	-713x3	-83945	-71152	-57713	-55461	-53065	-31497	9414	48940	61495
	153	147	93	51	62	85	60	-22	-104	-117
	1	1	1	1	1	1	1	1	1	1
-2	-13253	-27356	-37525	-56292	-80592	-89827	-65989	-11012	43137	66662
	32	19	17	58	131	177	136	18	-101	-139
	1	1	1	1	1	1	1	1	1	1
-3	55435	36972	2619	-47362	-96646	-118346	-94799	-34757	29391	64652
	-134	-124	-64	91	181	249	225	68	-78	-145
	1	1	1	1	1	1	1	1	1	1
-4	119771	76351	43244	-28919	-95796	-128507	-111699	-54979	10890	54575
	-272	-247	-142	27	19	280	245	111	-41	-130
	1	1	1	1	1	1	1	1	1	1
-5	169143	142549	81222	143	-73616	-114306	-108385	-64589	-7089	37770
	-166	-129	-222	-16	158	255	240	134	-97	-97
	1	1	1	1	1	1	1	1	1	1
-6	2,323	177352	118185	42768	-28276	-72765	-81017	-57925	-18349	17536
	-411	-377	-740	-92	76	13	185	123	27	-53
	1	1	1	1	1	1	1	1	1	1
-7	239371	218327	172411	107658	46229	-411	-24617	-27113	-15664	-1129
	-442	-427	-313	-14	-57	35	76	69	32	-8
	1	1	1	1	1	1	1	1	1	1
-8	363382	349475	310315	259348	231929	145507	95563	54306	20762	-8171
	-5	-5	-51	-51	-427	-315	-214	-129	-63	28
	1	1	1	1	1	1	1	1	1	1
-9	23,54721	23054911	22359493	20078494	18958693	10362881	13271672	9780312	5995818	2033617
	-37	53	-367	72	-32676	-29119	-24675	-19477	20	-1052
	7	7	68	62	53	44	32	71	7	-6

SUMMARY (Appendices A and B)  
Changes of the Correction Terms with Height

Equipotential surfaces						
elevation above sea level correction terms	0 km	1,000 km	10,000 km	100,000 km		
$\Delta r(m)$	60 -70	47 -51	16 -14	1.8	-1.8	
$\Delta g(\text{mgal})$	22 -26	8 -10	0.18 -0.17	0.000066	-0.000062	
$\xi''$	5.7 -5.4	3.1 -2.9	0.25 -0.29	0.0039	-0.0038	
$\eta''$	5.7 -6.8	2.9 -3.9	0.34 -0.42	0.0070	-0.0074	
$\Delta \frac{1}{\sqrt{K}}, \Delta \frac{1}{H}(m)$	567 -470	330 -270	43 -45	3.8	-4.1	
Equigravitational surfaces						
$\Delta r(m)$	116 -158	85 -104	26 -24	2.9	-2.7	
$\xi''$	15 -15	7.8 -7.7	0.43 -0.52	0.0061	-0.0060	
$\eta''$	16 -17	8.0 -9.1	0.54 -0.72	0.011	-0.011	
$\Delta \mu'' (\omega = 0)$	9.8 -9.6	4.9 -4.7	0.23 -0.20	0.0022	-0.0019	
Orthogonal trajectories						
$\Delta \delta''$	5.5 -5.8	3.0 -3.1	0.29 -0.23	0.0039	-0.0037	

APPENDIX C. Constants and Coefficients Used in the Numerical  
Computations

Constants and coefficients used in the numerical computations

$$a = 6,378,165 \text{ m equatorial radius}$$

$$GM = 3.986\ 032 \times 10^{20} \text{ cm}^3 \text{ sec}^{-2} \text{ gravitational constant} \times \text{mass of the Earth}$$

$$\omega^2 = 5.317\ 49 \times 10^{-9} \text{ rad sec}^{-2} \text{ (gravity field)}$$

and/or

$$\omega^2 = 0 \text{ (gravitational field)}$$

Zonal harmonic coefficients (Kozai, 1964)

$$\begin{aligned} c_{20} &= -1,082.645 \times 10^{-6} \\ c_{40} &= 1.649 \times 10^{-6} \\ c_{60} &= -0.646 \times 10^{-6} \\ c_{80} &= 0.270 \times 10^{-6} \\ c_{100} &= 0.054 \times 10^{-6} \\ c_{120} &= 0.357 \times 10^{-6} \\ c_{140} &= -0.179 \times 10^{-6} \end{aligned}$$

$$\text{with } c_{n0} = -J_n$$

$$\begin{aligned} c_{30} &= 2.546 \times 10^{-6} \\ c_{50} &= 0.210 \times 10^{-6} \\ c_{70} &= 0.333 \times 10^{-6} \\ c_{90} &= 0.053 \times 10^{-6} \\ c_{110} &= -0.302 \times 10^{-6} \\ c_{130} &= 0.114 \times 10^{-6} \end{aligned}$$

Conventional tesseral (sectorial) harmonic coefficients (Izsak, 1965)<sup>8</sup>

$$\begin{aligned} c_{22} &= 1.3446 \times 10^{-6} \\ c_{31} &= 1.7304 \times 10^{-6} \\ c_{32} &= 0.1298 \times 10^{-6} \\ c_{33} &= -0.2416 \times 10^{-7} \\ c_{41} &= -0.3617 \times 10^{-6} \\ c_{42} &= 0.4387 \times 10^{-7} \\ c_{43} &= 0.4103 \times 10^{-7} \\ c_{44} &= -0.2276 \times 10^{-8} \end{aligned}$$

$$\begin{aligned} s_{22} &= -8.0817 \times 10^{-7} \\ s_{31} &= -0.4158 \times 10^{-7} \\ s_{32} &= -0.2749 \times 10^{-6} \\ s_{33} &= 0.1959 \times 10^{-6} \\ s_{41} &= -0.3831 \times 10^{-6} \\ s_{42} &= 0.1292 \times 10^{-6} \\ s_{43} &= -0.6018 \times 10^{-8} \\ s_{44} &= 0.9117 \times 10^{-8} \end{aligned}$$

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<sup>8</sup>As obtained from a computer tape; the last two digits of a coefficient are, however, guarding figures.

$$\begin{aligned}
c_{51} &= -0.1177 \times 10^{-6} \\
c_{52} &= 0.3826 \times 10^{-7} \\
c_{53} &= -0.2204 \times 10^{-7} \\
c_{54} &= -0.1036 \times 10^{-8} \\
c_{55} &= 0.1907 \times 10^{-9} \\
c_{61} &= -0.1734 \times 10^{-7} \\
c_{62} &= 0.6073 \times 10^{-8} \\
c_{63} &= 0.1102 \times 10^{-8} \\
c_{64} &= 0.2782 \times 10^{-9} \\
c_{65} &= 0.3098 \times 10^{-9} \\
c_{66} &= -0.2696 \times 10^{-10}
\end{aligned}$$

$$\begin{aligned}
s_{51} &= -0.3355 \times 10^{-7} \\
s_{52} &= -0.4381 \times 10^{-7} \\
s_{53} &= 0.1538 \times 10^{-8} \\
s_{54} &= 0.1282 \times 10^{-8} \\
s_{55} &= -0.1022 \times 10^{-8} \\
s_{61} &= 0.9182 \times 10^{-7} \\
s_{62} &= -0.2857 \times 10^{-7} \\
s_{63} &= 0.5474 \times 10^{-12} \\
s_{64} &= -0.1459 \times 10^{-8} \\
s_{65} &= -0.3098 \times 10^{-9} \\
s_{66} &= -0.1373 \times 10^{-9}
\end{aligned}$$

The  $C_{21}$  and  $S_{21}$  coefficients were assumed to be zero.

### Coordinate system

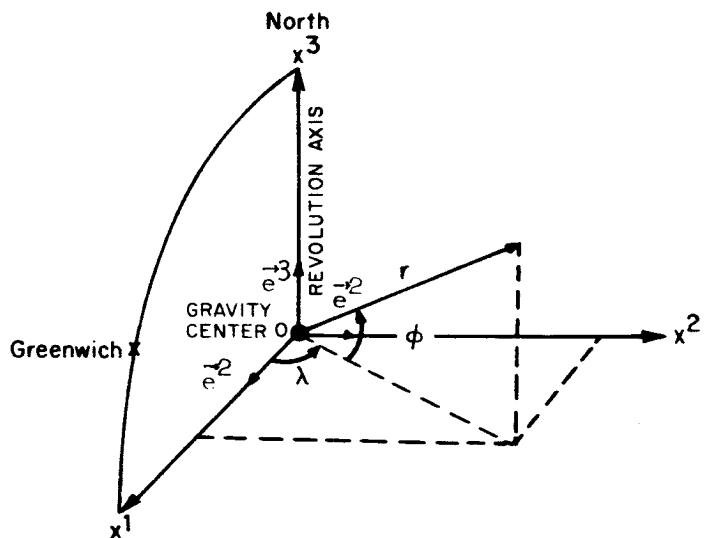


Figure 21.

Origin 0 is in the gravity center of the Earth ( $C_{10}=C_{11}=S_{11}=0$ )

$\varphi \gtrless 0$  Northern } Hemisphere  
Southern }

$\lambda \gtrless 0$  eastward  
westward

$\vec{e}^1, \vec{e}^2, \vec{e}^3$ , unit vectors

## NOTICE

This series of Special Reports was instituted under the supervision of Dr. F. L. Whipple, Director of the Astrophysical Observatory of the Smithsonian Institution, shortly after the launching of the first artificial earth satellite on October 4, 1957. Contributions come from the Staff of the Observatory.

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